ISTITUTO NAZIONALE DI GEOFISICA E VULCANOLOGIA

Site characterization report at the seismic station IT.BRSA – Brescia Città

Report di caratterizzazione di sito presso la stazione sismica IT.BRSA – Brescia Città

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Subject: **Final report illustrating the site characterization for seismic station IT.BRSA**	

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INTRODUCTION

In this report we present the geological setting and the geophysical measurements and results obtained in the framework of the 2019-2021 agreement between INGV and DPC, called *Allegato B2: Obiettivo 1 - TASK 2: Caratterizzazione siti accelerometrici (Responsabili: G. Cultrera, F. Pacor)* for the site characterization of station IT.BRSA (Brescia Città).

Location and coordinates are reported in Table 1.

Table 1

CODE	NAME	LAT [°]	LON [°]	ELEVATION [m]
IT.BRSA	Brescia Città	45.55270 *	10.22550 *	157 **
ADDRESS	Via Guglielmo Marconi, 25, 25128 Brescia (BS), Italy			

* Coordinates from ITACA (Nov. 2021)   ** Elevation from CTR 10k Regione Lombardia

A. Geological setting

A1. TOPOGRAPHIC AND GEOLOGICAL INFORMATION

Topographic information related to the site are reported in Table 2. Table 3 summarizes all available geological maps from literature for geological analyses.

Table 2

Topography	Description	Topography Class	Morphology Class
	Flat surfaces, isolated slope and reliefs with slope i ≤ 15°	T1	Valley edge (VE)

Table 3

Geological map	Source	Scale
IT.BRSA	Geological Map of Italy, sheet 47 (Brescia)	1:100.000

In Table 4 Geological and Lithotechnical Units (according to Seismic Microzonation classification; Technical Commission SM, 2015) are described and are concerned to maps of following chapters. The term “original” means the result comes from a preexisting cartography (Table 3); the term “deduced” means the result comes from an interpretation of a preexisting cartography according to the nomenclature of corresponding cartography.

Table 4

GEOLOGICAL UNITS		**LITHOLOGICAL UNITS**		**LITHOTECHNICAL UNITS**	
Geological Map of Italy 1:100.000, sheet 47 (Brescia) *original*		Amanti *et al.* (2008) *deduced*		(MZS) *deduced*	
code	**description**	**code**	**description**	**code**	**description**
f_g^w	Fluvial deposits	B4	Mixed grain size deposits	GM pd	Gravel-sand-silt mixture
dt^1	Debris, alluvial fans	B7	Undefined soils	GC cz	Gravel-sand-clay mixture
G_2^3	Monte Domaro Limestone	A3	Marly limestone	SFALS	Fractured /weathered layered alternance of lithotypes

A2. GEOLOGICAL MAP

In Figure 1 Geological Map is reported in a 1 $km \times 1$ km square around the station.

