

# Soil Dynamics and Earthquake Engineering

## A GIS procedure for the topographic classification of Italy, according to the seismic code provisions --Manuscript Draft--

<b>Manuscript Number:</b>	SOILDYN-D-21-00049R2
<b>Article Type:</b>	Research Paper
<b>Keywords:</b>	DEM; Topographic classification; EC8; NTC18; GIS; Morphometric analyses
<b>Corresponding Author:</b>	Claudia Mascandola INGV: Istituto Nazionale di Geofisica e Vulcanologia ITALY
<b>First Author:</b>	Claudia Mascandola
<b>Order of Authors:</b>	Claudia Mascandola Lucia Luzi Chiara Felicetta Francesca Pacor
<b>Abstract:</b>	<p>In this study we present a procedure for the topographic classification of Italy, taking advantage of existing high-resolution digital elevation models (DEM), with the support of routines embedded in Geographic Information Systems. The proposed method is based on morphometric analyses of a DEM, allowing for hilltop ridge detection, slope computation, identification of reliefs with potential topographic site effects, and topographic classification according to the indications of the current European and Italian seismic codes. The developed procedure can be applied worldwide and has a potential engineering interest for a fast and accurate topographic classification of a site, for both scientific and application purposes. The topographic classification of Italy is tested on several well-known cases of topographic amplification and it is adopted for the topographic classification of Italian accelerometric stations.</p>
<b>Suggested Reviewers:</b>	<p>Emilia Fiorini Eucentre emilia.fiorini@eucentre.it Expert in GIS and site effects</p> <p>Iunio Iervolino Universita degli Studi di Napoli Federico II Facolta di Ingegneria: Universita degli Studi di Napoli Federico II iunio.iervolino@unina.it Expert in Earthquake Engineering</p> <p>David Wald USGS: US Geological Survey wald@usgs.gov Expert in site effects and topographic data</p> <p>Nikolaos Theodoulidis ITSAK ntheo@itsak.gr Expert in site effects</p> <p>Maria Rosa Gallipoli IMAA-CNR: Istituto di Metodologie per l'Analisi Ambientale Consiglio Nazionale delle Ricerche mariarosaria.gallipoli@imaa.cnr.it Expert in site effects</p>
<b>Response to Reviewers:</b>	

## Response to Reviewer #2:

1. The answer to my first two concerns is not fully satisfactory since the main point is missed. I try here a new formulation of this point. The amplification coefficients reported in the seismic code is the result of numerical simulations relative to a limited set of morphological profiles. When one considers other configurations (e.g., the 'unclassified' ones in Pessina and Fiorini, 2014) applicability of the same coefficients is not warranted. Thus, the reduction of 'unclassified' sites is not necessarily a positive evidence encouraging the application of the new approach: actually, it may represent an improper application of the procedure outside its actual domain. Other terms, the question is: at what extent the amplification coefficients can be actually applied to the configurations you have identified and were missed by Pessina and Fiorini? This is the point, which, in my view, should be addressed. Of course, I not asking for a revision of the procedure the Authors propose, but only a clear statement (to be reported in the 'Conclusion' section also), about this key aspect.

*Pessina and Fiorini missed some configurations, since they did not use at the same time the concept of slope and ridge. In their tables 3 and 4, they make use of qualitative terms, like “close” and the classification of the site is driven by a rather subjective approach. This is why some of the configurations are missed. In our case, we make use of a quantitative approach, so that we can specify, for each pixel, the difference in height, the slope, the presence of a ridge, through the use of filter operators with specific dimensions that allow us to evaluate both ridge and slope, contemporarily. An example for the limit of the Pessina and Fiorini procedure is visible in their table 3, in case of the combination of R0-S3 codes (where sites are Not directly classified). The authors state that “stations are generally classified as T2, it is impossible to detect if they are at the base of slope, in the middle of slope, near terraces or in narrow valley”.*

2. I am not sure that when hazard (S1, S2, etc.) is evaluated at the same sites where maximum intensity were assessed, no "significant differences" are observed. A comparison between Figs. 2a and 2b in the answer puts in evidence important differences in the distribution of sites in the morphological classes T4 (and T2) relative to S3 and S4 hazard classes. When Figure 2b is compared with figure 7c, the respective patterns appear quite similar. Thus, the statistical significance of eventual differences (possibly induced by topographical amplification effects) should be assessed in some quantitative way. This is an important aspect to convince skeptical colleagues (including myself) that topographical features (at least the ones identified by the Authors) generally play a significant role in seismic hazard.

*To address this point, we have reformulated the paragraph ‘Statistics on the topographic classification map’. In particular, we have added a new figure (current Figure 7) with statistics on the 15332 Italian localities in DBMI15, grouped on the base of maximum observed macroseismic intensity at a site ( $I_{max}$ ).*

*As described in the manuscript, in our opinion Figure 7 shows two aspects.*

### *1. A relation between $I_{max}$ and the seismic zones (S1-S4).*

*“The localities with the lowest  $I_{max}$  (Figure 7a) are, of course, occurring in correspondence of the lowest seismic hazard zones (S3 and S4,  $a_g \leq 0.15$ ) in Northern Italy, whereas the highest intensities are observed in correspondence of seismic hazard zones S1 and S2 ( $a_g > 0.15$ ) in the Central-Southern Apennines (Figure 7d). “*

2. A relation between  $I_{max}$  and the topographic classes (T1-T4).

*“Based on the classification proposed in this study, about 40% of the localities with the  $I_{max} > 6$  could be affected by topographic amplification (T2-T4 classes, Figure 7c-d), in contrast to about 20% for  $I_{max} < 6$  (Figure 7a-b). Moreover, we observe a large occurrence of the sites with potential topographic amplification (T2-T4 classes) in correspondence of localities characterized by medium to high  $I_{max}$  (range 6 - 8), in seismic zone 2 (S2:  $0.15 < a_g \leq 0.25$ ), as shown in Figure 7c”.*

***In conclusion:*** based on our classification, the sites with potential topographic effects are clustered in zones with high seismic hazard (S1-S2 zones), where the higher macroseismic intensities are observed ( $I_{max} > 6$ ).

*These results do not prove the existence of topographic amplification at those sites classified in T2-T4, but the coherency between the topographic classification (T1-T4), the maximum observed microseismic intensity ( $I_{max}$ ) and the seismic hazard map, could be a hint of the occurrence of topographic effects. The concern of the reviewer about the significant role of topographic effects in seismic hazard will be further investigated in a future study that takes into account the residual distributions computed from Intensity Prediction Equations (IPE).*

- Topographic classification of Italy according to the Italian seismic code
- Novel GIS procedure
- Around one third of the Italian territory has sites with possible topographic amplification



29 mountain [1]; the focusing of seismic waves near the crest because of the reflection on a free surface and/or  
30 the interaction between incident and diffracted waves [2]. Moreover, topographic seismic amplification is  
31 considered as one of the possible causes of triggers of earthquake induced phenomena, like landslides, which  
32 represent one of the major causes of earthquake-related devastation (e.g. Refs. [3,4]). Several studies over the  
33 last four decades have discussed the effects of surface topography on seismic ground motions from  
34 instrumental investigations [5-12] and numerical models [13, 14, 4, 15, 16]. Relevant macroseismic case  
35 histories are reported by several authors [17, 18, 4, 19, 20, 21]. The proper assessment of topographic  
36 amplification is of key importance in the European context, where a broad range of exposure levels to  
37 earthquake hazard is encompassed [18] and many small towns are located at the top of isolated steep cliffs.  
38 Such unfavourable conditions often come together with high vulnerability of the buildings and the  
39 susceptibility to slope sliding [4]. Therefore, the topographic amplification, even if generally do not exceed  
40 the stratigraphic one (e.g. Refs. [4, 22, 23, 24]), plays a significant role in seismic design, microzonation studies  
41 and hazard analyses.

42 The current Italian (NTC, 2018; hereinafter NTC18 [25]) and European (CEN, 2004; hereinafter EC8 [26])  
43 seismic codes prescribe frequency-independent amplification factors not larger than 40%, based on the height  
44 and the slope of simple 2D relief configurations. The approach, introduced in the EC8, was further adopted by  
45 NTC (2008) [27] and the current NTC18, where four classes of topographic amplification are explicitly  
46 defined, as listed in Table 1. These categories refer to bidimensional configurations, such as elongated crests  
47 and ridges, and must be considered only when their height is greater than 30m.

48

49

50

51

52

53

54 **Table 1:** Topographic amplification factors prescribed by Italian seismic code (NTC18).

Topographic Class	Description	Amplification factor
T1	Flat surface, isolated slopes and cliffs with average slope angle $i \leq 15^\circ$	1
T2	Slopes with average slope angle $i > 15^\circ$	1.2
T3	Relief with ridge width much smaller than the base and average slope angle $15^\circ < i \leq 30^\circ$	1.2
T4	Relief with ridge width much smaller than the base and average slope angle $i > 30^\circ$	1.4

55

56 Following the classification of the Italian seismic code, Ref. [24] first implemented a procedure for the  
 57 topographic characterization of seismic recording stations archived in ITACA 2.0 database (Italian  
 58 Accelerometric Archive, [http://itaca20.mi.ingv.it/ItacaNet\\_20/](http://itaca20.mi.ingv.it/ItacaNet_20/); [28, 29]). The procedure is based on  
 59 morphometric analyses of high-resolution digital elevation models (DEMs), with the support of Geographic  
 60 Information Systems (GIS). In the specific, the procedure computes the slope and ridge layers from DEM,  
 61 besides different statistical indicators (i.e., maximum, minimum, sum, and mean) within a circular buffer area  
 62 around each recording station. These statistical indicators are combined to provide both a slope and a ridge  
 63 index that are further matched to define the topographic classification of the seismic station. However, some  
 64 combinations of the indices do not permit a univocal topographic classification because their arrangement may  
 65 lead to different topographic classes (in that case the site is not classified). Moreover, the procedure  
 66 implemented for the ridge detection is not straightforward and implies several noise removal steps to enhance  
 67 significant features and remove small topographic irregularities. Since the procedure is applied to the  
 68 surroundings of each seismic station, it needs to be run each time a new seismic station is added to the database.  
 69 These downsides led to the development of a new GIS-based procedure for topographic classification of the  
 70 whole Italian territory. However, the NTC18 prescribes the applicability of the topographic coefficients in case  
 71 of “simple” configurations, but the criteria to discriminate between “simple” and “complex” are not specified  
 72 and, for this reason, the applicability of the coefficients could be debated.

73 In this study, we propose a step-like procedure and provide some statistics of the resulting map, correlating it  
74 to the probabilistic seismic hazard [30, 31], the macroseismic intensity data in DBMI15 [32], and the seismic  
75 stations in the Italian Accelerometric Archive 3.1 (ITACA 3.1; [http://itaca.mi.ingv.it/ItacaNet\\_31/#/home](http://itaca.mi.ingv.it/ItacaNet_31/#/home);  
76 [33]). The topographic classification is tested at several sites, where topographic amplification has been  
77 observed. Finally, based on the proposed classification, we evaluate the impact of the topographic  
78 amplification factors (Table 1), on the probabilistic seismic hazard map calculated for Peak Ground  
79 Acceleration (PGA), a return period of 475 years, and stiff soil i.e. soil class EC8-A [30].

80

## 81 **2. Method**

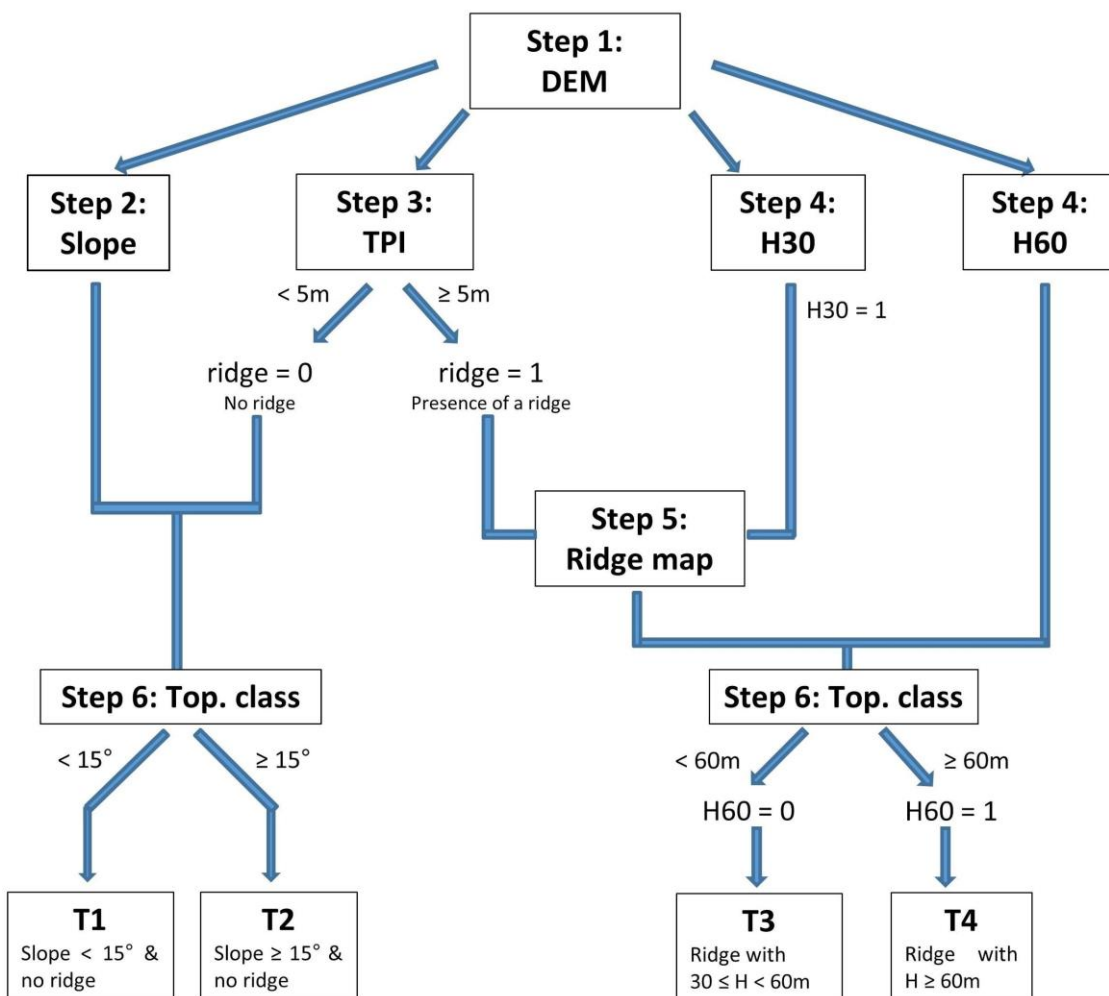
82

83 Digital elevation models (DEMs) are fundamental in environmental and morphological analyses. In this study,  
84 the topographic classification is applied to the so called *TinItaly DEM* [34], created for the whole Italian  
85 territory as a triangular irregular network (TIN), and converted into grid format of 10m cell size, according to  
86 a tiled structure composed of 193, 50-km side square elements. Ref. [34] carried out a comprehensive  
87 assessment of the accuracy of the TinItaly DEM, finding a root mean square error in elevation (RMSEz)  
88 between 0.8 and 6.0 m. For our analysis, we select the TinItaly DEM since, to date, it is the model with the  
89 highest resolution, freely available for scientific purposes for the whole Italian territory.

90 The morphological analyses applied on the DEM are summarized in Figure 1 and described in detail in the  
91 following subparagraphs. Starting from DEM (Figure 1, step 1), the slope angle is computed (step 2) along  
92 with the Topographic Position Index (TPI; [35], step 3) and the elevation range (maximum minus minimum  
93 elevation value in a moving window), applying a threshold height of 30 m (H30) and 60 m (H60), respectively  
94 (step 4). Since the positive TPI values allow detecting the ridge zones, an empirical threshold of 5 m is adopted  
95 to discriminate between ridges (ridge is assigned 1) and non-ridges (ridge is assigned 0). Considering the slope  
96 and TPI values, a first discrimination of topographic classes can be performed (Figure 1). Whereas T1 and T2  
97 sites are on slopes (ridge = 0), T3 and T4 sites should be on ridges (ridge = 1). Therefore, T1 and T2 categories  
98 can be identified by the combination of ridge = 0, and slope angle  $< 15^\circ$  or  $\geq 15^\circ$ , respectively (step 6); T3  
99 and T4 classes can be identified by ridge = 1 (Figure 1, step 5), on reliefs with difference in height  $> 30$  m. To  
100 further discriminate between T3 and T4 classes, the H60 map is considered. When H60 = 1 (elevation range



101  $H$  is between 30 and 60 m) the site is in class T3 (step 6). In the following, the step-like procedure applied to  
 102 the TinItaly DEM is described in detail, indicating the ESRI ArcGis ([www.esri.com](http://www.esri.com)) tools adopted for the  
 103 morphological analyses.  
 104



105  
 106 **Figure 1:** Workflow for topographic classification.

107  
 108 **Step 1: preliminary DEM processing**

109 To improve the computation time and data handling, the 193 TinItaly DEM tiles are resampled with a 40 m  
 110 resolution. Afterwards, the DEM tiles are merged to obtain a unique map for the Italian territory. Finally, to  
 111 remove minor irregularities, a smoothing is performed by applying a mean algorithm with a 3x3 moving  
 112 window (e.g. *Focal Statistics* tool in ESRI ArcGis), as shown in Figure 2a for a sample area in central  
 113 Apennines.

114

### 115 **Step 2: computation of topographic slope**

116 To compute the slope angle, we use the *Slope* function implemented on ESRI ArcGis (Figure 2b). The slope  
117 is computed in degree units in a 3 by 3 cell neighborhood (i.e., moving window) on the smoothed DEM,  
118 considering the planar method with the average maximum technique [36]. The slope is measured as the  
119 maximum rate of change in value from a cell to its immediate neighbors. The calculation is performed on a  
120 projected flat plane using a 2D Cartesian coordinate system. The processed cell is judged reliable if the  
121 computation is performed on at least seven cells in the 3 by 3 neighborhood.

