Marine geophysical methods for archaeological investigation of volcanic and bradyseismic areas

Francesco Giordano
Instituto Universitario Navale, Napoli, Italy

Abstract
The aim of this work was to test several geophysical methods for the identification and study of submerged prehistorical coastlines and archaeological sites. This research program was carried out in collaboration with the «Osservatorio Vesuviano» in the Gulf of Pozzuoli and dealt in particular with the bathymetric strips extending from -5 m to -50 m. Within these strips we identified the coastline dating from the Roman period, former beach boundaries associated with the vertical movements of the earth’s surface caused by seismic-volcanic activity, and the variations in sea level following the climatic changes throughout the last 15000 years. The UNIBOOM system was used for this part of the programme, permitting the identification of several coastlines and submerged beaches lying at different levels. The use of a modern Side Scan Sonar – for the morphological investigation of the sea bed – in a zone which had been the object of numerous archaeological surveys in the past, permitted previously unknown structures near the Lacus Baianus to be identified. Other features worth pointing out include the operating speed of the system (15000 m²/first minute approx) and its observation capacity in cloudy waters compared to visible radiation, as well as its ability to penetrate thin layers of mud which generally impede direct underwater observation.

Key words  Gulf of Pozzuoli (Naples) – marine archaeology – marine prospections – geophysical methods – marine archaeological survey

1. The results obtained with the UNIBOOM system

The UNIBOOM system is a marine acoustical system characterized by a high vertical resolution power supply and high working frequency 400-14000 Hz (1 Hz = 1 cycle/s). Two layers lying .20 m apart can be resolved provided that the surface which separates the two layers of different media is visible, i.e. if their acoustical impedences are sufficiently different (Stefanon, 1985a).

The UNIBOOM survey led to the discovery of former submerged beaches and covered coast lines. The area investigated lies to the north-east of the city of Naples and is marked on the map as the Gulf of Pozzuoli (fig. 1). The phenomena we have tried to discover occurred over the past 10000 years, during which time there has been a variation in the sea level of approximately +100 m (Pescatore et al., 1984, Luongo et al., 1991). There are several geological-structural maps of the Gulf of Pozzuoli in print, but the most complete one «the morphostructure of the Gulf of Pozzuoli» (1-20000) was drawn up by Segre (Segre, 1972).

In the last 25 years the area has been the subject of a variety of important geophysical surveys (reflection seismics), but none of these were carried out along the coastline and none had the same aims as our research. During the months of February and March 1993 we investigated strips of the coastline lying between the
bathymetrics of 5-50 m. Figure 1 shows the navigation route taken. The readings were preceded by numerous trial runs in order to obtain the best results from the measuring equipment, and in particular:

- the signal filters;
- the regulating of the geometry of the transducer-receiver couple of the UNIBOOM system.

When interpreting the data we initially used Segre’s map (mentioned above) as a reference for standardization purposes, but later we proceeded independently. The peculiarity of our measurements was due to the following factors:

- the precise positioning of the boat was facilitated by the GPS system;
- a small, manageable fishing boat was used;
- good signal penetration with little reverberation was achieved even in deep water.

Thus we were able to detect:

- areas of marine erosion, which were uncovered or covered by more recent sediments;

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**Fig. 2.** Stratigraphy of the sea bed in front of the town of Bacoli along the route M27 to M31. An accumulation of detritus 10 m thick is evident between bathymetrics 40 and 50. Traces of beach and erosions are also present.
these are characterized on the graph recording by the sudden interruption of the sedimentary lines;
  – former beaches which were superficial or covered, characterized by the presence of pebbles and chaotic material; these were easily recognizable in the recordings.

Further on we will analyse examples of the most important recordings.

The sea bed lying in front of the town of Bacoli (fig. 2) is characterized by the presence of a highly accentuated underwater slope. Along the route from M27 to M31 there is an obvious accumulation of material which is more than 10 m thick between bathymetries 40 and 50 in a parallel direction to the map. This recording picked up traces of beaches and slightly accentuated erosions.

Figure 3 shows the route from L35 to L39, including a 180° change of course at L37. This recording points out a strip of chaotic deposits which is more or less accentuated according to the morphological slant of the main basic horizontal. The maximum thickness of approximately 5-6 m is reached according to the –60 m bathymetric.

Another interesting route lay in front of Arco Felice, running almost parallel to the coastline, from P64 to P69 in fig. 4. A stratification which could be correlated to the presence of compact materials of a regular struc-

![Fig. 3. Route from L35 to L38 with a 180° change of course in L37; a stratification of chaotic material 5-6 m thick is present.](image-url)
Fig. 4. Route very near the coastline from P64 to P69 in front of the town of Arco Felice; the first horizon is made up of a highly regular stratification, thus suggesting the presence of artificial paving on top of uneven material.

The route opposite Arco Felice, shown in fig. 5, running south-northwards along L26 to L31 from bathymetric 50 m to 5 m, contained two examples of terracing (one from −50 m to −40 m, and the other from −25 m to −12 m), and traces of covering by material which from the recording appeared pebbly and not very even. This could therefore be correlated to an extended example of underwater beach. The bed-rock (substratum) is present underneath, and has areas where it is missing due to strong absorption by the layers on top, but nonetheless it is easy to follow. Considering that there is a 23 times vertical deformation the slopes, it can be deduced that the slope of each layer is in the region of 1.5-2 degrees, according to bibliographical data (Pescatore et al., 1984). The two detritic masses, one lying at approximately −20 m and the other at −40 m can respectively be correlated to units C and D of Segre’s Map, even though their extensions and positioning do not agree exactly with each other. We will mention this again later.