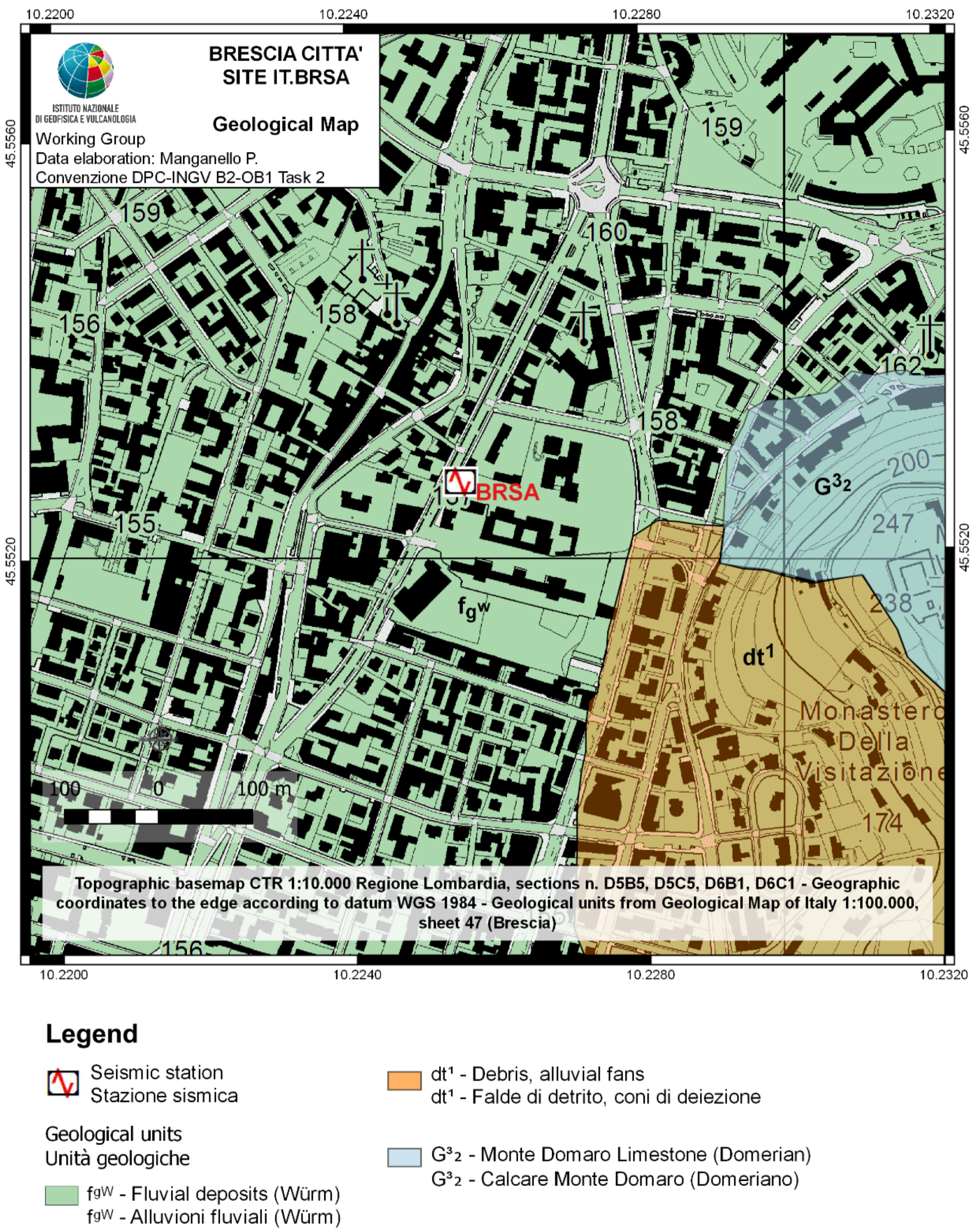

Figure 1: Geological map of seismic station IT.BRSA. Scale 1:5.000. Geological units come from Geological Map of Italy 1:100.000, sheet 47 (Brescia).

A3. LITHOLOGICAL MAP

In Figure 2 Lithological Map is reported in a 1 $km \times 1$ km square around the station.