122

### 123 **Step 3: ridge extraction**

124 For the ridge detection, we adopt the Topographic Position Index (TPI) algorithm proposed by Ref. [35]. This  
125 algorithm is implemented in the *Land Facet Analysis* toolbox [37] in ESRI ArcGis. The TPI is defined as a  
126 difference between the actual value of the elevation  $H$  at given location  $[x, y]$  and mean elevation of the  
127 neighboring area  $A$ :

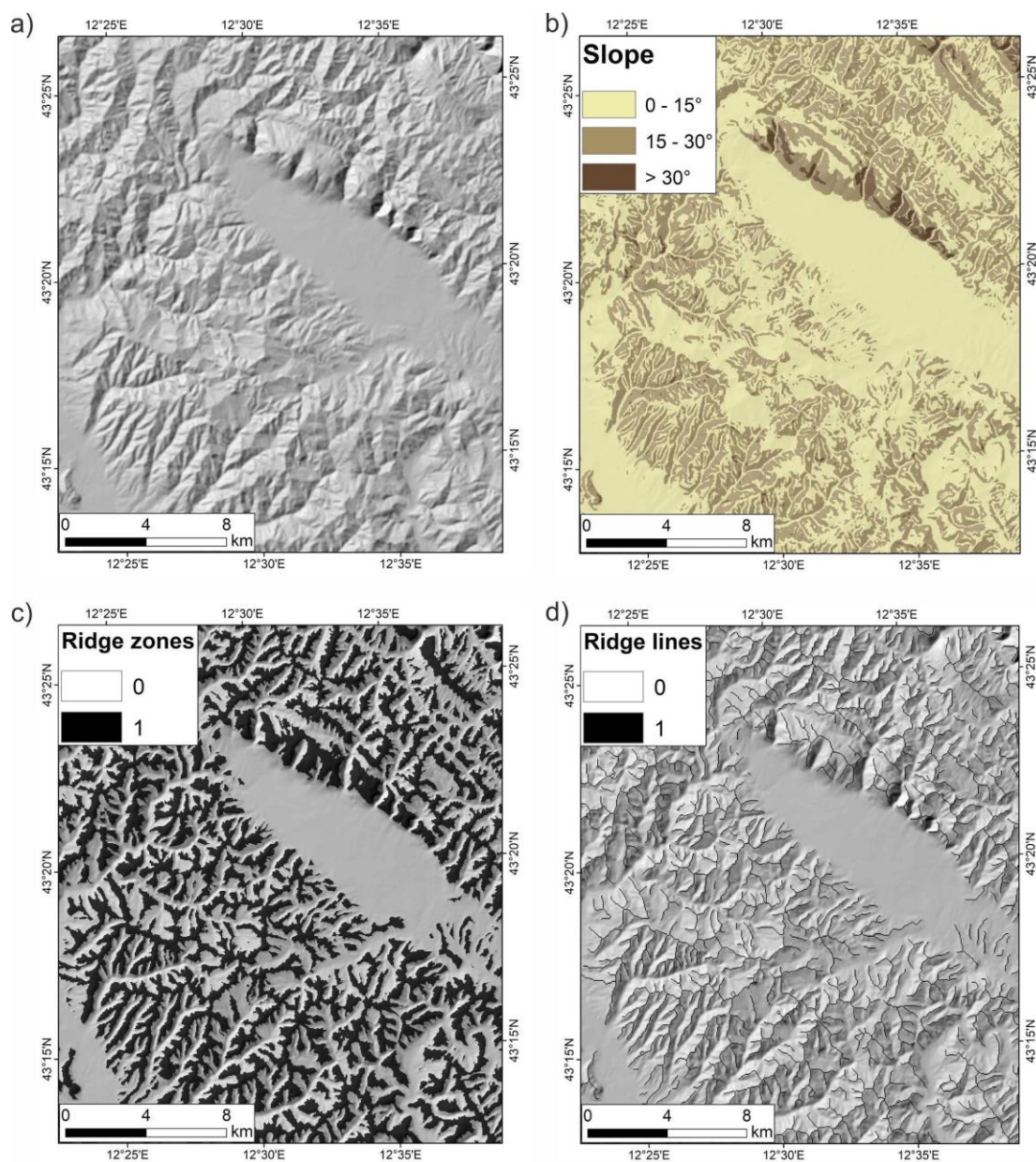
128

$$129 \quad TPI(x, y) = H(x, y) - \int_A H(x, y) \partial S / \int_A \partial S \quad \text{Eq. (1)}$$

130

131 The area  $A$  should be centered at the point  $[x, y]$ . Both the size and the shape of the neighboring area have to  
132 be defined in advance, depending on the specific application. In general, there is no constraint on both values.  
133 The TPIs reflect the differences between the elevation in a particular cell and the mean elevation of the  
134 surrounding cells: positive values mean the cell is higher than its surroundings, while negative values mean it  
135 is lower. Therefore, high positive values of TPI characterize ridgetops and hilltops, while negative values  
136 define basins or valley bottoms. TPI values close to zero indicate flat or mid-slope areas. This kind of analysis  
137 strictly depends on the scale used to count the surrounding cells. To set the interval for the TPI analysis, we  
138 carry out some tests, considering different radii (i.e. 300 m, 500 m and 800 m). For the Italian case, a  
139 neighborhood circle with a radius of 500 m is adopted, since it allows better discriminating the ridges out of  
140 the main reliefs, while ignoring the minor deflections [38]. The ridge zones are obtained from the TPI map  
141 (*Raster Calculator* tool) considering a threshold value equal to 5. The output is a binary raster that takes the  
142 value 1 when  $TPI > 5$  and 0 otherwise (Figure 2c). Since TPI units are in meters, a TPI value of 5 means that

143 this particular cell is 5 m higher than the average elevation of its neighborhood (i.e., a radius of 500 m). The  
 144 TPI threshold is selected empirically. A lower threshold (down to 1) does not significantly affect the final  
 145 results. The ridgelines are retrieved by applying a thinning function (*Thin* tool) that thins rasterized linear  
 146 features by reducing the number of cells representing the width of the features [39], as shown in Figure 2d.

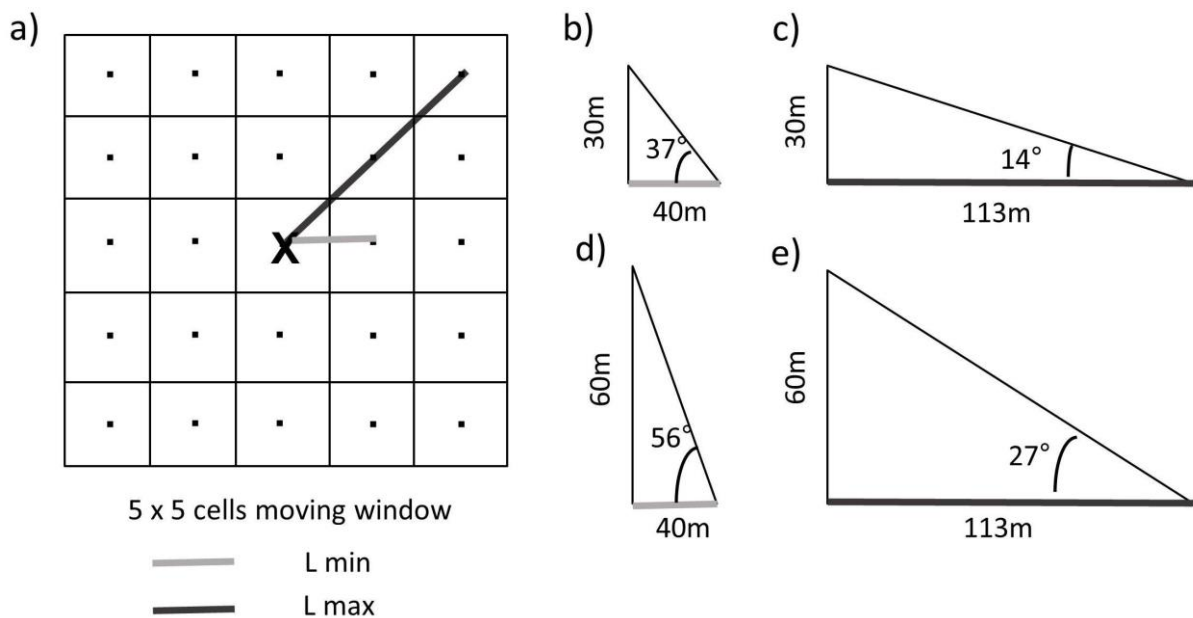


147  
 148 **Figure 2:** Extraction for an area of the Central Apennines of a) DEM, represented as hillshade; b) topographic  
 149 slope; c) ridge zones setting TPI > 5; and d) ridge lines after thinning.

150  
 151 **Step 4: detection of the difference in height to extract topographic categories T3 and T4**

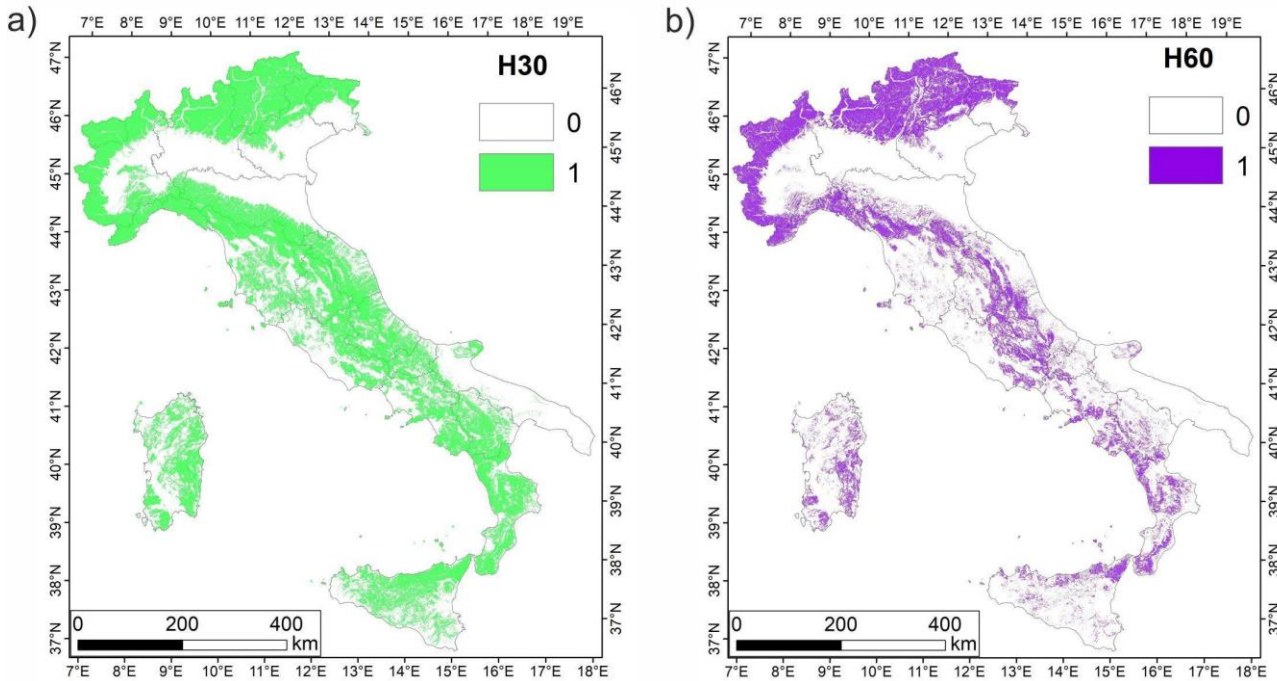
152 In the EC8 and NTC18 seismic codes, reliefs higher than 30 m are considered to have potential topographic  
 153 site effects (i.e., T3 or T4 topographic classes in NTC18). To individuate these reliefs, we compute the

154 elevation range, i.e., maximum minus minimum elevation value, in 5x5 cells moving windows (Figure 3a;  
 155 *Focal Statistics* tool). Considering that the DEM cell size is 40 m, the maximum and minimum semi-transverse  
 156 axis of a relief spanned by a 5x5 moving window is 113 m (e.g.  $80 \cdot \sqrt{2}$ m), as exemplified by the dark grey  
 157 line in Figure 3, and 40 m (light grey line in Figure 3), respectively. Considering a threshold difference in  
 158 height of 30 m (condition for topographic amplification), we obtain the slope angle interval, which ranges  
 159 between  $14^\circ$  and  $37^\circ$  (Figure 3b,c), thus including both T3 and T4 classes. To discriminate class T4, we  
 160 consider the map representing the difference in height of 60 m that individuates the slope angles in the interval  
 161  $27^\circ$ - $56^\circ$  (Figure 3d,e), including T4 sites (i.e, ridges with height greater than 30 m and average slope angle  $i >$   
 162  $30^\circ$ ). In the following sections, the elevation range computed with a threshold height of 30 m is indicated as  
 163 H30; whereas the elevation range computed with a threshold height of 60 m is indicated as H60 (Figure 4).



164  
 165  
 166 **Figure 3:** Trigonometry applied to H30 and H60 maps. a) 5x5 cells moving window with the smaller (light  
 167 grey line) and larger (dark grey line) semi-transverse axis spanned in the moving window. b)-c) maximum and  
 168 minimum slope angles, considering a difference in height of 30 m, d)-e) maximum and minimum slope angles  
 169 considering a difference in height of 60 m.

170



171

172

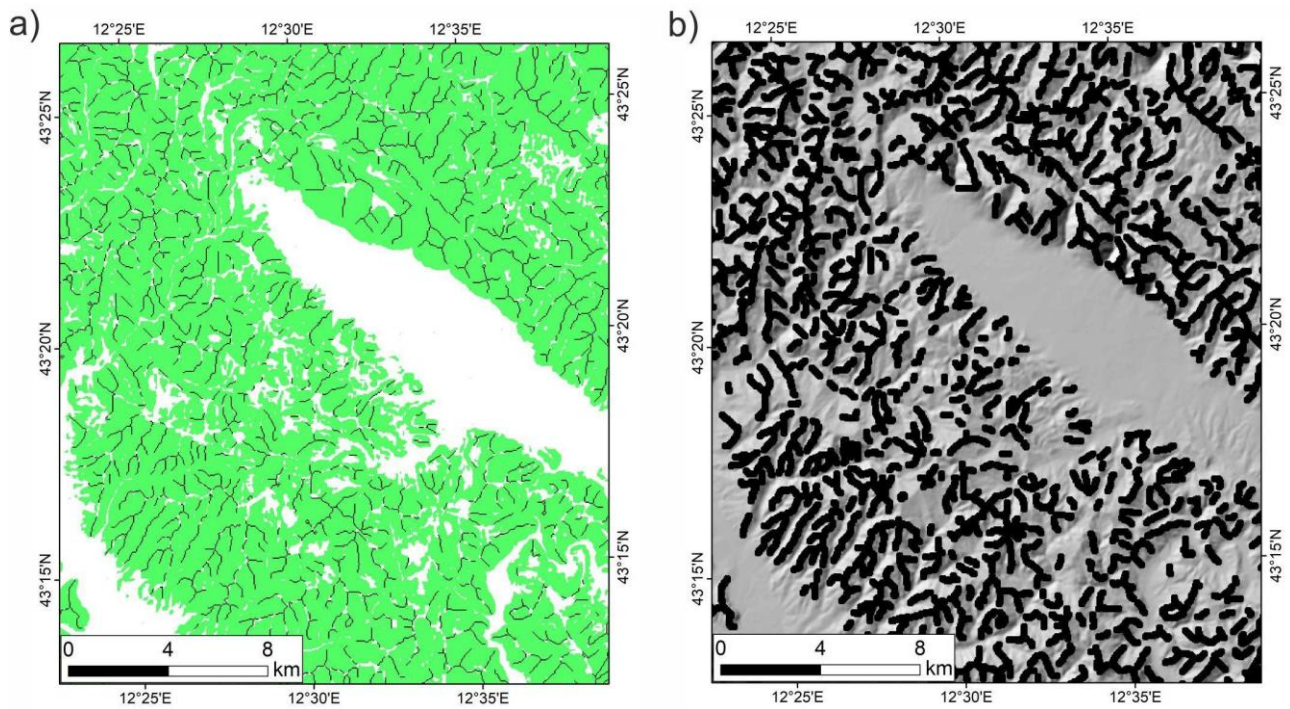
173 **Figure 4:** Map of a) difference in height of 30 m (H30) and b) 60 m (H60) for the Italian territory.

174

175 **Step 5: identification of T3 and T4 topographic categories**

176 After the identification of the difference in height to extract T3 and T4 topographic categories, the ridge lines  
 177 of the corresponding reliefs can be obtained by overlapping the ridge line map (evaluated in step 3) and the  
 178 H30 map (evaluated in step 4) (*Raster calculator* tool, imposing  $H30 = 1$  & ridge lines = 1), as shown in Figure  
 179 5a. To identify the sparse groups of points (< 5 cells), the *Region Group* tool is combined with the *Set Null*  
 180 operator. The *Region Group* counts the number of grouped cells within an 8-cells neighborhood (excluding 0  
 181 values), and the *Set Null* sets to null the cell values for counts less than 5. The sparse groups of points are  
 182 finally removed with the *Nibble* operator that allows filtering of the ridge line raster. Finally, the extracted  
 183 ridges are dilated (3 cells per side), using the *ArcScan* extension (Figure 5b). Being the raster cell size 40 m,  
 184 the ridge lines are dilated by 120 m per side to approximate the semi-transverse axis of a relief with a minimum  
 185 30 m of difference in height and a consequent minimum slope angle of 15°, which are the thresholds for  
 186 topographic amplification according to NTC18 and EC8 seismic codes (see Figure 3c as an example). In this  
 187 sense, the 120 m per side is conservative because it may include the whole relief down to the base.





188

189 **Figure 5:** a) intersection between ridge line map (black lines) and H30 map; b) ridge map on DEM represented  
 190 as hillshade.

191 **Step 6: topographic classification**

192 Once the identification of the morphological parameters is performed, the final topographic classification is  
 193 obtained by combining the following maps:

- 194 1. slope angles (*slope* map);  
 195 2. ridges of reliefs with potential topographic amplification (*ridge* map);  
 196 3. reliefs with average slope angle  $i > 30^\circ$  (*H60* map);

197 as reported in Table 2 (*Raster calculator* tool). The final topographic classification is reported in Figure 6 and  
 198 provided in the electronic supplement.

199

200

201

202

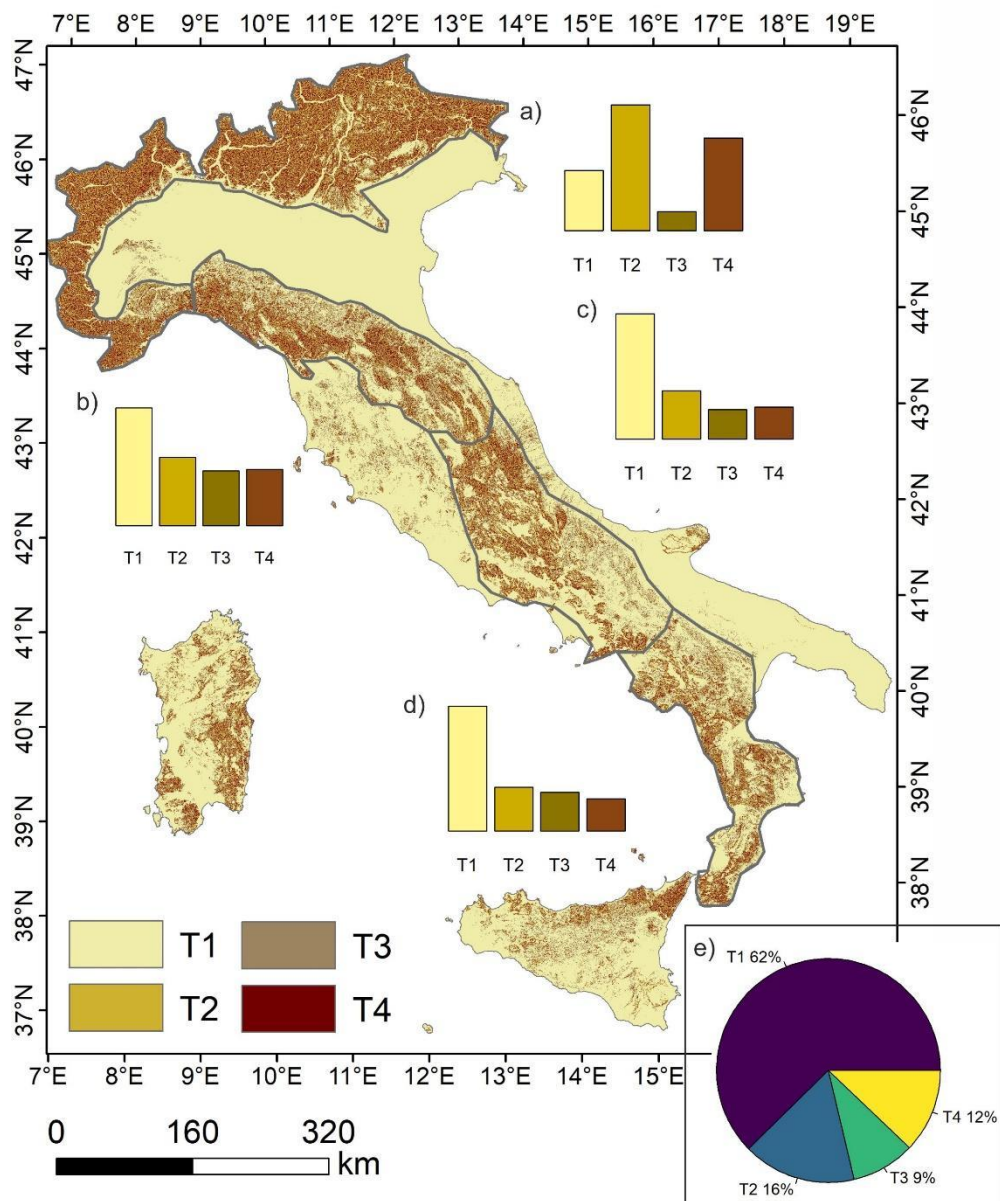
203

204 **Table 2:** Topographic classification (NTC18), based on the proposed GIS-procedure.

