RIASSUNTO

Modellizzazione geometrica a cinematica da dati geologici e geodetici delle strutture compressive nell’area Montello-Cansiglio (Catena Sudalpina Orientale, Italia nord-orientale).

Questo lavoro è dedicato allo studio delle geometrie e dei ratei di deformazione di breve e medio termine delle strutture compressive attive facenti parte dei fronti esterni della Catena Sudalpina, nel settore dell’anticlinale del Montello. Il metodo adottato utilizza informazioni derivate dall’analisi di una linea geodetica di primo ordine dell’IGM, combinate con osservazioni geofisiche, geologiche e geomorfologiche di superficie e di sottosuolo. La linea geodetica presa in esame mostra lungo alcuni suoi segmenti dei movimenti verticali relativi, positivi rispetto ai segmenti adiacenti (maggiori sollevamenti). Questi segnali geodetici, ottenuti dal confronto delle quote dei capisaldi misurate durante due distinte campagne separate da un intervallo di tempo di circa 50 anni, avvengono in corrispondenza dell’attraversamento di faglie cieche e sono stati quindi interpretati come dovuti all’attività di queste strutture sepolte. Per l’interpretazione, è stata costruita una sezione geologica che segue la traccia della linea di livellazione, ed è stato quindi modellizzato il segnale geodetico adottando un metodo diretto. Nel modello, le geometrie di partenza delle faglie sono state prese dalla sezione geologica, e sono state poi modificate per riprodurre il segnale geodetico. Una volta fissate le geometrie delle faglie, gli uplift rate sono stati convertiti in slip e shortening rate e comparati con: 1- i ratei di medio e lungo termine derivati dalle osservazioni geologiche e geomorfologiche per evidenziare eventuali cambiamenti nel tempo; e 2- con i tassi di convergenza GPS per studiare la partizione delle deformazione tra i diversi fronti. Infine sono state usate relazioni analitiche ed empiriche per stimare la massima magnitudo e i tempi di ricorrenza dei potenziali futuri terremoti.

Key words: Montello anticline, slip rates, Eastern Southalpine Chain, NE Italy.

INTRODUCTION

We present a study of the external thrust fronts of the Eastern Southalpine Chain (ESC) in the Montello-Cansiglio area (Fig. 1) using a combination of surface and subsurface geologic, morphotectonic and geodetic data. The aim of this work is to constrain the geometry of the active thrusts and to compute the corresponding rates of deformation. In addition we present an evaluation of the seismic potential of the seismogenic sources estimating the maximum magnitude of the potential earthquakes associated to the individual structures and their recurrence interval.

The active tectonics of the study area is the result of the relative motion of Africa and its northern Adriatic promontory with respect to Europe (e.g. CASTELLARIN, 2004). The analysis of the GPS velocities predicts a counter clockwise motion of the Adriatic block around a pole located in the western Alps (e.g. SERPELLONI et alii, 2005). This motion produces increasing convergence from west to east, which is matched by a similar increase of the seismic moment release. According to this model, in the Veneto Region N–S to NNW–SSE convergence is inferred to be in the order of about 1.5 mm/a (e.g. D’AGOSTINO et alii, 2005; GRENERCZY et alii, 2005).

The Neogene-Quaternary ESC is part of the S-verging backthrust chain of the Alps. West of the Tagliamento River, the ESC structures follow a WSW–ENE trend, and the activity of the main thrust fronts steadily migrated southwards (DOGGLIONI, 1992; CASTELLARIN & CANTELLI, 2000). The more external contractional structures are the Bassano–Valdobbiadene Thrust (BV Thrust in Fig. 1), associated with the uplift of the mountain front (DOGGLIONI, 1992), and in a
more external position, a system of growing mainly blind thrusts running at the boundary between the Prealpine relief and the plain areas (Fig. 1). This system is composed of several fault segments identified by GALADINI et alii (2005), which produce the uplift of Neogene–Quaternary deposits and along its eastern portion border the mountain front of the Carnian Prealps. The Montello-Conegliano Thrust (MC Thrust in Fig. 1) is the leading thrust of the ESC in the study sector. Its activity is expressed by the uplift and warping of Upper Miocene and Upper Pliocene Pleistocene deposits, forming the exposed Montello anticline (FERRARESE et alii, 1998; BENEDETTI et alii, 2000; FANTONI et alii, 2001; FANTONI et alii, 2002) one of the most impressive folds emerging from the Venetian and Friulian Plain. Surface and subsurface geological, geophysical and structural data show that the MC Thrust is a 35 km long fault that dies out to the east where it is overridden by the neighbouring Cansiglio Thrust (CA Thrust; GALADINI et alii, 2005). To the west the MC Thrust has a right-stepping echelon relationship with the adjoining Bassano-Cornuda Thrust (BC Thrust in Fig. 1). To the north of the MC Thrust the eastern portion of the BV Thrust is found at the base of the mountain front (Fig. 1). The area comprised between the two thrust fronts is characterized by the presence of a triangle zone formed by the BV Thrust and by a back-thrust spaying-off the MC Thrust.