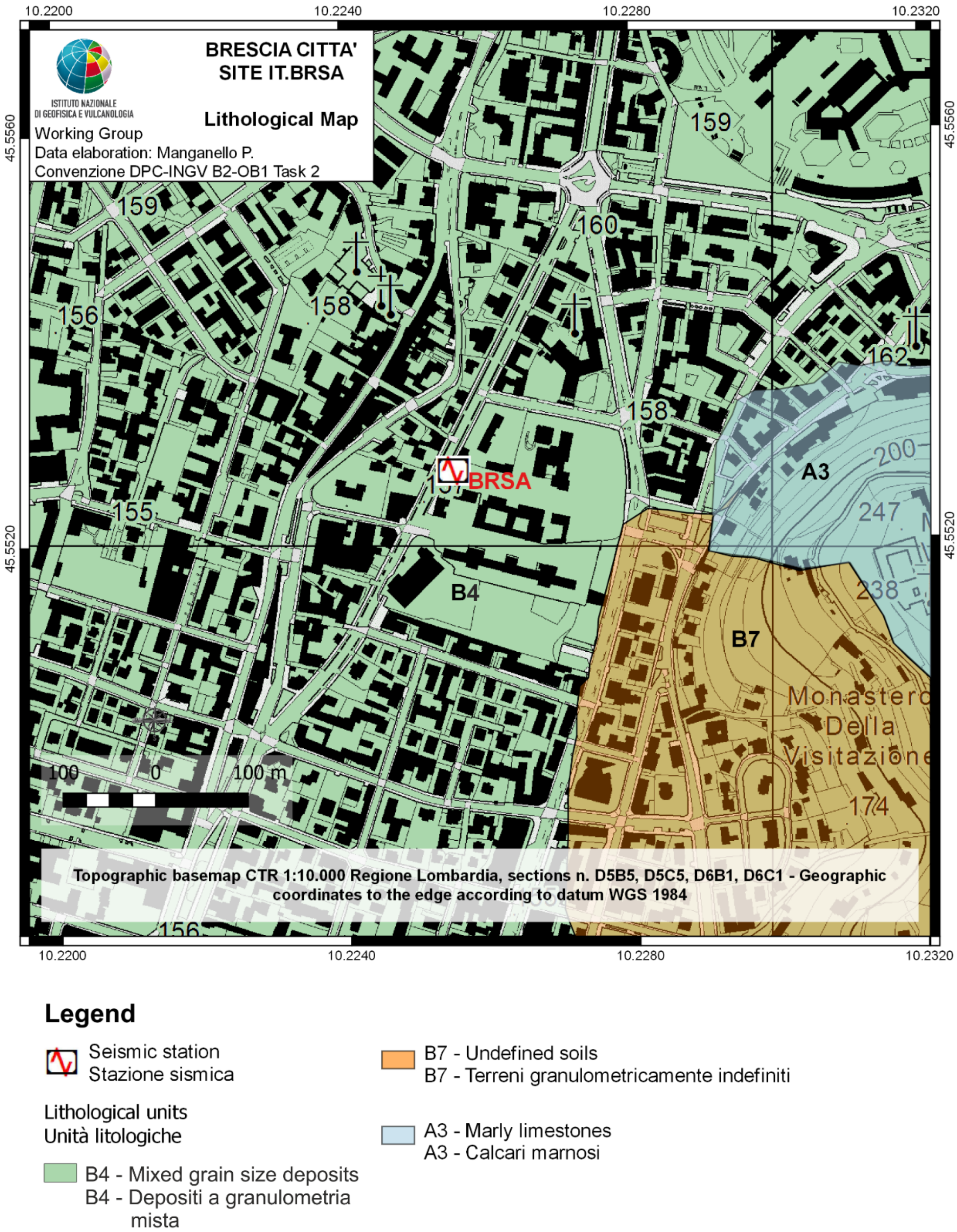

Figure 2: Lithological map of the seismic station IT.BRSA. Scale 1:5.000. The codes of the lithological units are assigned according to the nomenclature of the Lithological map of Italy ISPRA 1:100.000 (Amanti *et al.*, 2008).

A3. LITHOTECHNICAL MAP

In Figure 3 Lithotechnical Map is reported in a 1 $km \times 1$ km square around the station.