205

Condition	Class	Description
ridge = 0 & slope < 15°	T1	slope with average slope angle $i < 15^\circ$
ridge = 0 & slope $\geq 15^\circ$	T2	slope with average slope angle $i \geq 15^\circ$
ridge = 1 & H60 = 0	T3	ridge with difference in height in the range 30-60 m (i.e. base angle $\leq 30^\circ$ )
ridge = 1 & H60 = 1	T4	ridge with difference in height > 60 m (i.e. base angle > 30°)

206



207

208

209 **Figure 6:** Topographic classification of Italy according to the current Italian seismic code (NTC18), obtained  
210 following the proposed procedure. The histograms of the topographic classes are reported for the Alps (a) and  
211 the Apennine (b, c, d) mountain chains, in the Northern, Central, and Southern sectors respectively (grey  
212 outline). e) Pie chart of topographic classes (T1-T4) for the whole Italian territory. The statistics are computed  
213 on a sample of 10,000 random points distributed homogeneously throughout the Italian territory. The count is  
214 normalized to the number of points in each sector.

215

216

### 217 3. Statistics on the topographic classification map

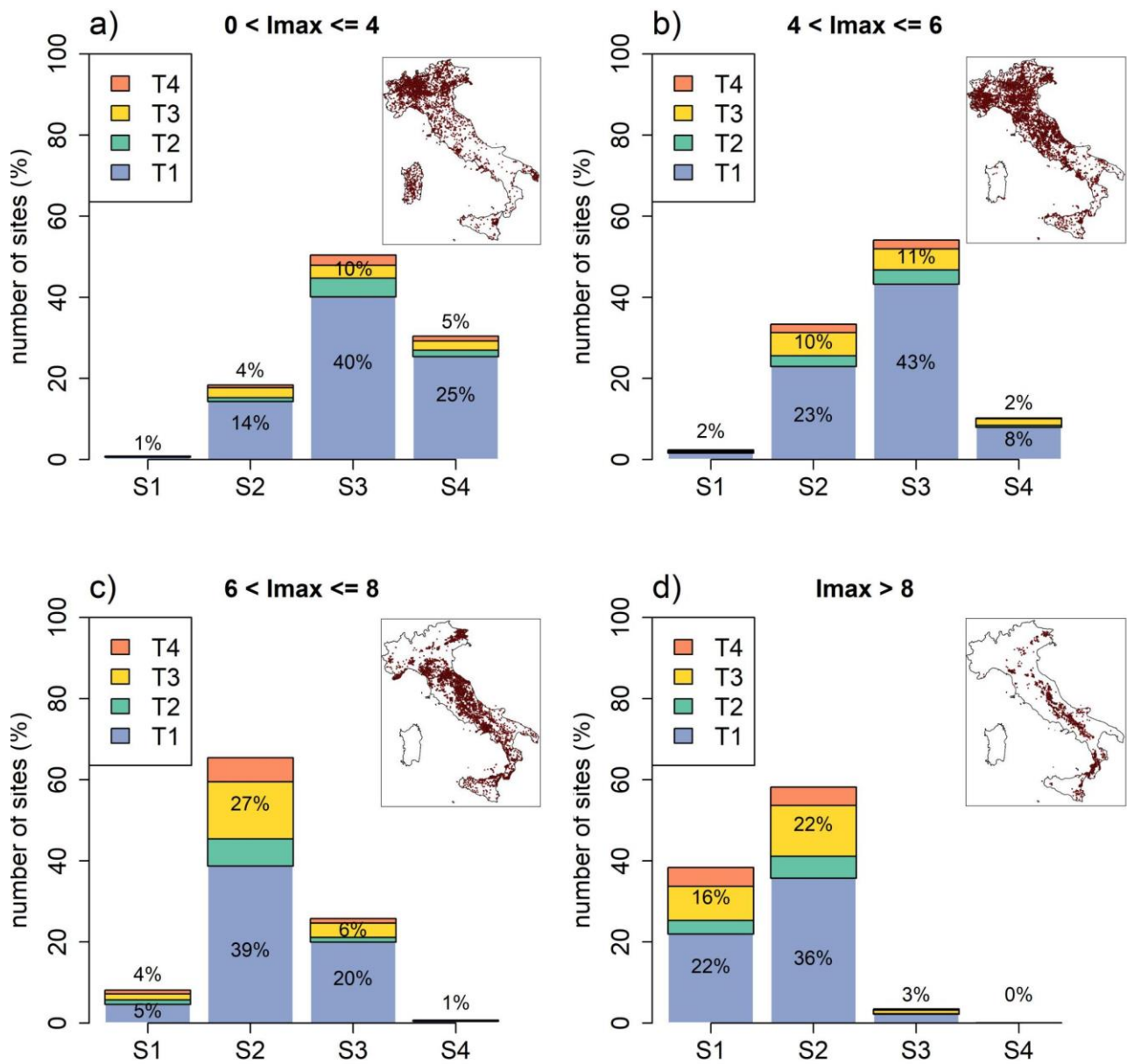
218

219 To provide an overview of the topographic classification map in Figure 6, some statistics are computed on a  
220 sample of 10,000 random points distributed homogeneously throughout the Italian territory, with a minimum  
221 interdistance of 50 m. Figure 6 shows the distribution of the topographic classes for different zones (i.e., Alps  
222 and Northern, Central, Southern Apennines). The Alps are characterized by prevalent T2 slopes (42%) and T4  
223 ridges (31%) (Figure 6a), whereas the Apennines are characterized by prevalent T1 slopes (40-50%) with  
224 equally distributed T3 and T4 ridges (both are around 15-20%) (Figure 6b, c, d). Most of the Apennines reliefs  
225 are concentrated to the North (Figure 6b), where the percentage of T2, T3 and T4 classes is larger (~60%) than  
226 in the southern sector of the Apennines (~45%) (Figure 6 c,d). Overall, about one third of the Italian territory  
227 is characterized by sites with possible topographic amplification (i.e., T2, T3 and T4 classes in Figure 6e). In  
228 detail, 25% of Italian territory is characterized by sites in T2 or T3 (topographic amplification factor of 1.2,  
229 Table 1), whereas 12% of the country area is in T4 (topographic amplification factor of 1.4, Table 1) (Figure  
230 6e).

231 In Figure 7, we show how the proposed topographic classes for the whole Italian territory are distributed with  
232 respect to the maximum observed macroseismic intensity ( $I_{max}$ , Mercalli-Cancani-Sieberg; [42]) at 15332  
233 Italian localities in DBMI15 [32] and the seismic zones in Ref. [40] (S1-S4). As expected, the localities with  
234 the lower  $I_{max}$  (Figure 7a) are concentrated in the lower seismic hazard zones (S3 and S4,  $a_g \leq 0.15$ ) of  
235 Northern Italy, moving gradually toward the higher seismic hazard zones (S1 and S2,  $a_g > 0.15$ ) of the Central-  
236 Southern Apennines for  $I_{max} > 8$  (Figure 7d). When we also consider the topographic classification of this



237 study, about 40% of the localities with the higher  $I_{max}$  ( $>6$ ) is affected by topographic amplification (T2-T4  
 238 classes, Figure 7c-d), in contrast to just about 20% for lower  $I_{max}$  ( $<6$ ) (Figure 7a-b). Moreover, we observe  
 239 a clusterization of the sites with possible topographic amplification (T2-T4 classes) in Figure 7c, with localities  
 240 characterized by medium high  $I_{max}$  (between 6 and 8), localized in seismic zone 2 ( $S2: 0.15 < a_g \leq 0.25$ ).  
 241 These results do not prove the existence of topographic amplification at those sites classified in T2-T4, but the  
 242 coherency between the topographic classification (T1-T4), the maximum observed microseismic intensity  
 243 ( $I_{max}$ ), and the seismic hazard map, represented by the seismic zones S1-S4.

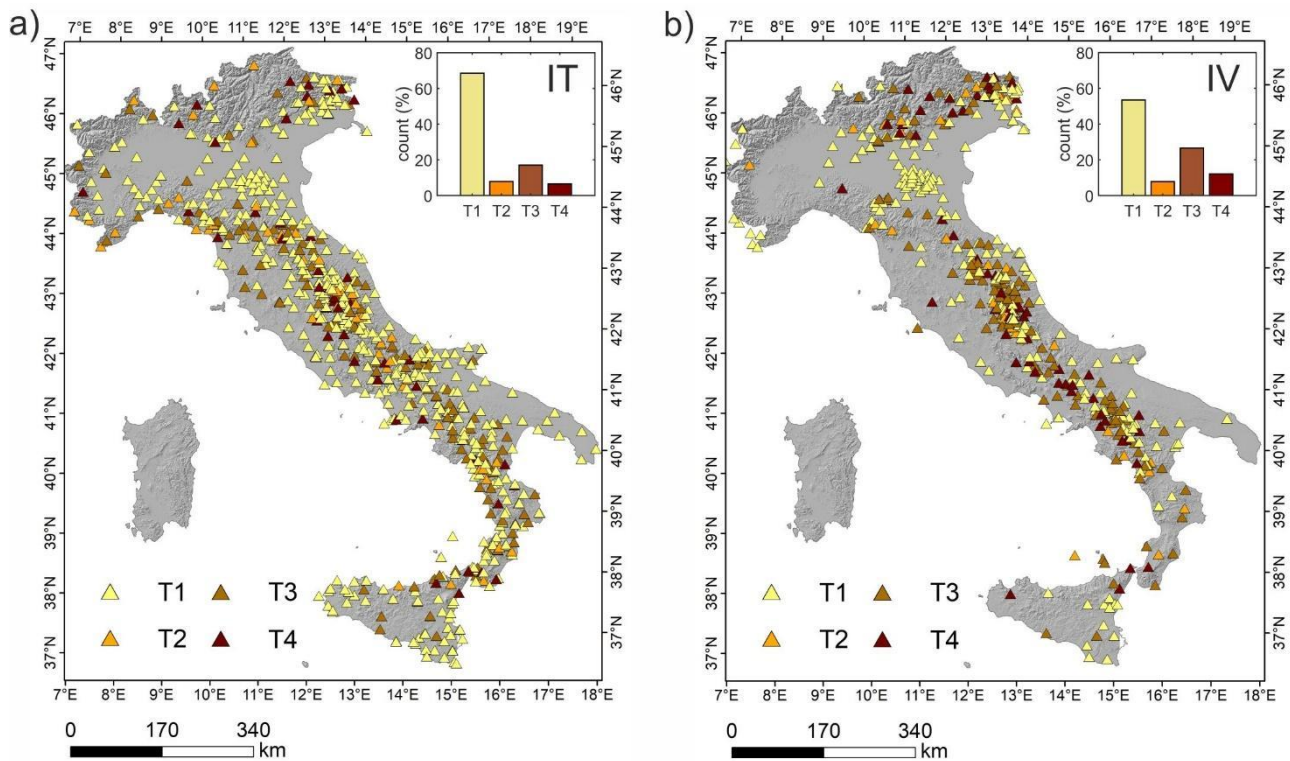


244

245 **Figure 7:** Statistics on the topographic classification of 15332 Italian localities in DBMI15 [32], grouped based  
246 on the maximum observed macroseismic intensity at a site ( $I_{max}$ , Mercalli-Cancani-Sieberg; [42]). a)  $I_{max}$   
247  $\leq 4$ , b)  $4 < I_{max} \leq 6$ , c)  $6 < I_{max} \leq 8$ , d)  $I_{max} > 8$ . The percentages are normalized to the number of sites in  
248 each group. S1-S4 are the seismic zones in Ref. [40], and T1-T4 are the topographic classes in NTC18. The  
249 geographic distribution of each dataset in panels a)-d) is reported in the insets.

250

251 Finally, we classify the seismic stations archived in the ITACA 3.1 database according to the map in Figure 6.  
252 The stations of the Italian Strong Motion Network (RAN; <http://ran.protezionecivile.it/IT/index.php>) operated  
253 by the Department of Civil Protection (IT network code), and of the Italian National Seismic Network (RSN;  
254 <http://cnt.rm.ingv.it/instruments/network/IV>), run by National Institute of Geophysics and Volcanology (IV  
255 network code), are considered in Figure 8a and Figure 8b, respectively. The accelerometric sites, classified as  
256 T1, reflect the distribution in Figure 6e (62% of Italian territory is in T1); however the percentage of T1 sites  
257 in IV and IT networks is quite different (53% versus 68%, respectively, as in Figure 8). Indeed, most seismic  
258 stations of the IT network are installed in urban areas for civil protection purposes (Figure 8a), whereas many  
259 stations of the IV network are installed outside urban areas for seismic monitoring or scientific purposes  
260 (Figure 8b). Among the other topographic classes, both networks have a majority of T3 class for those stations  
261 on the Appennine ridges (Figure 8), whereas the T2 and T4 classes are distributed between the Alps and  
262 Apennine mountain chains (Figure 8). The percentage of T2 sites in the IT and IV networks is equivalent  
263 (~7%), but the percentage of T3 (17% for IT; 26% for IV) and T4 (6% for IT; 12% for IV) sites nearly  
264 duplicates moving from IT to IV network (Figure 8).



265

266 **Figure 8:** topographic classification for the seismic stations of the ITACA 3.1 database for the RAN network

267 (a), and the RSN network (b). **IT: network code of the RAN stations; IV: network code of the RSN stations.**

268

269

#### 270 **4. Testing of the topographic classification on case studies**

271

272 The proposed topographic classification is tested at different sites for which topographic amplification is well  
 273 known (e.g. Refs. [43, 44, 10, 45, 46, 22, 23, 47]). In this study we consider Narni, Navelli, Nocera Umbra,  
 274 and Assisi (Figure 9).

275 The seismic amplification of the Narni ridge was investigated through numerical modelling and empirical  
 276 observations (e.g. Refs. [10, 46]), which proved the dependence of amplification on morphological features.

277 The experimental evidence shows an amplification of the reference motion up to 2.2-3, for the spectral  
 278 ordinates in the range 0.2-0.3 s (4-5 Hz). The largest amplifications are observed when the direction of the  
 279 ground motion is perpendicular to the main axis of the ridge. Ref. [10] describe the Narni ridge as 220-m-high,  
 280 with slopes ranging from 22° to 35°. According to the authors, the topographic class of the Narni ridge is T3

281 at the eastern end, and T4 at the western end, in agreement with the topographic classification of this study  
282 (Figure 9a).

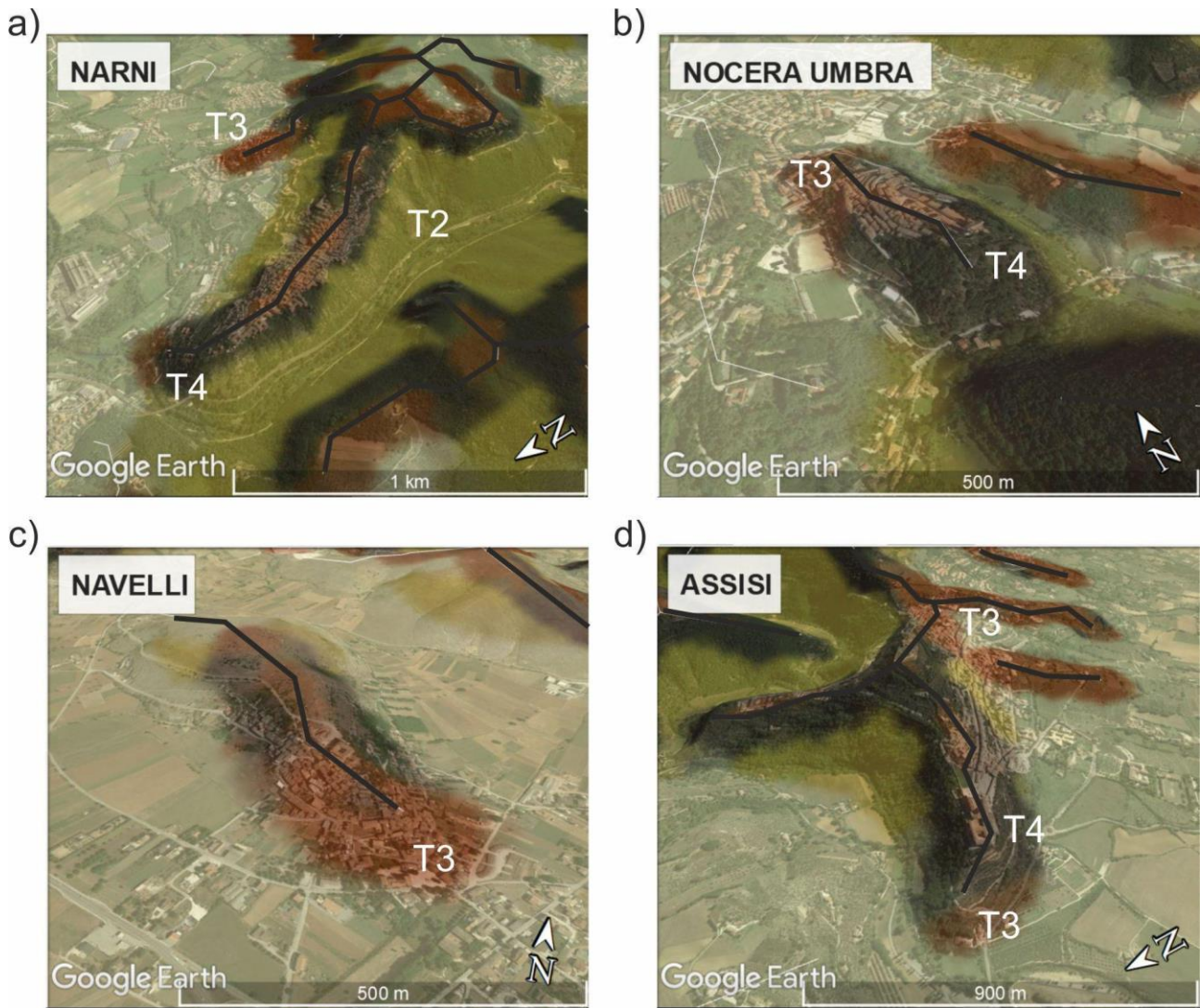
283 The Nocera Umbra hill was studied by several authors (e.g. Refs. [43-45]). The historical centre of Nocera  
284 Umbra suffered MCS intensity VII–VIII, due to the  $M_w$  5.7 and 6.0 Umbria–Marche earthquakes on 1997,  
285 September 26<sup>th</sup>. The hill is elongated in the NNW–SSE direction, with a maximum height of 574 m a.s.l., and  
286 a difference in height equal to 144 m. The northern side of the hill is quite gentle, while the southern one is  
287 very steep, therefore the topographic profile along the major axis is asymmetric with a narrow ridge. In  
288 contrast, the transverse section is more regular and symmetric, with a base width of about 400 m. Topographic  
289 amplification effects are observed at about 3.5 Hz, where the polarization is transversal to the major axis of  
290 the hill. The topographic classification map of this study (Figure 9b) reproduces the morphological features  
291 described in Ref. [45], with a T3 class at the northern end of the relief, and a T4 class at the southern end.