The recent activity of the MC Thrust is testified by several Middle and Upper Pleistocene warped terraces of the Piave River paleo-course flanking the western termination of the fold (e.g. BENEDETTI et alii, 2000 and references therein), and by the eastward deflection of the Piave River around the growing anticline. Conversely, evidence of recent activity of the BV Thrust are more sparse and not conclusive (GALADINI et alii, 2001). To the east, Late Quaternary activity of the CA Thrust is shown by morphotectonic evidence and folding and faulting of LGM slope deposits (GALADINI et alii, 2005).

Ongoing seismic activity of the ESC results in several destructive M 6+ earthquakes that have been positively associated to individual segments of the external thrust fronts (BASILI et alii, 2008; BURRATO et alii, 2008; DISS WORKING GROUP, 2007; GALADINI et alii, 2005). In addition, shallower M 5–6 events generated by smaller secondary structures pose a not negligible hazard to the region at a more local scale. In this context, the almost continuous seismogenic belt that follows the most external ESC thrust front from the Tagliamento River (epicentral area of the Mmax 6.3 1976 seismic sequence) westward to Bassano (epicentral area of the Mw 6.6 1695 earthquake) is interrupted in the area of the Montello anticline (CPTI WORKING GROUP, 2004), with no earthquakes during the last 700 years (considered the interval of completeness of the catalogue for events of M 6+) (Fig. 2).

**METHOD**

Our analysis focused along the trace of the high precision IGM geodetic levelling line “Mestre-Polpet” that runs in a N-S direction from the plain near Venice to the inner sector of the Venetian Prealps near Belluno, and crosses both the MC-CA and BV thrust fronts (Fig. 1). We analyzed relative elevation changes measured by the geodetic line in a 50 years long time interval, referring them to the first nodal benchmark of the line considered as having a stable elevation (see D’ANASTASIO et alii, 2006 for an explanation of the method).

This study highlighted the occurrence of some segments of the line characterized by positive vertical relative motions with respect to nearby segments. These vertical movements, occurring at the crossing of the ESC thrust fronts, are local signals with a wavelength up to several kilometers superposed to a regional uplift trend. One of these sectors was already studied and linked to the active growing of the Montello anticline (DE MARTINI et alii, 1998). In this study, we adopted a forward modeling procedure for the local geodetic uplift signals to estimate the thrusts recent activity.

To reconstruct the starting geometry of our model we used available seismic exploration data combined with surface geologic and morphotectonic observations, and produced a detailed geologic section along the trace of the leveling line. The geometric parameters of the thrust derived from this section (i.e. strike, angle of dip, depth) were used as input parameters for the modeling of their expected vertical surface dislocation. The results of the modeling were compared to the registration of the geodetic line to derive range values of the principal geometrical fault parameters.

Once the geometry of the active thrust faults were known, we converted the short-term uplift rates obtained from the analysis of the geodetic line into slip and shortening rates. Then, we compared them with: 1- the long- and mid-term geologic/geomorphic slip rates to check for variations through time; and 2- to the GPS shortening budget in order to examine the partitioning of the deformation across the different ESC thrust fronts.

Finally, since the area we studied is a seismic gap (Fig. 2) we also used analytical (e.g. HANKS & KANAMORI, 1979) and empirical (e.g. WELLS & COPPERSMITH, 1994) relationships to figure out maximum magnitude and recurrence interval of possible future earthquakes.

![Fig. 2](http://www.crs.inogs.it/) – Map of historical and instrumental seismicity from the CPTI04 Catalogue (CPTI Working Group, 2004) and the 1977–2001 OGS Annual Bulletin (available from: http://www.crs.inogs.it/). (from BURRATO et alii, 2008, modified).
REFERENCES


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