As can be seen from the navigation map (fig. 1), readings were also taken in the area around the island of Nisida and beyond, but their quality was not sufficient to obtain proper coverage of the area.
2. Marine archaeological surveys using the Side Scan Sonar

The Side Scan Sonar (SSS) is usually employed using a moving ship which tows the transducer ("fish", i.e. a transmitter and receiver) at a working speed of 2 to 6 knots. It is therefore possible to investigate large areas in a relatively short time. The instrumentation is fully portable and can be used in small boats. The Side Scan Sonar acoustically records in real time an image of the sea bed which is comparable to aerial photographs or scanner images of land. The range of lateral coverage is inversely proportional to the outgoing frequency: a working frequency of 100 kHz gives a maximum range of around 600 m (300 m on both sides of the track) and lower frequencies (50 kHz) easily reach the km (Stefanon, 1985b).

We also used the SSS to investigate the presence of archaeological structures in shallow water (approximately 10 m). It is, in actual fact, possible for divers to take direct readings using a camera or a video camera, the only limitation being that a wide-angle panorama of the situation cannot be obtained. However, difficulties can also arise in shallow water due to
Fig. 6. Side Scan Sonar investigated area: Gulf of Pozzuoli after Camodeca, modified with structures 2-3-6.
Fig. 7. Side Scan Sonar image of a strip of sea in front of the trading port of Baia running north-eastwards.  
1) An outline can be seen in the depression suggesting a ship wreck just emerging from the sea bed; 4) and 5) are images of the two impressive walls of the access canal to the Lacus Baianus, (Camodeca, 1987); 2) and 3) are wall structures unknown to bibliographies and their position would seem to indicate that they were used at a different period of time to (4) and (5) as they correspond probably to a different sea level.
Fig. 8. Side Scan Sonar image of a part of Calpurnio Pisone's villa in the area of water opposite Epitachio Head.
murminkess and to the presence of lightweight mud on the sea bed. This does not occur with SSS.

SSS allows the sea bed to be explored in 100, 150 and 200 m wide strips of unlimited length and with a covering speed of over 10000 m² per first minute. The recordings can be read simply and almost immediately as will be demonstrated in the examples below.

We carried out a brief trial run of the SSS KLEIN (double frequency, 100-500 kHz) in the port area of Baia and Epitaffio Head. Exploring with a double frequency has the added advantage of being able to penetrate slightly beneath the underwater vegetation and the muddy blanket on the sea bed using the lowest frequency, but the highest frequency of 500 kHz produces more accurate readings. The recording strips shown in figs. 7 and 8 represent the acoustic images which correspond to a working frequency 500 kHz; the theoretical resolution power is:

\[ T. \text{ RES. POW.} \approx \text{WAVELENGTH} \approx \frac{\text{Velocity of Sound/Frequency}}{(500 \text{ kHz})} = 3 \cdot 10^{-3} \text{ m.} \]

The recordings in figs. 7 and 8 were obtained along a route starting from the area facing the Baia Castle and ending around Epitaffio Head, also fig. 6. The sea beds explored lie at a depth of approximately −10 m.

The recording in fig. 7 shows part of the route running parallel to the coast in front of Baia Port (proceeding north eastwards). The first part of the recording gives evidence of a depression – the clear zone in the recording – and in the middle there is a trace (mark 1)

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Fig. 9. A probable evolution of the Phlegraean coastline in the last 10000 years (after Albore Livadie, 1986, modified).
which looks like a sunken ship, approximately 20 m long, half covered with muddy sand. At the end of this depression, corresponding to mark 2, there is a sort of wall which lies semi-transversally compared to the route. Running perpendicularly to this wall lie tracks almost certainly caused by the dragging of anchors along the sea bed, since merchant ships dock near this area. Moving along, several impressive wall structures corresponding to 3, 4, 5, 6 are noted, and in particular numbers 4 and 5. They correspond to the walled sides of the access canal to the Lacus Baianus, dating from Roman times (Camodeca, 1987). The traces in 2, 3 and 6 do not appear to have been noticed by previous authors (see also fig. 6). In the route shown in fig. 8, which was approximately 300 m long and took 3 min to cover, a portion of the ruins of Calpurnio Pisone’s villa, dating from the Roman Empire, was clearly visible. Close inspection showed the walls which were still standing were less than one metre tall. The resolution power of the instrument was even able to pick up the bricks which had fallen away and were lying on the sea bed.

3. Conclusions

The results obtained in this brief geophysical survey demonstrate the immense value of this type of sedimentological and evolutionary research in volcanic areas with the following phenomena: the deposits of eruptive products above and below water; upward and downward bradyseismic phenomena; massive erosions and collapses of various materials due to wave action. Geophysical methodologies are also able to discover anthropical structures and pick up detailed readings from them in a quick, direct and total manner, and are therefore useful for underwater archaeology. Marine geophysical surveys make continual spatial distribution data available, which are not easily obtainable on land. Defining the physical evolution of marine-continental-marine environments makes it possible to define the relationship between them and the presence of anthropical settlements.

A probable evolution of Phlegraean Field coastline in the last 10000 years is showed in fig. 9, merged with archaeological lines of the C. Albore Livadie map (Albore Livadie, 1986).

REFERENCES