292 The Navelli ridge was studied by Ref. [23]. The historical centre of Navelli is located along the south-western  
293 slope of a NW-SE-trending narrow ridge, which is lithologically characterized by the outcropping crystalline  
294 limestones. For this case study, a combination of topographic and stratigraphic amplifications was observed,  
295 given a station on a rocky slope (in the historical center of Navelli) and one on a flat alluvial valley (at the base  
296 of the hill). The mean slope angle along the historical center is  $19^\circ$  and considering that this morphological  
297 ridge is much narrow at the top with respect to its base, the Navelli historical centre can be ascribed to the  
298 topographic category T2 [23]. The topographic classification map of this study reproduces the morphological  
299 features described in Ref. [23], but extends the ridge zone to the historical center of Navelli (Figure 9c) that,  
300 for this reason, is classified in T3 instead of T2 class. In this regard, the topographic classification map is  
301 conservative with respect to a more detailed observation at the site.

302 Finally, the ridge of Assisi, studied by Ref. [22], is presented. The station ASS  
303 ([http://itaca.mi.ingv.it/ItacaNet\\_31/#/station/IT/ASS](http://itaca.mi.ingv.it/ItacaNet_31/#/station/IT/ASS)) is located inside the Sacro Convento, near the historical  
304 Basilica of San Francesco, at the WNW edge of a 3D relief featuring two 2D appendixes pointing NNE and  
305 WNW. The ASS station shows an amplification at 3.5 Hz, polarized at  $25^\circ$ , perpendicularly to the main axis  
306 of the ridge. In Ref. [22] the ASS station is indicated in T3 class. The topographic classification map of this



307 study (Figure 9d) reproduces the morphological features described in Ref. [22], with a T3-T4 class for the  
308 WWN edge of the 3D relief.



309  
310 **Figure 9.** 3D view of the topographic classification map. Narni (a), Nocera Umbra (b), Navelli (c), and Assisi  
311 (d). The ridge lines are reported as black solid lines. The legend for color scale is the same as Figure 6.

312  
313

## 314 5. Conclusions

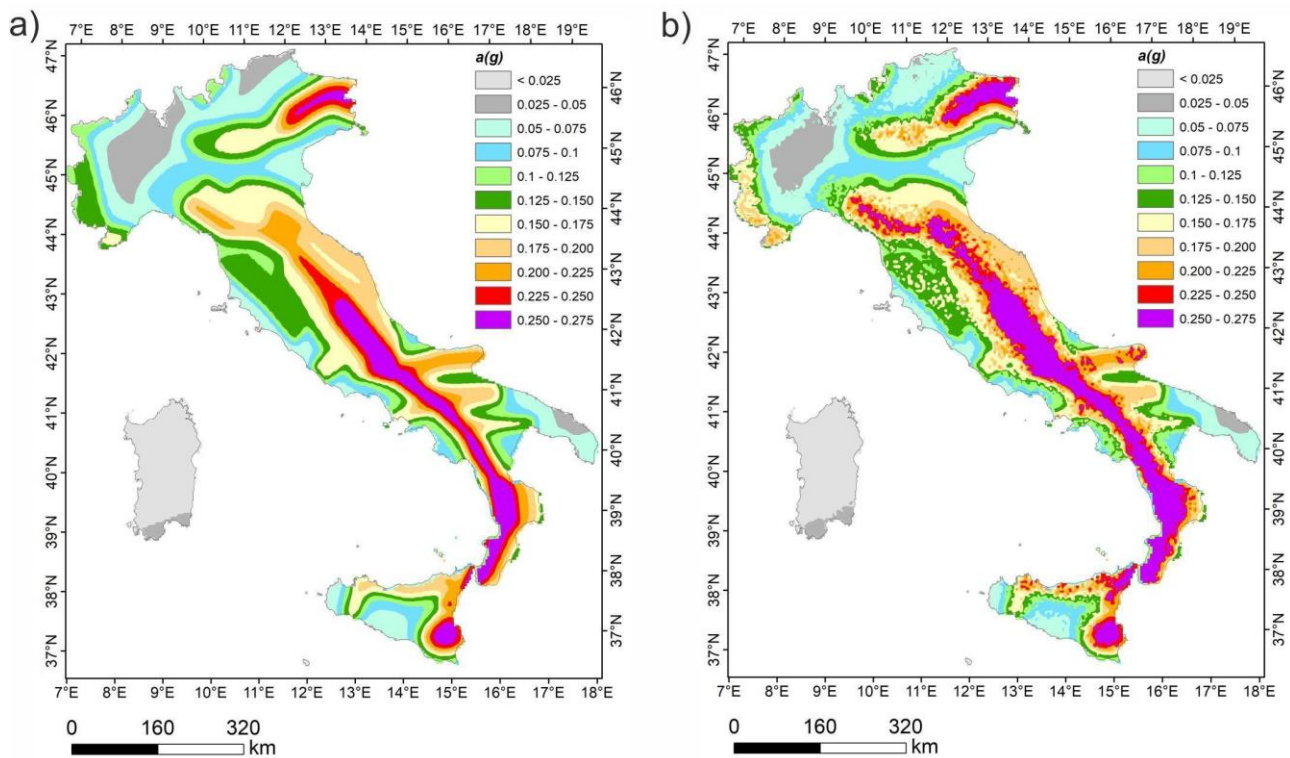
315

316 The paper describes a GIS procedure for the topographic classification of a DEM, according to the NTC18  
317 seismic code. The resulting map, with the topographic classification of Italy, is provided in the electronic  
318 supplement.

319 The GIS procedure is easily replicable: it makes use of (i) TinItaly high-resolution DEMs, (ii) the ridges  
320 identification with the Topographic Position Index (TPI; [35]), and (iii) slope angle and elevation differences  
321 calculated with built-in ESRI ArcGIS functions. The procedure allows to univocally classify the Italian  
322 territory and improves previous studies [24], where the classification was only performed in the neighbourhood  
323 of a single accelerometric station. Moreover, the results of this study do not need a manual revision, usually  
324 performed on Google Earth [48, 49, 38]. In this case, not applying quantitative and systematic morphological  
325 analyses based on digital elevation models, the topographic classification is affected by a certain level of  
326 subjectivity, related to the point of view, and to the complex geomorphology of the site. The overall check  
327 between the topographic map and the case studies reported in this work indicates a satisfactory representation  
328 of the morphological features and a correct classification of the sites. In case of mismatch, the proposed  
329 classification is generally conservative.

330 The statistics computed on the topographic classification map show that one third of the Italian territory is  
331 characterized by sites with possible topographic amplification (i.e., T2, T3 and T4 classes, Table 1). In  
332 particular, following the NTC18 and EC8 seismic codes, 25% of Italian territory is characterized by a  
333 topographic amplification factor of 1.2 (T2 and T3 classes), whereas 12% of the country area presents a  
334 topographic amplification factor of 1.4 (T4 class). The NTC18 prescribes the applicability of the topographic  
335 coefficients in case of “simple” configurations, but the criteria to discriminate between “simple” and  
336 “complex” are not specified and, for this reason, the applicability of the coefficients could be debated. Here  
337 we propose the application to the entire Italian territory, extending the classification prescribed in seismic  
338 codes for “simple” 2D relief configurations to all 3D geomorphological settings. The statistics computed on  
339 the 15332 Italian localities in DBMI [32] support the effectiveness of this extrapolation from 2D to 3D, but do  
340 not prove the existence of topographic amplification effects. Based on the topographic classification proposed  
341 in this study, about 40% of the localities that experienced the highest macroseismic intensities ( $I_{max} > 6$ ) are  
342 potentially affected by topographic amplification (T2-T4 classes). The impact of the topographic amplification  
343 (Table 1) on the hazard map [30], calculated as PGA for a return period of 475 years and soil category A  
344 (NTC18), is shown in Figure 10. Being the topographic amplification period-independent according to NTC18  
345 (Table 1), the effects are the same at all spectral periods.

346 The map of the topographic classification of Italy can be used for both scientific and application purposes. In  
347 this regard, the topographic classification map may be used for the characterization of seismic stations, and  
348 may support the identification of sites with systematic topographic effects. The map is also useful for the  
349 selection of accelerometric time histories that meet the spectrum-compatibility requirement indicated in the  
350 Italian (NTC18) and European (EC8) seismic codes. In the latter case, it allows selecting those stations in T1  
351 class, whose recordings are not affected by topographic amplification. Moreover, the proposed map may guide  
352 future investigations of T2, T3 and T4 sites, in particular to study the period-dependent topographic  
353 amplifications, as opposed to the constant amplification factors provided in NTC18 and EC8 seismic codes.  
354



355  
356 **Figure 10:** a) probabilistic seismic hazard map for PGA, soil class A and topographic class T1 (return period  
357 of 475 years), as in Ref. [30]; b) application of topographic amplification factors according to the classification  
358 shown in Figure 6.

359  
360  
361  
362

363 **Acknowledgment**

364 The study has partially benefited from funding provided by the Italian Presidenza del Consiglio dei Ministri–  
365 Dipartimento della Protezione Civile (DPC)—Agreement B2, DPC-INGV. This paper does not necessarily  
366 represent DPC official opinions and policies.

367

368 **CRediT authorship contribution statement**

369 **Claudia Mascandola:** Conceptualization, Methodology, Software, Writing - original draft, Writing - review  
370 & editing, Data curation, Formal analysis. **Lucia Luzi:** Conceptualization, Methodology, Software, Writing -  
371 original draft, Writing - review & editing, Data curation, Formal analysis. **Chiara Felicetta:** Writing - original  
372 draft, Writing - review & editing, Data curation, Formal analysis. **Francesca Pacor:** Writing - original draft,  
373 Writing - review & editing, Data curation, Formal analysis.

374

375 **References**

376 [1] Boore, D. M. (1972). A note on the effect of simple topography on seismic SH waves, *Bull. Seismol. Soc.*  
377 *Am.* 62, 275–284.

378 [2] Bard, P. Y. (1982). Diffracted waves and displacement fields over two dimensional elevated topographies,  
379 *Geophys. J. Int.* 71, 731–760.

380 [3] Panizza, M. (1991). Geomorphology and seismic risk. *Earth-Sci Rev*; 31:11–20.

381 [4] Paolucci, R. (2002). Amplification of earthquake ground motion by steep topographic irregularities. *Earthq.*  
382 *Eng. Struct. Dyn.*; 31:1831–53.

383 [5] Griffiths, D.W., and Bollinger, G.A. (1979). The effect of Appalachian Mountain topography on seismic  
384 waves. *Bull. Seism. Soc. Am.*; 69:1081–105.

385 [6] Kawase, H, and Aki, K. (1990). Topography effect at the critical SV-wave incidence: possible explanation  
386 of damage pattern by the Whittier Narrows, California, earthquake of 1 October 1987. *Bull. Seism. Soc. Am.*;  
387 80:1–22.



- 388 [7] Marsan, P, Milana, G, Pugliese, A, and Sanò, T. (2000). Local amplification effects recorded by a local  
389 strong motion network during the 1997 Umbria-Marche earthquake. In: Proceedings of 12th world conference  
390 on earthquake engineering, paper no. 1046. Auckland, New Zealand.
- 391 [8] Marra, F., Azzara, R., Bellucci, F., Caserta, A., Cultrera, G., Mele, B., Palombo, B., Rovelli, A., and  
392 Boschi, E. (2000). Large amplification of ground motion at rock sites within a fault zone in Nocera Umbra  
393 (central Italy). *J Seismol.*;4:543–54.
- 394 [9] Cara, F., Rovelli, A., Di Giulio, G., Marra, F., Braun, T., Cultrera, G., Azzara R., and Boschi E.(2005).  
395 The role of site effects on the intensity anomaly of San Giuliano di Puglia inferred from aftershocks of the  
396 Molise, Central Southern Italy, sequence, *Bull. Seism. Soc. Am.*;95:1457–68.
- 397 [10] Massa, M., Lovati, S., D’Alema, E., Ferretti, G., and Bakavoli, M. (2010). An experimental approach for  
398 estimating seismic amplification effects at the top of a ridge, and the implication for ground-motion  
399 predictions: the case of Narni, Central Italy. *Bull. Seism. Soc. Am.*; 100(6): 286–301.
- 400 [11] Buech, F., Davies, T.R., and Pettinga, J.R. (2010). The little red hill seismic experimental study:  
401 topographic effects on ground motion at a bedrock-dominated mountain Edifice. *Bull. Seism. Soc.*  
402 *Am.*;100:2219–29.
- 403 [12] Marzorati, S., Ladina, C., Falcucci, E., Gori, S., Saroli, M., Ameri, G., and Galadini F. (2011). Site effects  
404 on the rock: the case of Castelvechio Subequo (L’Aquila,central Italy). *Bull. Earthq. Eng.*;9:841–68.
- 405 [13] Paolucci, R., Faccioli, E., and Maggio, F. (1999). 3D Response analysis of an instrumented hill at  
406 Matsuzaki, Japan, by a spectral method. *J. Seismol.*;3:191–209.
- 407 [14] Gazetas, G., Kallou, P.V., and Psarropoulos, P.N. (2002). Topography and soil effects in the Ms 5.9  
408 Parnitha (Athens) earthquake: the case of Adámes. *Nat. Hazards*; 27:133–69.

- 409 [15] Assimaki, D., Gazetas, G., and Kausel, E.. (2005). Effects of local soil conditions and topographic  
410 aggravation seismic motion: parametric investigation and recorded field evidence from the 1999 Athens  
411 earthquake. *Bull. Seism. Soc. Am.*; 95:1059–89.
- 412 [16] Scandella, L, Davì, M, Paolucci, R, and Pessina, V. (2008). Numerical simulation and Gis mapping of  
413 seismic amplification due to topographic effects. In: *Proceedings of 31st European Seismological Commission*  
414 (ESC). Crete, Greece.
- 415 [17] Faccioli, E. (Eds.) (1986). *Elementi Per Una Guida Alle Indagini Di Microzonazione sismica – Progetto*  
416 *finalizzato “Geodinamica”, Quaderni de “La ricerca scientifica”, n.114, CNR, 7, 72–82, (in Italian).*
- 417 [18] Faccioli, E., Vanini, M., Frassinè, L. (2002). “Complex” site effects in earthquake ground motion,  
418 including topography. In: *Proceedings of 12th European Conference on earthquake engineering*. Paperno. 844.
- 419 [19] Galli, P, and Molin, D. (2004). Macro seismic Survey of the 2002 Molise, Italy, Earthquake and Historical  
420 Seismicity of San Giuliano di Puglia. *Earthq. Spectra*; 20(S1):39–52.
- 421 [20] Tertulliani, A, Arcoraci, L, Berardi, M, Bernardini, F, Camassi, R, Castellano, C, Del Mese, S., Ercolani,  
422 E., Graziani, L., Leschiutta, I., Rossi A., and Vecchi M. (2011). An application of EMS98 in a medium-sized  
423 city: the case of L’Aquila (Central Italy) after the April 6, 2009 Mw 6.3 earthquake. *Bull. Earthq. Eng*; 9:67–  
424 80.
- 425 [21] Gizzi, F.T., Potenza, M.R., and Zotta, C. (2012). 3 November 1980 Irpinia-Basilicata earthquake  
426 (Southern Italy): towards a full knowledge of the seismic effects. *Bull. Earth. Eng*; 10:1109–31.
- 427 [22] Cauzzi, C., Fäh, D., Pessina, V., Faccioli, E., and Smerzini, C. (2012). Topographic amplification from  
428 recorded earthquake data and numerical simulations. In *15th World Conference on Earthquake Engineering*.
- 429 [23] Gallipoli, M. R., Bianca, M., Mucciarelli, M., Parolai, S., and Picozzi, M. (2013). Topographic versus  
430 stratigraphic amplification: mismatch between code provisions and observations during the L’Aquila (Italy,  
431 2009) sequence. *Bull. of Earth. Eng.*, 11(5), 1325-1336.

- 432 [24] Pessina, V., and Fiorini, E. (2014). A GIS procedure for fast topographic characterization of seismic  
433 recording stations. *Soil Dynamics and Earthquake Engineering*, 63, 248-258.
- 434 [25] NTC (2018). Ministero delle Infrastrutture e dei Trasporti. Aggiornamento delle Norme Tecniche per le  
435 Costruzioni. Part 3.2.2: Categorie di sottosuolo e condizioni topografiche, *Gazzetta Ufficiale* n. 42 del 20  
436 febbraio 2018 (in Italian).
- 437 [26] CEN (Comite Europe en de Normalisation), 2004. Eurocode 8. Design Provisions for Earthquake  
438 Resistance of Structures—Part 5: Foundations, Retaining Structures and Geotechnical Aspects. ENV 1998-5,  
439 CEN European Committee for Standardisation, Brussels, 1994.
- 440 [27] NTC (2008). Ministero delle Infrastrutture e dei Trasporti. Nuove Norme Tecniche per le Costruzioni.  
441 Part 3: Categorie di sottosuolo e condizioni topografiche, *Gazzetta Ufficiale*, no. 29, del 4 febbraio 2008 (in  
442 Italian).
- 443 [28] Luzi, L., Hailemikael, S., Bindi, D., Pacor, F., Mele, F., and Sabetta, F. (2008). ITACA (ITalian  
444 ACcelerometric Archive): A Web Portal for the Dissemination of Italian Strong-motion Data, *Seismological*  
445 *Research Letters*, 79(5), 716–722.
- 446 [29] Pacor, F., Paolucci, R., Luzi, L., Sabetta, F., Spinelli, A., Gorini, A., Nicoletti, M., Marcucci, S., Filippi,  
447 L., and Dolce M. (2011), Overview of the Italian strong motion database ITACA 1.0, *Bull. Earth. Eng.*, 9(6),  
448 1723–1739.
- 449 [30] MPS Working Group (2004). Redazione della mappa di pericolosità sismica prevista dall’Ordinanza  
450 PCM del 20 marzo 2003, *Rapporto Conclusivo per il Dipartimento della Protezione Civile* (in Italian).
- 451 [31] Stucchi, M., C. Meletti, V. Montaldo, H. Crowley, G. M. Calvi, and E. Boschi (2011). Seismic hazard  
452 assessment (2003–2009) for the Italian building code, *Bull. Seismol. Soc. Am.* 101, no. 4, 1885–1911.
- 453 [32] Locati, M., Camassi, R., Rovida, A., Ercolani, E., Bernardini, F., Castelli, V., Caracciolo, C.H.,  
454 Tertulliani, A., Rossi, A., Azzaro, R., D’Amico, S., Conte, S., Rocchetti, E., Antonucci, A. (2019). Database

455 Macrosismico Italiano (DBMI15), versione 2.0. Istituto Nazionale di Geofisica e Vulcanologia (INGV).  
456 <https://doi.org/10.13127/DBMI/DBMI15.2>

457 [33] D'Amico, M., Felicetta, C., Russo, E., Sgobba, S., Lanzano, G., Pacor, F., and Luzi, L. (2020). Italian  
458 Accelerometric Archive v 3.1 - Istituto Nazionale di Geofisica e Vulcanologia, Dipartimento della Protezione  
459 Civile Nazionale. doi: 10.13127/itaca.3.1

460 [34] Tarquini, S., Isola, I., Favalli, M., Mazzarini, F., Bisson, M., Pareschi, M.T., and Boschi, E. (2007).  
461 TINITALY/01: a new triangular irregular network of Italy. *Annals of Geophysics* 50,407–425.

462 [35] Weiss, A. (2001). Topographic position and landform analysis. Poster presentation, ESRI Users  
463 Conference, San Diego, CA.

464 [36] Burrough, P. A., and McDonell, R. A., (1998). *Principles of Geographical Information Systems* (Oxford  
465 University Press, New York), 190 pp.

466 [37] Jenness, J., Brost, B., and Beier, P. (2013). Land facet corridor designer. USDA forest service rocky  
467 mountain research station.

468 [38] Mascandola, C., Massa, M., Lovati, S., and Augliera, P. (2018). The site characterization scheme of the  
469 INGV strong motion database (ISMD): Overview and site classification. *Seism. Res. Lett.*, 89(1), 86-98.

470 [39] Zhan, C. (1993). A Hybrid Line Thinning Approach. *Proceedings Auto-Carto 11*, Minneapolis, pp. 396-  
471 405.

472 [40] OPCM n. 3519 (2006): criteri generali per l'individuazione delle zone sismiche e per la formazione e  
473 l'aggiornamento degli elenchi delle stesse zone; *Gazzetta Ufficiale* n. 108 dell'11 maggio 2006 (in Italian).

474 [41] ISTAT, 2011. 154° Censimento della Popolazione e delle Abitazioni. [http://www.istat.it/it/censimento-](http://www.istat.it/it/censimento-popolazione/censimento-popolazione-2011)  
475 [popolazione/censimento-popolazione-2011](http://www.istat.it/it/censimento-popolazione/censimento-popolazione-2011).

476 [42] Sieberg, A. (1930). The earthquake geology (*Geologie der Erdbeben*). *Handbuch der Geophysik*  
477 2(4):550–555 (in German).

- 478 [43] Caserta, A., Bellucci, F., Cultrera, G., Donati, S., Marra, F., Mele, G., Palombo, B. and Rovelli, A. (2000).  
479 Study of site effects in the area of Nocera Umbra (Central Italy) during the 1997 Umbria-Marche seismic  
480 sequence, *J. Seismol.*, 4, 555–565.
- 481 [44] Donati, S., Marra, F. and Rovelli, A. (2001). Damage and ground shaking in the town of Nocera Umbra  
482 during Umbria-Marche, central Italy, earthquakes: the special effect of a fault zone, *Bull. seism. Soc. Am.*, 91,  
483 511–519.
- 484 [45] Pischietta, M., Cultrera, G., Caserta, A., Luzi, L., and Rovelli, A. (2010). Topographic effects on the hill  
485 of Nocera Umbra, central Italy. *Geophysical Journal International*, 182(2), 977-987.
- 486 [46] Lovati, S., Bakavoli, M. K. H., Massa, M., Ferretti, G., Pacor, F., Paolucci, R., Haghshenas E., and  
487 Kamalian, M. (2011). Estimation of topographical effects at Narni ridge (Central Italy): comparisons between  
488 experimental results and numerical modelling. *Bull. of Earth. Eng.*, 9(6), 1987-2005.
- 489 [47] Pischietta, M., Cianfarra, P., Salvini, F., Cara, F., and Vannoli, P. (2018). A systematic analysis of  
490 directional site effects at stations of the Italian seismic network to test the role of local topography. *Geophysical*  
491 *Journal International*, 214(1), 635-650.
- 492 [48] Pessina, V. (2010). Appendix E. EC8 subsoil and topographic classification of ITACA stations  
493 Deliverable D10, Project S4 The Italian strong motion database.  
494 [http://esse4.mi.ingv.it/files/images/stories/d10\\_appendix%20e.pdf](http://esse4.mi.ingv.it/files/images/stories/d10_appendix%20e.pdf)
- 495 [49] Di Capua, G., Lanzo, G., Pessina, V., Peppoloni, S., and Scasserra, G. (2011). The recording stations of  
496 the Italian strong motion network: Geological information and site classification. *Bulletin of Earthquake*  
497 *Engineering*, 9(6), 1779-1796.



29 mountain [1]; the focusing of seismic waves near the crest because of the reflection on a free surface and/or  
30 the interaction between incident and diffracted waves [2]. Moreover, topographic seismic amplification is  
31 considered as one of the possible causes of triggers of earthquake induced phenomena, like landslides, which  
32 represent one of the major causes of earthquake-related devastation (e.g. Refs. [3,4]). Several studies over the  
33 last four decades have discussed the effects of surface topography on seismic ground motions from  
34 instrumental investigations [5-12] and numerical models [13, 14, 4, 15, 16]. Relevant macroseismic case  
35 histories are reported by several authors [17, 18, 4, 19, 20, 21]. The proper assessment of topographic  
36 amplification is of key importance in the European context, where a broad range of exposure levels to  
37 earthquake hazard is encompassed [18] and many small towns are located at the top of isolated steep cliffs.  
38 Such unfavourable conditions often come together with high vulnerability of the buildings and the  
39 susceptibility to slope sliding [4]. Therefore, the topographic amplification, even if generally do not exceed  
40 the stratigraphic one (e.g. Refs. [4, 22, 23, 24]), plays a significant role in seismic design, microzonation studies  
41 and hazard analyses.

42 The current Italian (NTC, 2018; hereinafter NTC18 [25]) and European (CEN, 2004; hereinafter EC8 [26])  
43 seismic codes prescribe frequency-independent amplification factors not larger than 40%, based on the height  
44 and the slope of simple 2D relief configurations. The approach, introduced in the EC8, was further adopted by  
45 NTC (2008) [27] and the current NTC18, where four classes of topographic amplification are explicitly  
46 defined, as listed in Table 1. These categories refer to bidimensional configurations, such as elongated crests  
47 and ridges, and must be considered only when their height is greater than 30m.

48

49

50

51

52

53

54 **Table 1:** Topographic amplification factors prescribed by Italian seismic code (NTC18).

Topographic Class	Description	Amplification factor
T1	Flat surface, isolated slopes and cliffs with average slope angle $i \leq 15^\circ$	1
T2	Slopes with average slope angle $i > 15^\circ$	1.2
T3	Relief with ridge width much smaller than the base and average slope angle $15^\circ < i \leq 30^\circ$	1.2
T4	Relief with ridge width much smaller than the base and average slope angle $i > 30^\circ$	1.4

55

56 Following the classification of the Italian seismic code, Ref. [24] first implemented a procedure for the  
 57 topographic characterization of seismic recording stations archived in ITACA 2.0 database (Italian  
 58 Accelerometric Archive, [http://itaca20.mi.ingv.it/ItacaNet\\_20/](http://itaca20.mi.ingv.it/ItacaNet_20/); [28, 29]). The procedure is based on  
 59 morphometric analyses of high-resolution digital elevation models (DEMs), with the support of Geographic  
 60 Information Systems (GIS). In the specific, the procedure computes the slope and ridge layers from DEM,  
 61 besides different statistical indicators (i.e., maximum, minimum, sum, and mean) within a circular buffer area  
 62 around each recording station. These statistical indicators are combined to provide both a slope and a ridge  
 63 index that are further matched to define the topographic classification of the seismic station. However, some  
 64 combinations of the indices do not permit a univocal topographic classification because their arrangement may  
 65 lead to different topographic classes (in that case the site is not classified). Moreover, the procedure  
 66 implemented for the ridge detection is not straightforward and implies several noise removal steps to enhance  
 67 significant features and remove small topographic irregularities. Since the procedure is applied to the  
 68 surroundings of each seismic station, it needs to be run each time a new seismic station is added to the database.  
 69 These downsides led to the development of a new GIS-based procedure for topographic classification of the  
 70 whole Italian territory. However, the NTC18 prescribes the applicability of the topographic coefficients in case  
 71 of “simple” configurations, but the criteria to discriminate between “simple” and “complex” are not specified  
 72 and, for this reason, the applicability of the coefficients could be debated.



73 In this study, we propose a step-like procedure and provide some statistics of the resulting map, correlating it  
74 to the probabilistic seismic hazard [30, 31], the macroseismic intensity data in DBMI15 [32], and the seismic  
75 stations in the Italian Accelerometric Archive 3.1 (ITACA 3.1; [http://itaca.mi.ingv.it/ItacaNet\\_31/#/home](http://itaca.mi.ingv.it/ItacaNet_31/#/home);  
76 [33]). The topographic classification is tested at several sites, where topographic amplification has been  
77 observed. Finally, based on the proposed classification, we evaluate the impact of the topographic  
78 amplification factors (Table 1), on the probabilistic seismic hazard map calculated for Peak Ground  
79 Acceleration (PGA), a return period of 475 years, and stiff soil i.e. soil class EC8-A [30].

80

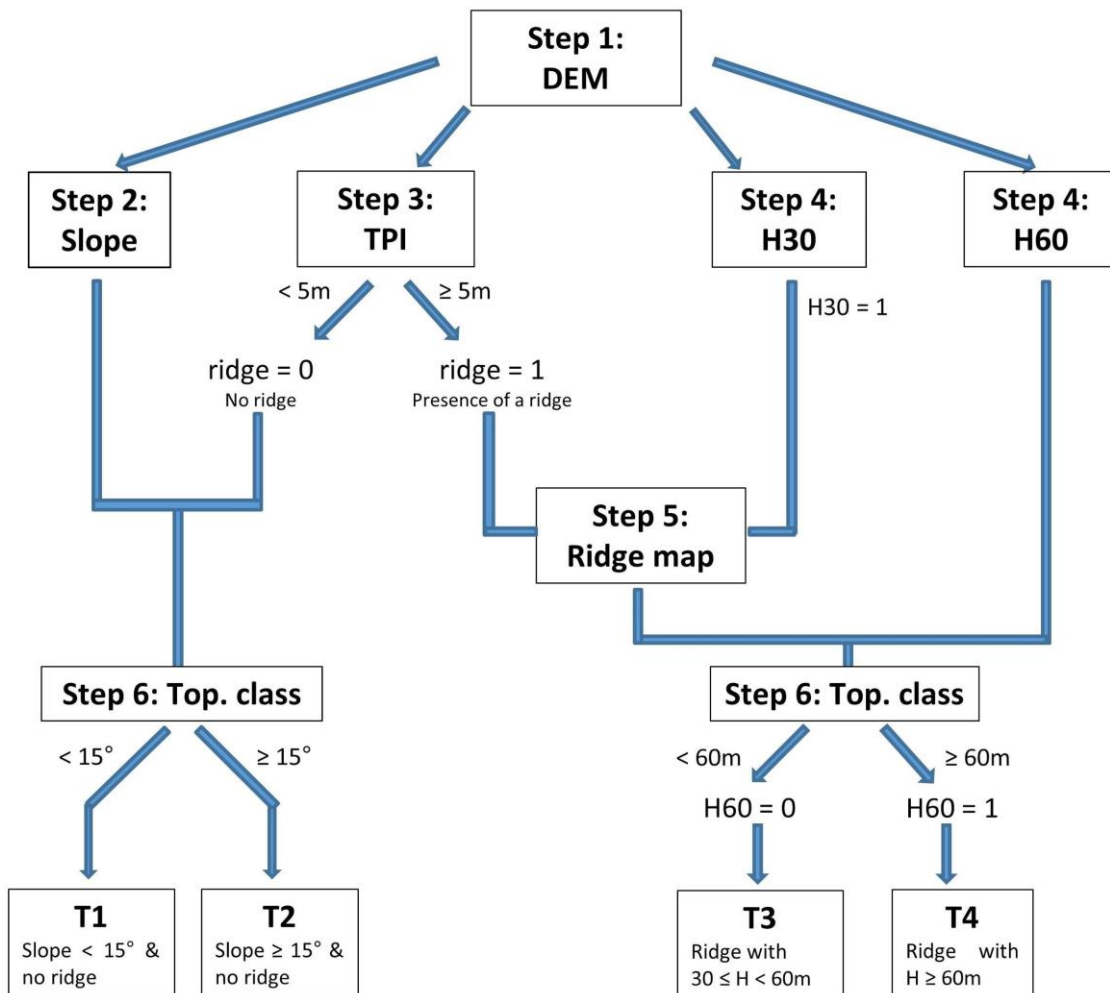
## 81 **2. Method**

82

83 Digital elevation models (DEMs) are fundamental in environmental and morphological analyses. In this study,  
84 the topographic classification is applied to the so called *TinItaly DEM* [34], created for the whole Italian  
85 territory as a triangular irregular network (TIN), and converted into grid format of 10m cell size, according to  
86 a tiled structure composed of 193, 50-km side square elements. Ref. [34] carried out a comprehensive  
87 assessment of the accuracy of the TinItaly DEM, finding a root mean square error in elevation (RMSEz)  
88 between 0.8 and 6.0 m. For our analysis, we select the TinItaly DEM since, to date, it is the model with the  
89 highest resolution, freely available for scientific purposes for the whole Italian territory.

90 The morphological analyses applied on the DEM are summarized in Figure 1 and described in detail in the  
91 following subparagraphs. Starting from DEM (Figure 1, step 1), the slope angle is computed (step 2) along  
92 with the Topographic Position Index (TPI; [35], step 3) and the elevation range (maximum minus minimum  
93 elevation value in a moving window), applying a threshold height of 30 m (H30) and 60 m (H60), respectively  
94 (step 4). Since the positive TPI values allow detecting the ridge zones, an empirical threshold of 5 m is adopted  
95 to discriminate between ridges (ridge is assigned 1) and non-ridges (ridge is assigned 0). Considering the slope  
96 and TPI values, a first discrimination of topographic classes can be performed (Figure 1). Whereas T1 and T2  
97 sites are on slopes (ridge = 0), T3 and T4 sites should be on ridges (ridge = 1). Therefore, T1 and T2 categories  
98 can be identified by the combination of ridge = 0, and slope angle  $< 15^\circ$  or  $\geq 15^\circ$ , respectively (step 6); T3  
99 and T4 classes can be identified by ridge = 1 (Figure 1, step 5), on reliefs with difference in height  $> 30$  m. To  
100 further discriminate between T3 and T4 classes, the H60 map is considered. When H60 = 1 (elevation range

101  $H$  is between 30 and 60 m) the site is in class T3 (step 6). In the following, the step-like procedure applied to  
 102 the TinItaly DEM is described in detail, indicating the ESRI ArcGis ([www.esri.com](http://www.esri.com)) tools adopted for the  
 103 morphological analyses.  
 104



105  
 106 **Figure 1:** Workflow for topographic classification.

107  
 108 **Step 1: preliminary DEM processing**

109 To improve the computation time and data handling, the 193 TinItaly DEM tiles are resampled with a 40 m  
 110 resolution. Afterwards, the DEM tiles are merged to obtain a unique map for the Italian territory. Finally, to  
 111 remove minor irregularities, a smoothing is performed by applying a mean algorithm with a 3x3 moving  
 112 window (e.g. *Focal Statistics* tool in ESRI ArcGis), as shown in Figure 2a for a sample area in central  
 113 Apennines.

114

### 115 **Step 2: computation of topographic slope**

116 To compute the slope angle, we use the *Slope* function implemented on ESRI ArcGis (Figure 2b). The slope  
117 is computed in degree units in a 3 by 3 cell neighborhood (i.e., moving window) on the smoothed DEM,  
118 considering the planar method with the average maximum technique [36]. The slope is measured as the  
119 maximum rate of change in value from a cell to its immediate neighbors. The calculation is performed on a  
120 projected flat plane using a 2D Cartesian coordinate system. The processed cell is judged reliable if the  
121 computation is performed on at least seven cells in the 3 by 3 neighborhood.

122

### 123 **Step 3: ridge extraction**

124 For the ridge detection, we adopt the Topographic Position Index (TPI) algorithm proposed by Ref. [35]. This  
125 algorithm is implemented in the *Land Facet Analysis* toolbox [37] in ESRI ArcGis. The TPI is defined as a  
126 difference between the actual value of the elevation  $H$  at given location  $[x, y]$  and mean elevation of the  
127 neighboring area  $A$ :

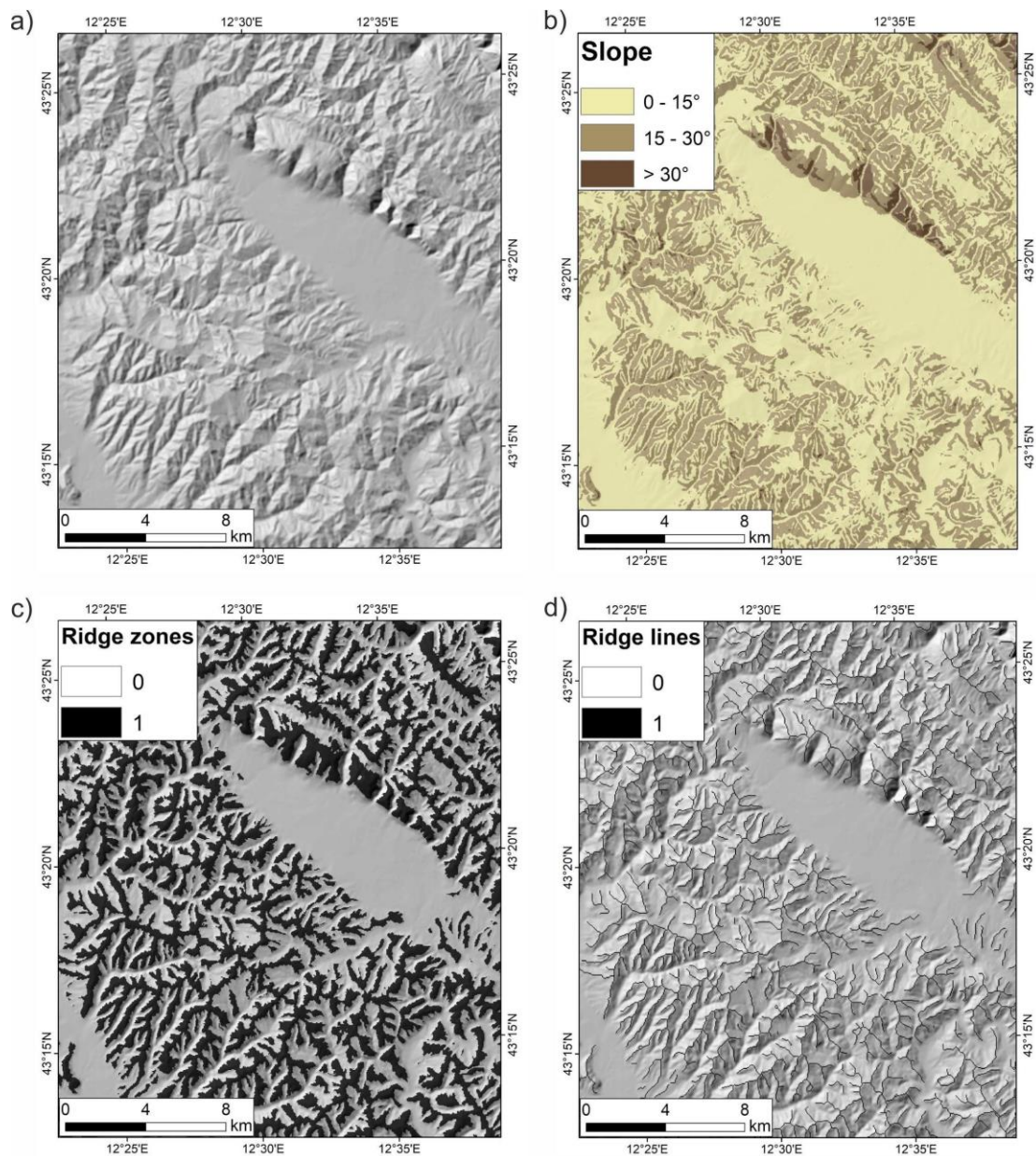
128

$$129 \quad TPI(x, y) = H(x, y) - \int_A H(x, y) \partial S / \int_A \partial S \quad \text{Eq. (1)}$$

130

131 The area  $A$  should be centered at the point  $[x, y]$ . Both the size and the shape of the neighboring area have to  
132 be defined in advance, depending on the specific application. In general, there is no constraint on both values.  
133 The TPIs reflect the differences between the elevation in a particular cell and the mean elevation of the  
134 surrounding cells: positive values mean the cell is higher than its surroundings, while negative values mean it  
135 is lower. Therefore, high positive values of TPI characterize ridgetops and hilltops, while negative values  
136 define basins or valley bottoms. TPI values close to zero indicate flat or mid-slope areas. This kind of analysis  
137 strictly depends on the scale used to count the surrounding cells. To set the interval for the TPI analysis, we  
138 carry out some tests, considering different radii (i.e. 300 m, 500 m and 800 m). For the Italian case, a  
139 neighborhood circle with a radius of 500 m is adopted, since it allows better discriminating the ridges out of  
140 the main reliefs, while ignoring the minor deflections [38]. The ridge zones are obtained from the TPI map  
141 (*Raster Calculator* tool) considering a threshold value equal to 5. The output is a binary raster that takes the  
142 value 1 when  $TPI > 5$  and 0 otherwise (Figure 2c). Since TPI units are in meters, a TPI value of 5 means that

143 this particular cell is 5 m higher than the average elevation of its neighborhood (i.e., a radius of 500 m). The  
144 TPI threshold is selected empirically. A lower threshold (down to 1) does not significantly affect the final  
145 results. The ridgelines are retrieved by applying a thinning function (*Thin* tool) that thins rasterized linear  
146 features by reducing the number of cells representing the width of the features [39], as shown in Figure 2d.

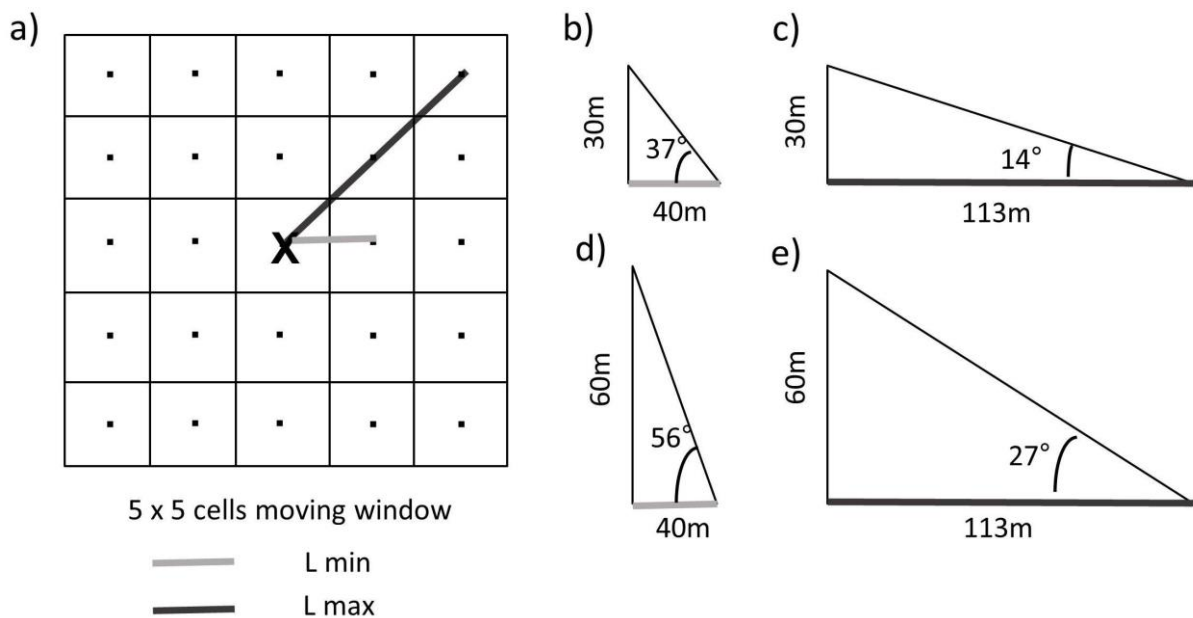


147  
148 **Figure 2:** Extraction for an area of the Central Apennines of a) DEM, represented as hillshade; b) topographic  
149 slope; c) ridge zones setting TPI > 5; and d) ridge lines after thinning.

150  
151 **Step 4: detection of the difference in height to extract topographic categories T3 and T4**

152 In the EC8 and NTC18 seismic codes, reliefs higher than 30 m are considered to have potential topographic  
153 site effects (i.e., T3 or T4 topographic classes in NTC18). To individuate these reliefs, we compute the

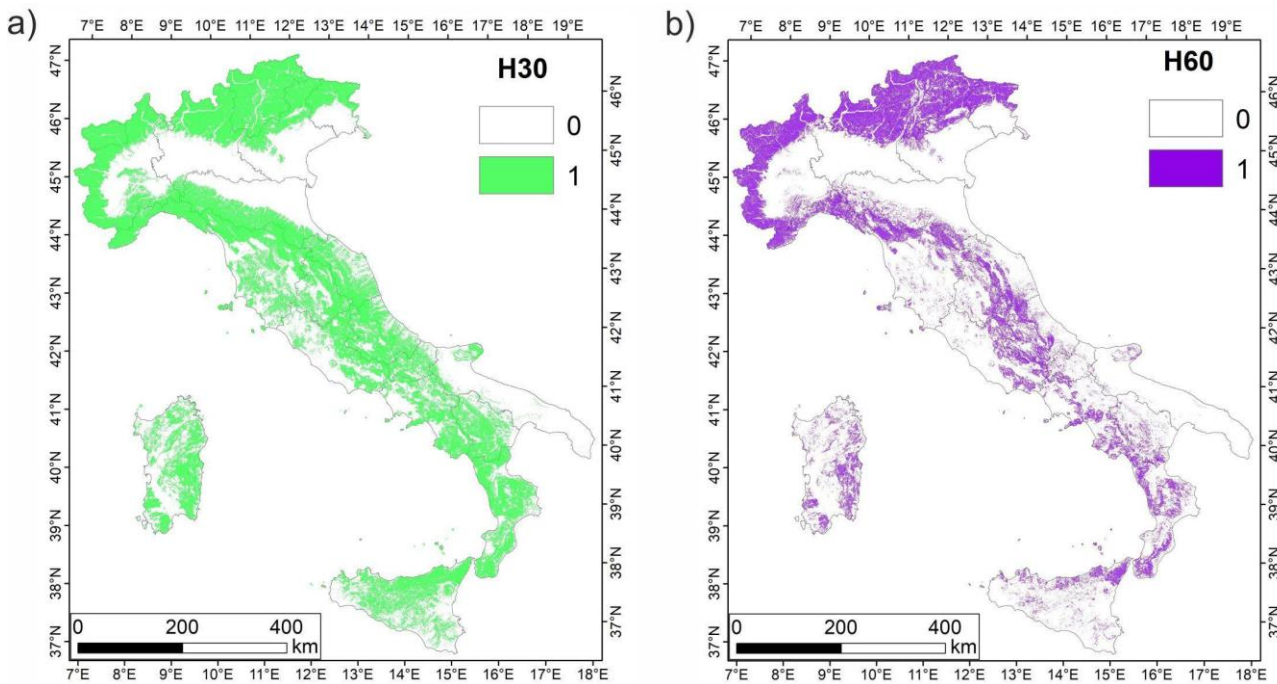
154 elevation range, i.e., maximum minus minimum elevation value, in 5x5 cells moving windows (Figure 3a;  
 155 *Focal Statistics* tool). Considering that the DEM cell size is 40 m, the maximum and minimum semi-transverse  
 156 axis of a relief spanned by a 5x5 moving window is 113 m (e.g.  $80 \cdot \sqrt{2}$ m), as exemplified by the dark grey  
 157 line in Figure 3, and 40 m (light grey line in Figure 3), respectively. Considering a threshold difference in  
 158 height of 30 m (condition for topographic amplification), we obtain the slope angle interval, which ranges  
 159 between  $14^\circ$  and  $37^\circ$  (Figure 3b,c), thus including both T3 and T4 classes. To discriminate class T4, we  
 160 consider the map representing the difference in height of 60 m that individuates the slope angles in the interval  
 161  $27^\circ$ - $56^\circ$  (Figure 3d,e), including T4 sites (i.e, ridges with height greater than 30 m and average slope angle  $i >$   
 162  $30^\circ$ ). In the following sections, the elevation range computed with a threshold height of 30 m is indicated as  
 163 H30; whereas the elevation range computed with a threshold height of 60 m is indicated as H60 (Figure 4).



164  
 165  
 166 **Figure 3:** Trigonometry applied to H30 and H60 maps. a) 5x5 cells moving window with the smaller (light  
 167 grey line) and larger (dark grey line) semi-transverse axis spanned in the moving window. b)-c) maximum and  
 168 minimum slope angles, considering a difference in height of 30 m, d)-e) maximum and minimum slope angles  
 169 considering a difference in height of 60 m.

170





171

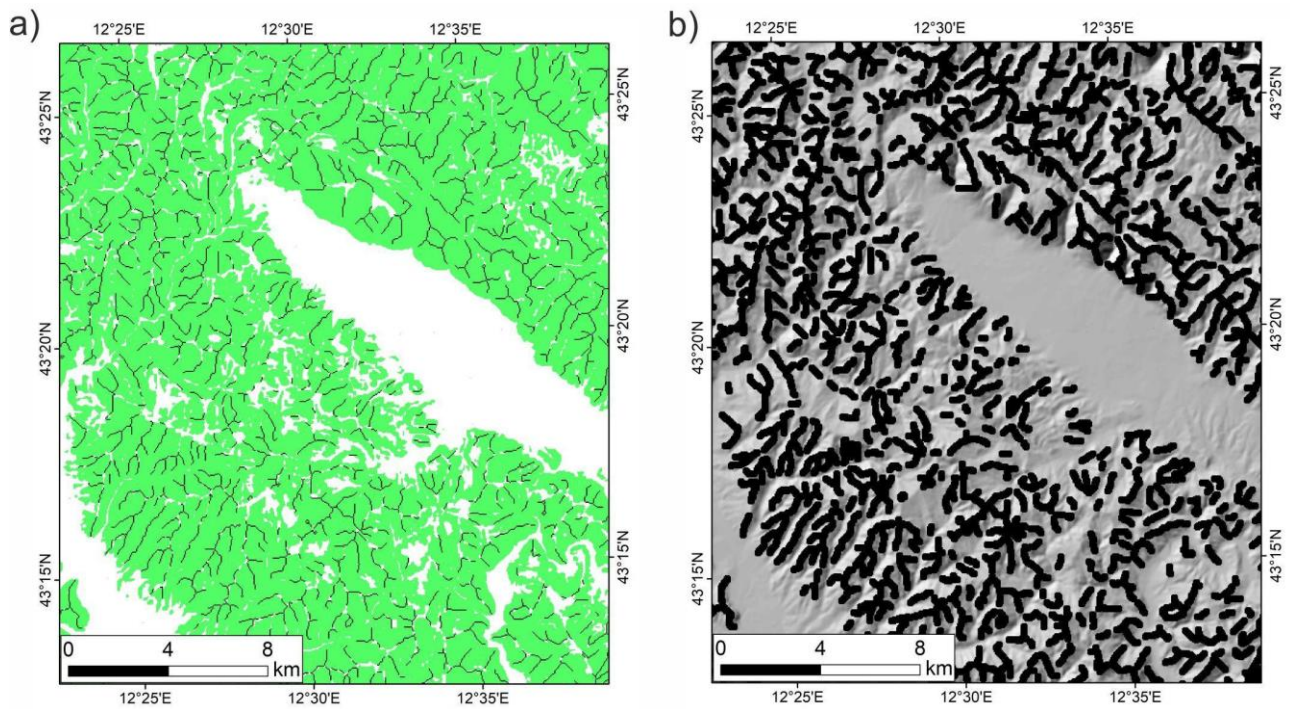
172

173 **Figure 4:** Map of a) difference in height of 30 m (H30) and b) 60 m (H60) for the Italian territory.

174

175 **Step 5: identification of T3 and T4 topographic categories**

176 After the identification of the difference in height to extract T3 and T4 topographic categories, the ridge lines  
 177 of the corresponding reliefs can be obtained by overlapping the ridge line map (evaluated in step 3) and the  
 178 H30 map (evaluated in step 4) (*Raster calculator* tool, imposing  $H30 = 1$  & ridge lines = 1), as shown in Figure  
 179 5a. To identify the sparse groups of points (< 5 cells), the *Region Group* tool is combined with the *Set Null*  
 180 operator. The *Region Group* counts the number of grouped cells within an 8-cells neighborhood (excluding 0  
 181 values), and the *Set Null* sets to null the cell values for counts less than 5. The sparse groups of points are  
 182 finally removed with the *Nibble* operator that allows filtering of the ridge line raster. Finally, the extracted  
 183 ridges are dilated (3 cells per side), using the *ArcScan* extension (Figure 5b). Being the raster cell size 40 m,  
 184 the ridge lines are dilated by 120 m per side to approximate the semi-transverse axis of a relief with a minimum  
 185 30 m of difference in height and a consequent minimum slope angle of 15°, which are the thresholds for  
 186 topographic amplification according to NTC18 and EC8 seismic codes (see Figure 3c as an example). In this  
 187 sense, the 120 m per side is conservative because it may include the whole relief down to the base.



188

189 **Figure 5:** a) intersection between ridge line map (black lines) and H30 map; b) ridge map on DEM represented  
 190 as hillshade.

191 **Step 6: topographic classification**

192 Once the identification of the morphological parameters is performed, the final topographic classification is  
 193 obtained by combining the following maps:

- 194 1. slope angles (*slope* map);  
 195 2. ridges of reliefs with potential topographic amplification (*ridge* map);  
 196 3. reliefs with average slope angle  $i > 30^\circ$  (*H60* map);

197 as reported in Table 2 (*Raster calculator* tool). The final topographic classification is reported in Figure 6 and  
 198 provided in the electronic supplement.

199

200

201

202

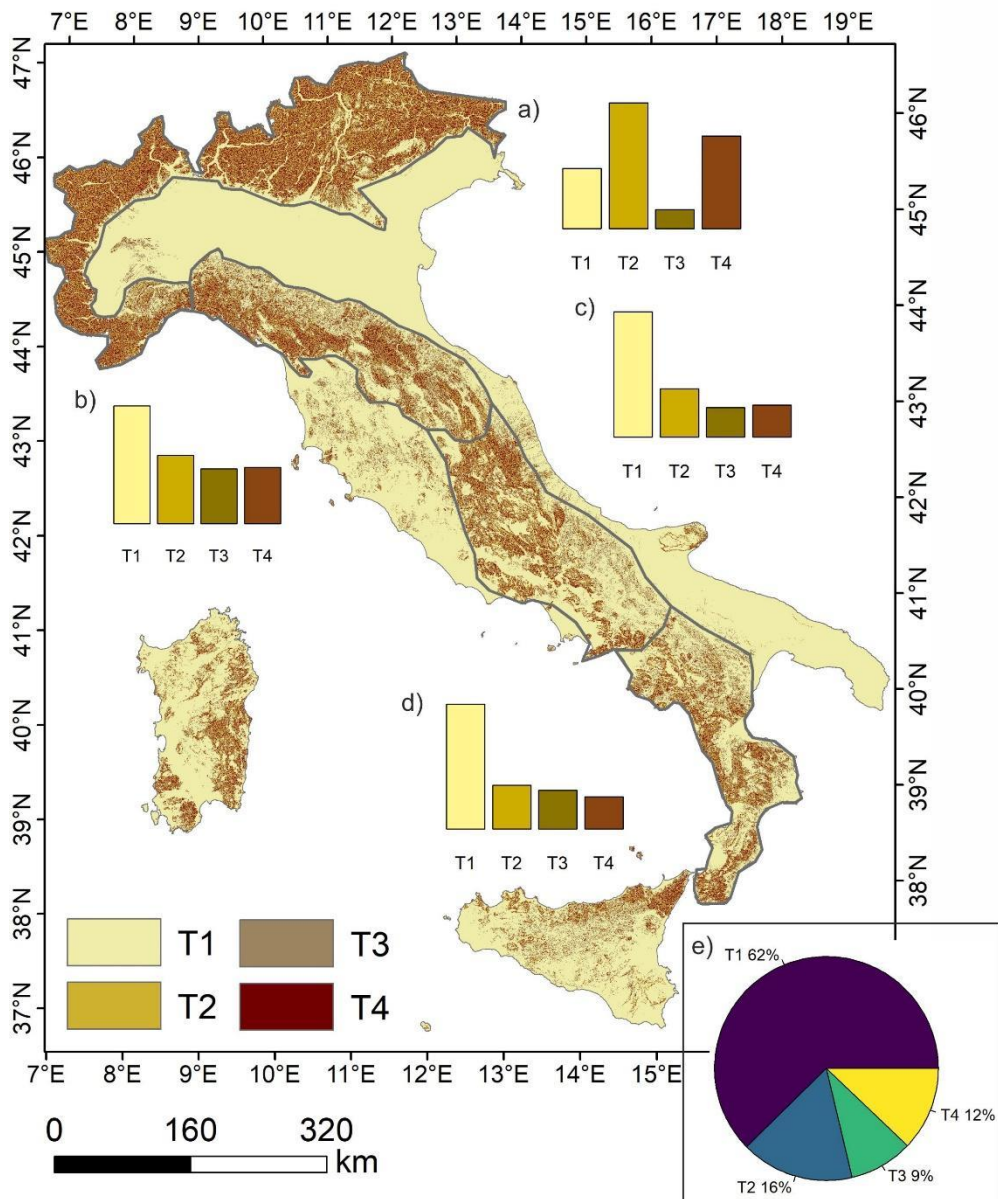
203

204 **Table 2:** Topographic classification (NTC18), based on the proposed GIS-procedure.

205

Condition	Class	Description
ridge = 0 & slope < 15°	T1	slope with average slope angle $i < 15^\circ$
ridge = 0 & slope $\geq 15^\circ$	T2	slope with average slope angle $i \geq 15^\circ$
ridge = 1 & H60 = 0	T3	ridge with difference in height in the range 30-60 m (i.e. base angle $\leq 30^\circ$ )
ridge = 1 & H60 = 1	T4	ridge with difference in height > 60 m (i.e. base angle > 30°)

206



207

208



209 **Figure 6:** Topographic classification of Italy according to the current Italian seismic code (NTC18), obtained  
210 following the proposed procedure. The histograms of the topographic classes are reported for the Alps (a) and  
211 the Apennine (b, c, d) mountain chains, in the Northern, Central, and Southern sectors respectively (grey  
212 outline). e) Pie chart of topographic classes (T1-T4) for the whole Italian territory. The statistics are computed  
213 on a sample of 10,000 random points distributed homogeneously throughout the Italian territory. The count is  
214 normalized to the number of points in each sector.

215

216

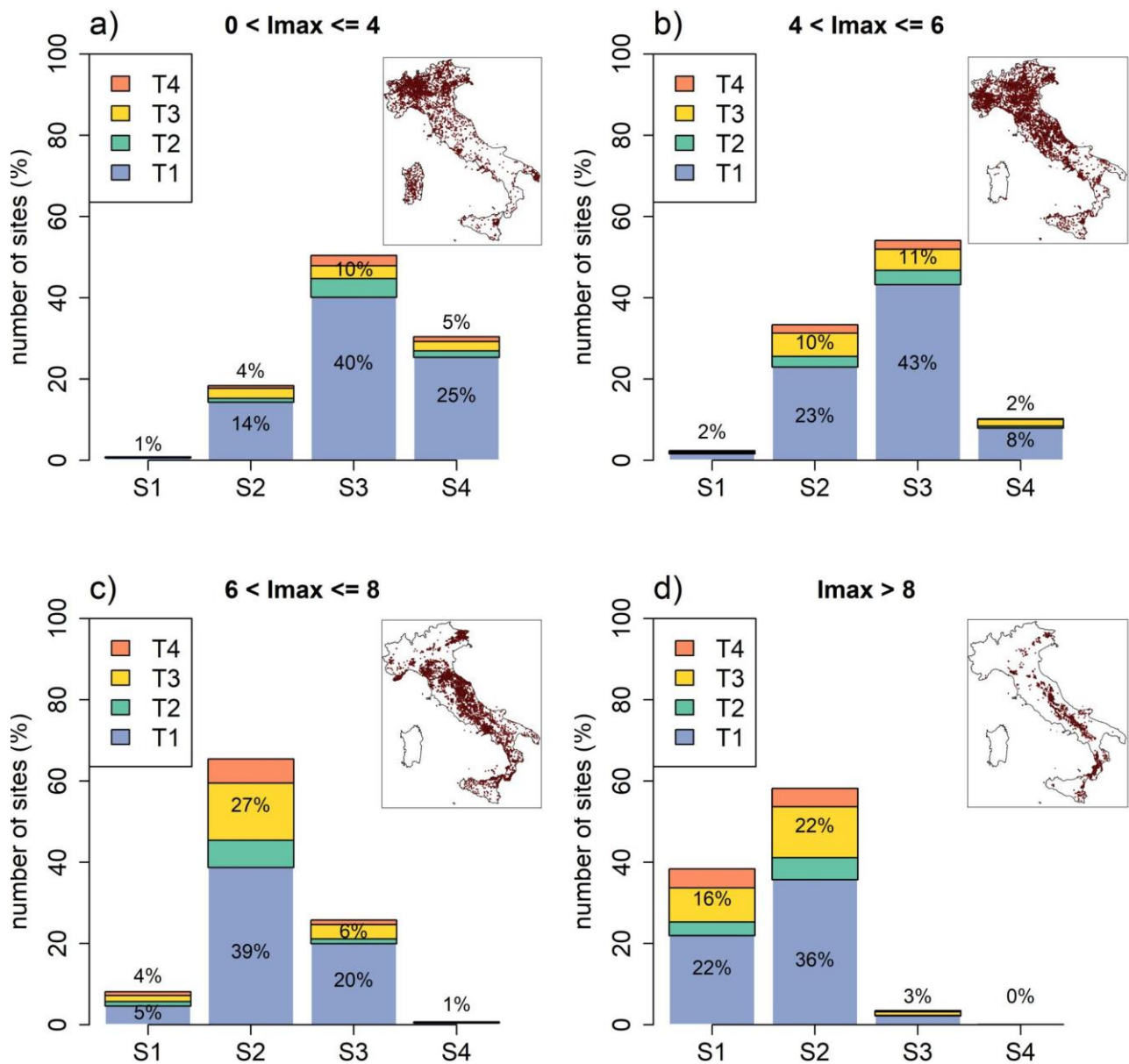
### 217 **3. Statistics on the topographic classification map**

218

219 To provide an overview of the topographic classification map in Figure 6, some statistics are computed on a  
220 sample of 10,000 random points distributed homogeneously throughout the Italian territory, with a minimum  
221 interdistance of 50 m. Figure 6 shows the distribution of the topographic classes for different zones (i.e., Alps  
222 and Northern, Central, Southern Apennines). The Alps are characterized by prevalent T2 slopes (42%) and T4  
223 ridges (31%) (Figure 6a), whereas the Apennines are characterized by prevalent T1 slopes (40-50%) with  
224 equally distributed T3 and T4 ridges (both are around 15-20%) (Figure 6b, c, d). Most of the Apennines reliefs  
225 are concentrated to the North (Figure 6b), where the percentage of T2, T3 and T4 classes is larger (~60%) than  
226 in the southern sector of the Apennines (~45%) (Figure 6 c,d). Overall, about one third of the Italian territory  
227 is characterized by sites with possible topographic amplification (i.e., T2, T3 and T4 classes in Figure 6e). In  
228 detail, 25% of Italian territory is characterized by sites in T2 or T3 (topographic amplification factor of 1.2,  
229 Table 1), whereas 12% of the country area is in T4 (topographic amplification factor of 1.4, Table 1) (Figure  
230 6e).

231 In Figure 7, we show how the proposed topographic classes for the whole Italian territory are distributed with  
232 respect to the maximum observed macroseismic intensity ( $I_{max}$ , Mercalli-Cancani-Sieberg; [42]) at 15332  
233 Italian localities in DBMI15 [32] and the seismic zones in Ref. [40] (S1-S4). As expected, the localities with  
234 the lower  $I_{max}$  (Figure 7a) are concentrated in the lower seismic hazard zones (S3 and S4,  $a_g \leq 0.15$ ) of  
235 Northern Italy, moving gradually toward the higher seismic hazard zones (S1 and S2,  $a_g > 0.15$ ) of the Central-  
236 Southern Apennines for  $I_{max} > 8$  (Figure 7d). When we also consider the topographic classification of this

237 study, about 40% of the localities with the higher  $I_{max}$  ( $>6$ ) is affected by topographic amplification (T2-T4  
 238 classes, Figure 7c-d), in contrast to just about 20% for lower  $I_{max}$  ( $<6$ ) (Figure 7a-b). Moreover, we observe  
 239 a clusterization of the sites with possible topographic amplification (T2-T4 classes) in Figure 7c, with localities  
 240 characterized by medium high  $I_{max}$  (between 6 and 8), localized in seismic zone 2 (S2:  $0.15 < a_g \leq 0.25$ ).  
 241 These results do not prove the existence of topographic amplification at those sites classified in T2-T4, but the  
 242 coherency between the topographic classification (T1-T4), the maximum observed microseismic intensity  
 243 ( $I_{max}$ ), and the seismic hazard map, represented by the seismic zones S1-S4.

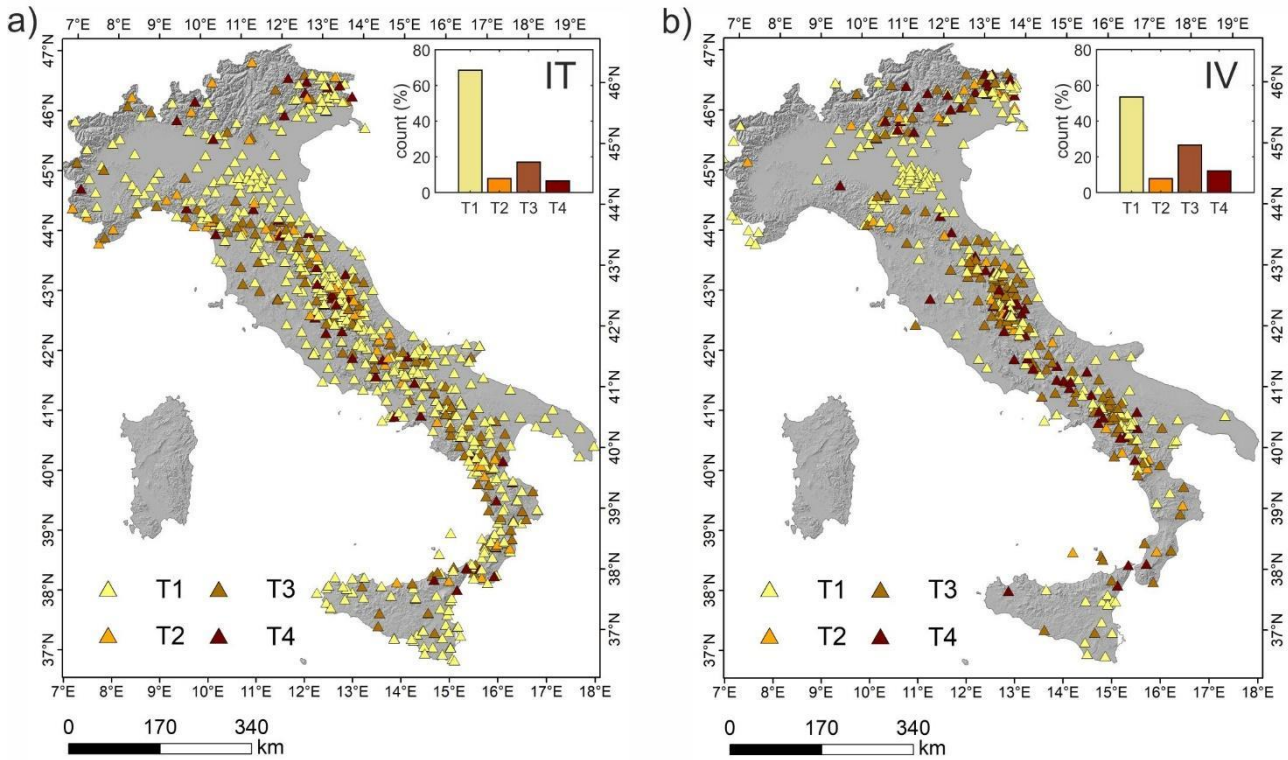


244

245 **Figure 7:** Statistics on the topographic classification of 15332 Italian localities in DBMI15 [32], grouped based  
246 on the maximum observed macroseismic intensity at a site ( $I_{max}$ , Mercalli-Cancani-Sieberg; [42]). a)  $I_{max}$   
247  $\leq 4$ , b)  $4 < I_{max} \leq 6$ , c)  $6 < I_{max} \leq 8$ , d)  $I_{max} > 8$ . The percentages are normalized to the number of sites in  
248 each group. S1-S4 are the seismic zones in Ref. [40], and T1-T4 are the topographic classes in NTC18. The  
249 geographic distribution of each dataset in panels a)-d) is reported in the insets.

250

251 Finally, we classify the seismic stations archived in the ITACA 3.1 database according to the map in Figure 6.  
252 The stations of the Italian Strong Motion Network (RAN; <http://ran.protezionecivile.it/IT/index.php>) operated  
253 by the Department of Civil Protection (IT network code), and of the Italian National Seismic Network (RSN;  
254 <http://cnt.rm.ingv.it/instruments/network/IV>), run by National Institute of Geophysics and Volcanology (IV  
255 network code), are considered in Figure 8a and Figure 8b, respectively. The accelerometric sites, classified as  
256 T1, reflect the distribution in Figure 6e (62% of Italian territory is in T1); however the percentage of T1 sites  
257 in IV and IT networks is quite different (53% versus 68%, respectively, as in Figure 8). Indeed, most seismic  
258 stations of the IT network are installed in urban areas for civil protection purposes (Figure 8a), whereas many  
259 stations of the IV network are installed outside urban areas for seismic monitoring or scientific purposes  
260 (Figure 8b). Among the other topographic classes, both networks have a majority of T3 class for those stations  
261 on the Appennine ridges (Figure 8), whereas the T2 and T4 classes are distributed between the Alps and  
262 Apennine mountain chains (Figure 8). The percentage of T2 sites in the IT and IV networks is equivalent  
263 ( $\sim 7\%$ ), but the percentage of T3 (17% for IT; 26% for IV) and T4 (6% for IT; 12% for IV) sites nearly  
264 duplicates moving from IT to IV network (Figure 8).



265

266 **Figure 8:** topographic classification for the seismic stations of the ITACA 3.1 database for the RAN network  
 267 (a), and the RSN network (b). IT: network code of the RAN stations; IV: network code of the RSN stations.

268

269

#### 270 **4. Testing of the topographic classification on case studies**

271

272 The proposed topographic classification is tested at different sites for which topographic amplification is well  
 273 known (e.g. Refs. [43, 44, 10, 45, 46, 22, 23, 47]). In this study we consider Narni, Navelli, Nocera Umbra,  
 274 and Assisi (Figure 9).

275 The seismic amplification of the Narni ridge was investigated through numerical modelling and empirical  
 276 observations (e.g. Refs. [10, 46]), which proved the dependence of amplification on morphological features.

277 The experimental evidence shows an amplification of the reference motion up to 2.2-3, for the spectral  
 278 ordinates in the range 0.2-0.3 s (4-5 Hz). The largest amplifications are observed when the direction of the  
 279 ground motion is perpendicular to the main axis of the ridge. Ref. [10] describe the Narni ridge as 220-m-high,  
 280 with slopes ranging from 22° to 35°. According to the authors, the topographic class of the Narni ridge is T3

281 at the eastern end, and T4 at the western end, in agreement with the topographic classification of this study  
282 (Figure 9a).

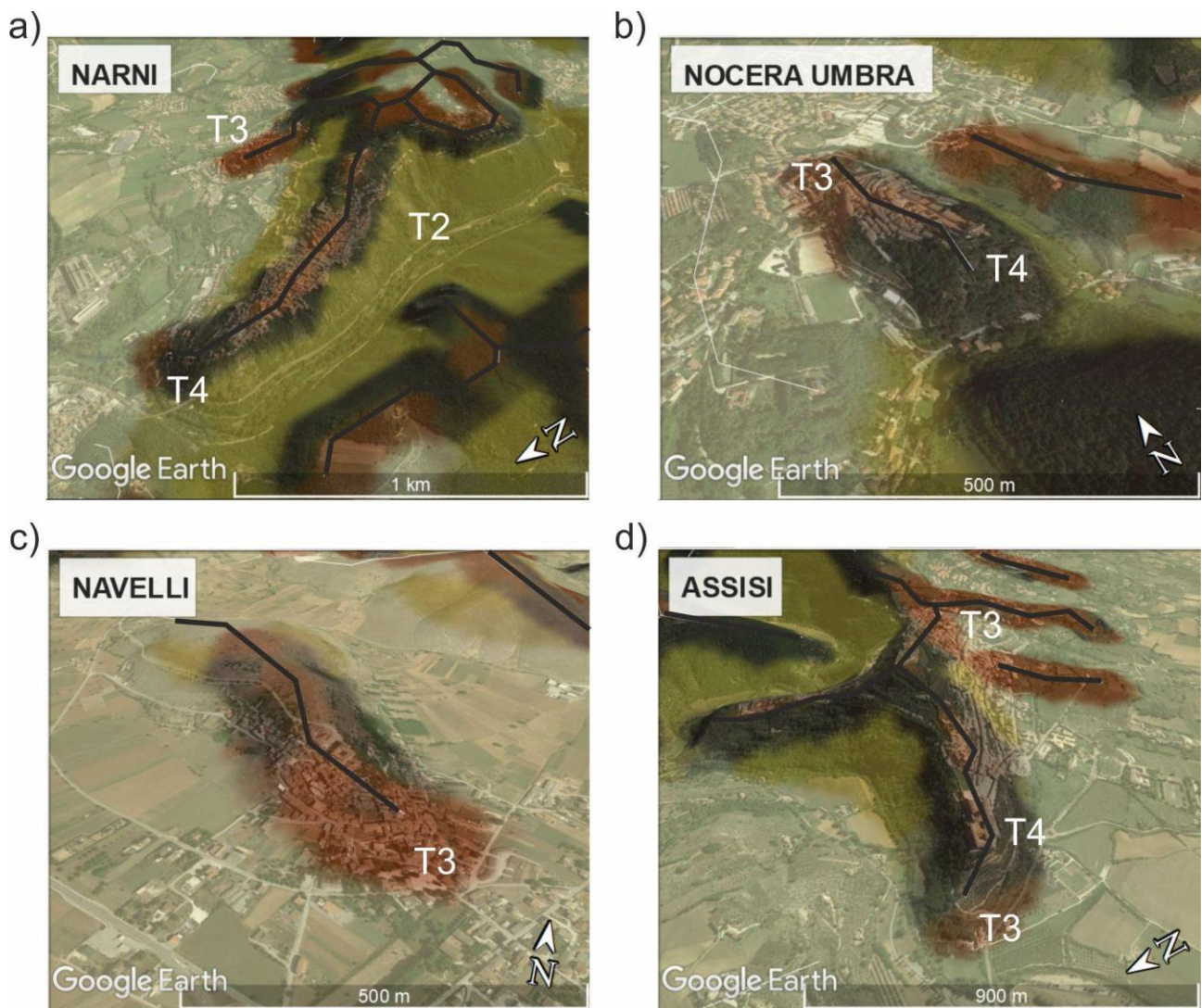
283 The Nocera Umbra hill was studied by several authors (e.g. Refs. [43-45]). The historical centre of Nocera  
284 Umbra suffered MCS intensity VII–VIII, due to the  $M_w$  5.7 and 6.0 Umbria–Marche earthquakes on 1997,  
285 September 26<sup>th</sup>. The hill is elongated in the NNW–SSE direction, with a maximum height of 574 m a.s.l., and  
286 a difference in height equal to 144 m. The northern side of the hill is quite gentle, while the southern one is  
287 very steep, therefore the topographic profile along the major axis is asymmetric with a narrow ridge. In  
288 contrast, the transverse section is more regular and symmetric, with a base width of about 400 m. Topographic  
289 amplification effects are observed at about 3.5 Hz, where the polarization is transversal to the major axis of  
290 the hill. The topographic classification map of this study (Figure 9b) reproduces the morphological features  
291 described in Ref. [45], with a T3 class at the northern end of the relief, and a T4 class at the southern end.

292 The Navelli ridge was studied by Ref. [23]. The historical centre of Navelli is located along the south-western  
293 slope of a NW-SE-trending narrow ridge, which is lithologically characterized by the outcropping crystalline  
294 limestones. For this case study, a combination of topographic and stratigraphic amplifications was observed,  
295 given a station on a rocky slope (in the historical center of Navelli) and one on a flat alluvial valley (at the base  
296 of the hill). The mean slope angle along the historical center is  $19^\circ$  and considering that this morphological  
297 ridge is much narrow at the top with respect to its base, the Navelli historical centre can be ascribed to the  
298 topographic category T2 [23]. The topographic classification map of this study reproduces the morphological  
299 features described in Ref. [23], but extends the ridge zone to the historical center of Navelli (Figure 9c) that,  
300 for this reason, is classified in T3 instead of T2 class. In this regard, the topographic classification map is  
301 conservative with respect to a more detailed observation at the site.

302 Finally, the ridge of Assisi, studied by Ref. [22], is presented. The station ASS  
303 ([http://itaca.mi.ingv.it/ItacaNet\\_31/#/station/IT/ASS](http://itaca.mi.ingv.it/ItacaNet_31/#/station/IT/ASS)) is located inside the Sacro Convento, near the historical  
304 Basilica of San Francesco, at the WNW edge of a 3D relief featuring two 2D appendixes pointing NNE and  
305 WNW. The ASS station shows an amplification at 3.5 Hz, polarized at  $25^\circ$ , perpendicularly to the main axis  
306 of the ridge. In Ref. [22] the ASS station is indicated in T3 class. The topographic classification map of this



307 study (Figure 9d) reproduces the morphological features described in Ref. [22], with a T3-T4 class for the  
308 WWN edge of the 3D relief.



309  
310 **Figure 9.** 3D view of the topographic classification map. Narni (a), Nocera Umbra (b), Navelli (c), and Assisi  
311 (d). The ridge lines are reported as black solid lines. The legend for color scale is the same as Figure 6.

312  
313

## 314 5. Conclusions

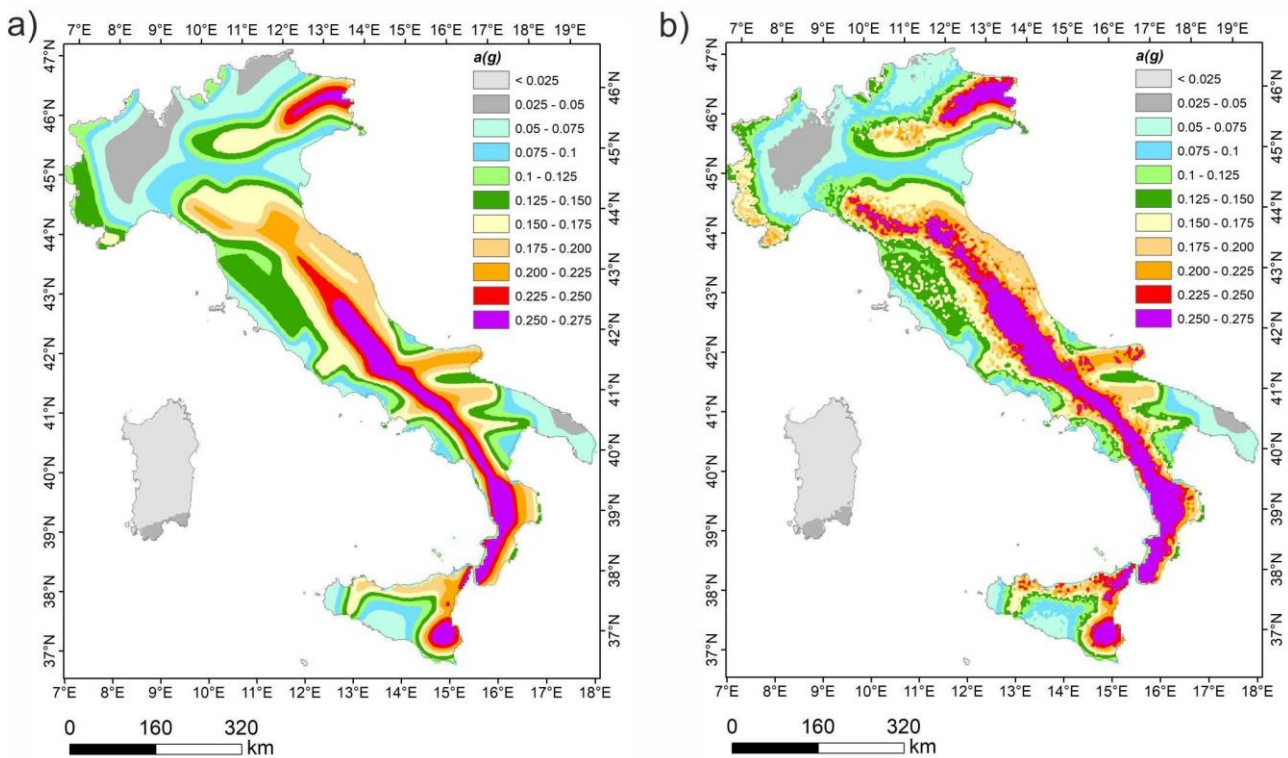
315

316 The paper describes a GIS procedure for the topographic classification of a DEM, according to the NTC18  
317 seismic code. The resulting map, with the topographic classification of Italy, is provided in the electronic  
318 supplement.

319 The GIS procedure is easily replicable: it makes use of (i) TinItaly high-resolution DEMs, (ii) the ridges  
320 identification with the Topographic Position Index (TPI; [35]), and (iii) slope angle and elevation differences  
321 calculated with built-in ESRI ArcGIS functions. The procedure allows to univocally classify the Italian  
322 territory and improves previous studies [24], where the classification was only performed in the neighbourhood  
323 of a single accelerometric station. Moreover, the results of this study do not need a manual revision, usually  
324 performed on Google Earth [48, 49, 38]. In this case, not applying quantitative and systematic morphological  
325 analyses based on digital elevation models, the topographic classification is affected by a certain level of  
326 subjectivity, related to the point of view, and to the complex geomorphology of the site. The overall check  
327 between the topographic map and the case studies reported in this work indicates a satisfactory representation  
328 of the morphological features and a correct classification of the sites. In case of mismatch, the proposed  
329 classification is generally conservative.

330 The statistics computed on the topographic classification map show that one third of the Italian territory is  
331 characterized by sites with possible topographic amplification (i.e., T2, T3 and T4 classes, Table 1). In  
332 particular, following the NTC18 and EC8 seismic codes, 25% of Italian territory is characterized by a  
333 topographic amplification factor of 1.2 (T2 and T3 classes), whereas 12% of the country area presents a  
334 topographic amplification factor of 1.4 (T4 class). The NTC18 prescribes the applicability of the topographic  
335 coefficients in case of “simple” configurations, but the criteria to discriminate between “simple” and  
336 “complex” are not specified and, for this reason, the applicability of the coefficients could be debated. Here  
337 we propose the application to the entire Italian territory, extending the classification prescribed in seismic  
338 codes for “simple” 2D relief configurations to all 3D geomorphological settings. The statistics computed on  
339 the 15332 Italian localities in DBMI [32] support the effectiveness of this extrapolation from 2D to 3D, but do  
340 not prove the existence of topographic amplification effects. Based on the topographic classification proposed  
341 in this study, about 40% of the localities that experienced the highest macroseismic intensities ( $I_{max} > 6$ ) are  
342 potentially affected by topographic amplification (T2-T4 classes). The impact of the topographic amplification  
343 (Table 1) on the hazard map [30], calculated as PGA for a return period of 475 years and soil category A  
344 (NTC18), is shown in Figure 10. Being the topographic amplification period-independent according to NTC18  
345 (Table 1), the effects are the same at all spectral periods.

346 The map of the topographic classification of Italy can be used for both scientific and application purposes. In  
347 this regard, the topographic classification map may be used for the characterization of seismic stations, and  
348 may support the identification of sites with systematic topographic effects. The map is also useful for the  
349 selection of accelerometric time histories that meet the spectrum-compatibility requirement indicated in the  
350 Italian (NTC18) and European (EC8) seismic codes. In the latter case, it allows selecting those stations in T1  
351 class, whose recordings are not affected by topographic amplification. Moreover, the proposed map may guide  
352 future investigations of T2, T3 and T4 sites, in particular to study the period-dependent topographic  
353 amplifications, as opposed to the constant amplification factors provided in NTC18 and EC8 seismic codes.  
354



355  
356 **Figure 10:** a) probabilistic seismic hazard map for PGA, soil class A and topographic class T1 (return period  
357 of 475 years), as in Ref. [30]; b) application of topographic amplification factors according to the classification  
358 shown in Figure 6.

359  
360  
361  
362



363 **Acknowledgment**

364 The study has partially benefited from funding provided by the Italian Presidenza del Consiglio dei Ministri–  
365 Dipartimento della Protezione Civile (DPC)—Agreement B2, DPC-INGV. This paper does not necessarily  
366 represent DPC official opinions and policies.

367

368 **CRedit authorship contribution statement**

369 **Claudia Mascandola:** Conceptualization, Methodology, Software, Writing - original draft, Writing - review  
370 & editing, Data curation, Formal analysis. **Lucia Luzi:** Conceptualization, Methodology, Software, Writing -  
371 original draft, Writing - review & editing, Data curation, Formal analysis. **Chiara Felicetta:** Writing - original  
372 draft, Writing - review & editing, Data curation, Formal analysis. **Francesca Pacor:** Writing - original draft,  
373 Writing - review & editing, Data curation, Formal analysis.

374

375 **References**

376 [1] Boore, D. M. (1972). A note on the effect of simple topography on seismic SH waves, *Bull. Seismol. Soc.*  
377 *Am.* 62, 275–284.

378 [2] Bard, P. Y. (1982). Diffracted waves and displacement fields over two dimensional elevated topographies,  
379 *Geophys. J. Int.* 71, 731–760.

380 [3] Panizza, M. (1991). Geomorphology and seismic risk. *Earth-Sci Rev*; 31:11–20.

381 [4] Paolucci, R. (2002). Amplification of earthquake ground motion by steep topographic irregularities. *Earthq.*  
382 *Eng. Struct. Dyn.*; 31:1831–53.

383 [5] Griffiths, D.W., and Bollinger, G.A. (1979). The effect of Appalachian Mountain topography on seismic  
384 waves. *Bull. Seism. Soc. Am.*; 69:1081–105.

385 [6] Kawase, H, and Aki, K. (1990). Topography effect at the critical SV-wave incidence: possible explanation  
386 of damage pattern by the Whittier Narrows, California, earthquake of 1 October 1987. *Bull. Seism. Soc. Am.*;  
387 80:1–22.

- 388 [7] Marsan, P, Milana, G, Pugliese, A, and Sanò, T. (2000). Local amplification effects recorded by a local  
389 strong motion network during the 1997 Umbria-Marche earthquake. In: Proceedings of 12th world conference  
390 on earthquake engineering, paper no. 1046. Auckland, New Zealand.
- 391 [8] Marra, F., Azzara, R., Bellucci, F., Caserta, A., Cultrera, G., Mele, B., Palombo, B., Rovelli, A., and  
392 Boschi, E. (2000). Large amplification of ground motion at rock sites within a fault zone in Nocera Umbra  
393 (central Italy). *J Seismol.*;4:543–54.
- 394 [9] Cara, F., Rovelli, A., Di Giulio, G., Marra, F., Braun, T., Cultrera, G., Azzara R., and Boschi E.(2005).  
395 The role of site effects on the intensity anomaly of San Giuliano di Puglia inferred from aftershocks of the  
396 Molise, Central Southern Italy, sequence, *Bull. Seism. Soc. Am.*;95:1457–68.
- 397 [10] Massa, M., Lovati, S., D’Alema, E., Ferretti, G., and Bakavoli, M. (2010). An experimental approach for  
398 estimating seismic amplification effects at the top of a ridge, and the implication for ground-motion  
399 predictions: the case of Narni, Central Italy. *Bull. Seism. Soc. Am.*; 100(6): 286–301.
- 400 [11] Buech, F., Davies, T.R., and Pettinga, J.R. (2010). The little red hill seismic experimental study:  
401 topographic effects on ground motion at a bedrock-dominated mountain Edifice. *Bull. Seism. Soc.*  
402 *Am.*;100:2219–29.
- 403 [12] Marzorati, S., Ladina, C., Falcucci, E., Gori, S., Saroli, M., Ameri, G., and Galadini F. (2011). Site effects  
404 on the rock: the case of Castelvechio Subequo (L’Aquila,central Italy). *Bull. Earthq. Eng.*;9:841–68.
- 405 [13] Paolucci, R., Faccioli, E., and Maggio, F. (1999). 3D Response analysis of an instrumented hill at  
406 Matsuzaki, Japan, by a spectral method. *J. Seismol.*;3:191–209.
- 407 [14] Gazetas, G., Kallou, P.V., and Psarropoulos, P.N. (2002). Topography and soil effects in the Ms 5.9  
408 Parnitha (Athens) earthquake: the case of Adámes. *Nat. Hazards*; 27:133–69.

- 409 [15] Assimaki, D., Gazetas, G., and Kausel, E.. (2005). Effects of local soil conditions and topographic  
410 aggravation seismic motion: parametric investigation and recorded field evidence from the 1999 Athens  
411 earthquake. *Bull. Seism. Soc. Am.*; 95:1059–89.
- 412 [16] Scandella, L, Davì, M, Paolucci, R, and Pessina, V. (2008). Numerical simulation and Gis mapping of  
413 seismic amplification due to topographic effects. In: *Proceedings of 31st European Seismological Commission*  
414 (ESC). Crete, Greece.
- 415 [17] Faccioli, E. (Eds.) (1986). *Elementi Per Una Guida Alle Indagini Di Microzonazione sismica – Progetto*  
416 *finalizzato “Geodinamica”, Quaderni de “La ricerca scientifica”, n.114, CNR, 7, 72–82, (in Italian).*
- 417 [18] Faccioli, E., Vanini, M., Frassinè, L. (2002). “Complex” site effects in earthquake ground motion,  
418 including topography. In: *Proceedings of 12th European Conference on earthquake engineering*. Paperno. 844.
- 419 [19] Galli, P, and Molin, D. (2004). *Macroseismic Survey of the 2002 Molise, Italy, Earthquake and Historical*  
420 *Seismicity of San Giuliano di Puglia. Earthq. Spectra*; 20(S1):39–52.
- 421 [20] Tertulliani, A, Arcoraci, L, Berardi, M, Bernardini, F, Camassi, R, Castellano, C, Del Mese, S., Ercolani,  
422 E., Graziani, L., Leschiutta, I., Rossi A., and Vecchi M. (2011). An application of EMS98 in a medium-sized  
423 city: the case of L’Aquila (Central Italy) after the April 6, 2009 Mw 6.3 earthquake. *Bull. Earthq. Eng*; 9:67–  
424 80.
- 425 [21] Gizzi, F.T., Potenza, M.R., and Zotta, C. (2012). 3 November 1980 Irpinia-Basilicata earthquake  
426 (Southern Italy): towards a full knowledge of the seismic effects. *Bull. Earth. Eng*; 10:1109–31.
- 427 [22] Cauzzi, C., Fäh, D., Pessina, V., Faccioli, E., and Smerzini, C. (2012). Topographic amplification from  
428 recorded earthquake data and numerical simulations. In *15th World Conference on Earthquake Engineering*.
- 429 [23] Gallipoli, M. R., Bianca, M., Mucciarelli, M., Parolai, S., and Picozzi, M. (2013). Topographic versus  
430 stratigraphic amplification: mismatch between code provisions and observations during the L’Aquila (Italy,  
431 2009) sequence. *Bull. of Earth. Eng.*, 11(5), 1325-1336.

- 432 [24] Pessina, V., and Fiorini, E. (2014). A GIS procedure for fast topographic characterization of seismic  
433 recording stations. *Soil Dynamics and Earthquake Engineering*, 63, 248-258.
- 434 [25] NTC (2018). Ministero delle Infrastrutture e dei Trasporti. Aggiornamento delle Norme Tecniche per le  
435 Costruzioni. Part 3.2.2: Categorie di sottosuolo e condizioni topografiche, *Gazzetta Ufficiale* n. 42 del 20  
436 febbraio 2018 (in Italian).
- 437 [26] CEN (Comite Europe en de Normalisation), 2004. Eurocode 8. Design Provisions for Earthquake  
438 Resistance of Structures—Part 5: Foundations, Retaining Structures and Geotechnical Aspects. ENV 1998-5,  
439 CEN European Committee for Standardisation, Brussels, 1994.
- 440 [27] NTC (2008). Ministero delle Infrastrutture e dei Trasporti. Nuove Norme Tecniche per le Costruzioni.  
441 Part 3: Categorie di sottosuolo e condizioni topografiche, *Gazzetta Ufficiale*, no. 29, del 4 febbraio 2008 (in  
442 Italian).
- 443 [28] Luzi, L., Hailemikael, S., Bindi, D., Pacor, F., Mele, F., and Sabetta, F. (2008). ITACA (ITalian  
444 ACcelerometric Archive): A Web Portal for the Dissemination of Italian Strong-motion Data, *Seismological*  
445 *Research Letters*, 79(5), 716–722.
- 446 [29] Pacor, F., Paolucci, R., Luzi, L., Sabetta, F., Spinelli, A., Gorini, A., Nicoletti, M., Marcucci, S., Filippi,  
447 L., and Dolce M. (2011), Overview of the Italian strong motion database ITACA 1.0, *Bull. Earth. Eng.*, 9(6),  
448 1723–1739.
- 449 [30] MPS Working Group (2004). Redazione della mappa di pericolosità sismica prevista dall’Ordinanza  
450 PCM del 20 marzo 2003, *Rapporto Conclusivo per il Dipartimento della Protezione Civile* (in Italian).
- 451 [31] Stucchi, M., C. Meletti, V. Montaldo, H. Crowley, G. M. Calvi, and E. Boschi (2011). Seismic hazard  
452 assessment (2003–2009) for the Italian building code, *Bull. Seismol. Soc. Am.* 101, no. 4, 1885–1911.
- 453 [32] Locati, M., Camassi, R., Rovida, A., Ercolani, E., Bernardini, F., Castelli, V., Caracciolo, C.H.,  
454 Tertulliani, A., Rossi, A., Azzaro, R., D’Amico, S., Conte, S., Rocchetti, E., Antonucci, A. (2019). Database

455 Macrosismico Italiano (DBMI15), versione 2.0. Istituto Nazionale di Geofisica e Vulcanologia (INGV).  
456 <https://doi.org/10.13127/DBMI/DBMI15.2>

457 [33] D'Amico, M., Felicetta, C., Russo, E., Sgobba, S., Lanzano, G., Pacor, F., and Luzi, L. (2020). Italian  
458 Accelerometric Archive v 3.1 - Istituto Nazionale di Geofisica e Vulcanologia, Dipartimento della Protezione  
459 Civile Nazionale. doi: 10.13127/itaca.3.1

460 [34] Tarquini, S., Isola, I., Favalli, M., Mazzarini, F., Bisson, M., Pareschi, M.T., and Boschi, E. (2007).  
461 TINITALY/01: a new triangular irregular network of Italy. *Annals of Geophysics* 50,407–425.

462 [35] Weiss, A. (2001). Topographic position and landform analysis. Poster presentation, ESRI Users  
463 Conference, San Diego, CA.

464 [36] Burrough, P. A., and McDonell, R. A., (1998). *Principles of Geographical Information Systems* (Oxford  
465 University Press, New York), 190 pp.

466 [37] Jenness, J., Brost, B., and Beier, P. (2013). Land facet corridor designer. USDA forest service rocky  
467 mountain research station.

468 [38] Mascandola, C., Massa, M., Lovati, S., and Augliera, P. (2018). The site characterization scheme of the  
469 INGV strong motion database (ISMD): Overview and site classification. *Seism. Res. Lett.*, 89(1), 86-98.

470 [39] Zhan, C. (1993). A Hybrid Line Thinning Approach. *Proceedings Auto-Carto 11*, Minneapolis, pp. 396-  
471 405.

472 [40] OPCM n. 3519 (2006): criteri generali per l'individuazione delle zone sismiche e per la formazione e  
473 l'aggiornamento degli elenchi delle stesse zone; *Gazzetta Ufficiale* n. 108 dell'11 maggio 2006 (in Italian).

474 [41] ISTAT, 2011. 154° Censimento della Popolazione e delle Abitazioni. [http://www.istat.it/it/censimento-](http://www.istat.it/it/censimento-popolazione/censimento-popolazione-2011)  
475 [popolazione/censimento-popolazione-2011](http://www.istat.it/it/censimento-popolazione/censimento-popolazione-2011).

476 [42] Sieberg, A. (1930). The earthquake geology (*Geologie der Erdbeben*). *Handbuch der Geophysik*  
477 2(4):550–555 (in German).

- 478 [43] Caserta, A., Bellucci, F., Cultrera, G., Donati, S., Marra, F., Mele, G., Palombo, B. and Rovelli, A. (2000).  
479 Study of site effects in the area of Nocera Umbra (Central Italy) during the 1997 Umbria-Marche seismic  
480 sequence, *J. Seismol.*, 4, 555–565.
- 481 [44] Donati, S., Marra, F. and Rovelli, A. (2001). Damage and ground shaking in the town of Nocera Umbra  
482 during Umbria-Marche, central Italy, earthquakes: the special effect of a fault zone, *Bull. seism. Soc. Am.*, 91,  
483 511–519.
- 484 [45] Pischietta, M., Cultrera, G., Caserta, A., Luzi, L., and Rovelli, A. (2010). Topographic effects on the hill  
485 of Nocera Umbra, central Italy. *Geophysical Journal International*, 182(2), 977-987.
- 486 [46] Lovati, S., Bakavoli, M. K. H., Massa, M., Ferretti, G., Pacor, F., Paolucci, R., Haghshenas E., and  
487 Kamalian, M. (2011). Estimation of topographical effects at Narni ridge (Central Italy): comparisons between  
488 experimental results and numerical modelling. *Bull. of Earth. Eng.*, 9(6), 1987-2005.
- 489 [47] Pischietta, M., Cianfarra, P., Salvini, F., Cara, F., and Vannoli, P. (2018). A systematic analysis of  
490 directional site effects at stations of the Italian seismic network to test the role of local topography. *Geophysical*  
491 *Journal International*, 214(1), 635-650.
- 492 [48] Pessina, V. (2010). Appendix E. EC8 subsoil and topographic classification of ITACA stations  
493 Deliverable D10, Project S4 The Italian strong motion database.  
494 [http://esse4.mi.ingv.it/files/images/stories/d10\\_appendix%20e.pdf](http://esse4.mi.ingv.it/files/images/stories/d10_appendix%20e.pdf)
- 495 [49] Di Capua, G., Lanzo, G., Pessina, V., Peppoloni, S., and Scasserra, G. (2011). The recording stations of  
496 the Italian strong motion network: Geological information and site classification. *Bulletin of Earthquake*  
497 *Engineering*, 9(6), 1779-1796.



**Declaration of interests**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

**CRedit authorship contribution statement**

**Claudia Mascandola:** Conceptualization, Methodology, Software, Writing - original draft, Writing - review & editing, Data curation, Formal analysis. **Lucia Luzi:** Conceptualization, Methodology, Software, Writing - original draft, Writing - review & editing, Data curation, Formal analysis. **Chiara Felicetta:** Writing - original draft, Writing - review & editing, Data curation, Formal analysis. **Francesca Pacor:** Writing - original draft, Writing - review & editing, Data curation, Formal analysis.