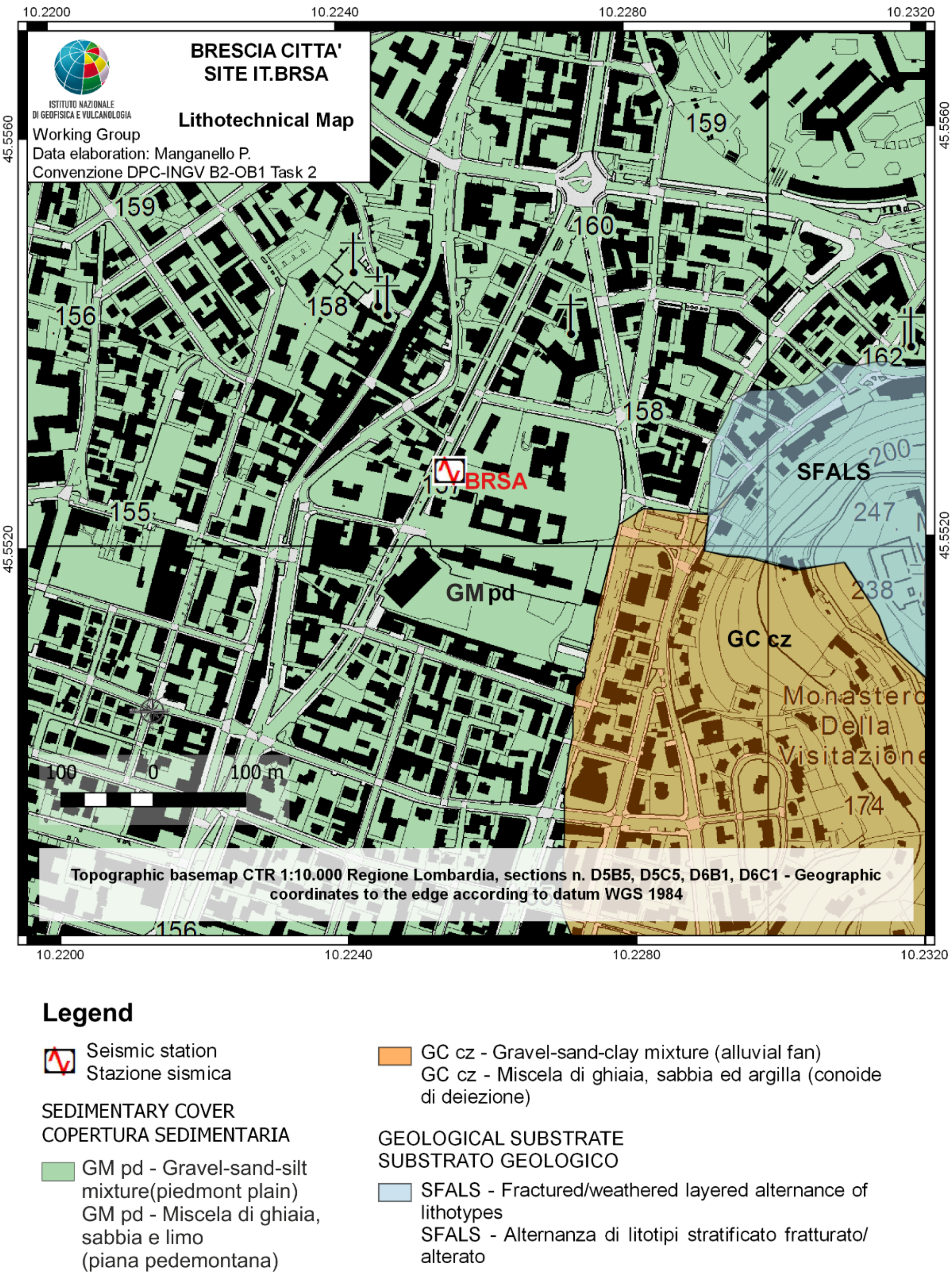

Figure 3: Lithotechnical map of the seismic station IT.BRSA. Scale 1:5.000. The lithotechnical units are deduced according to the nomenclature of Seismic Microzonation (Technical Commission SM, 2015).

A5. SURVEY MAP

Figure 4 shows the Survey Map reporting both previous investigations and geophysical surveys conducted by INGV Working Group.

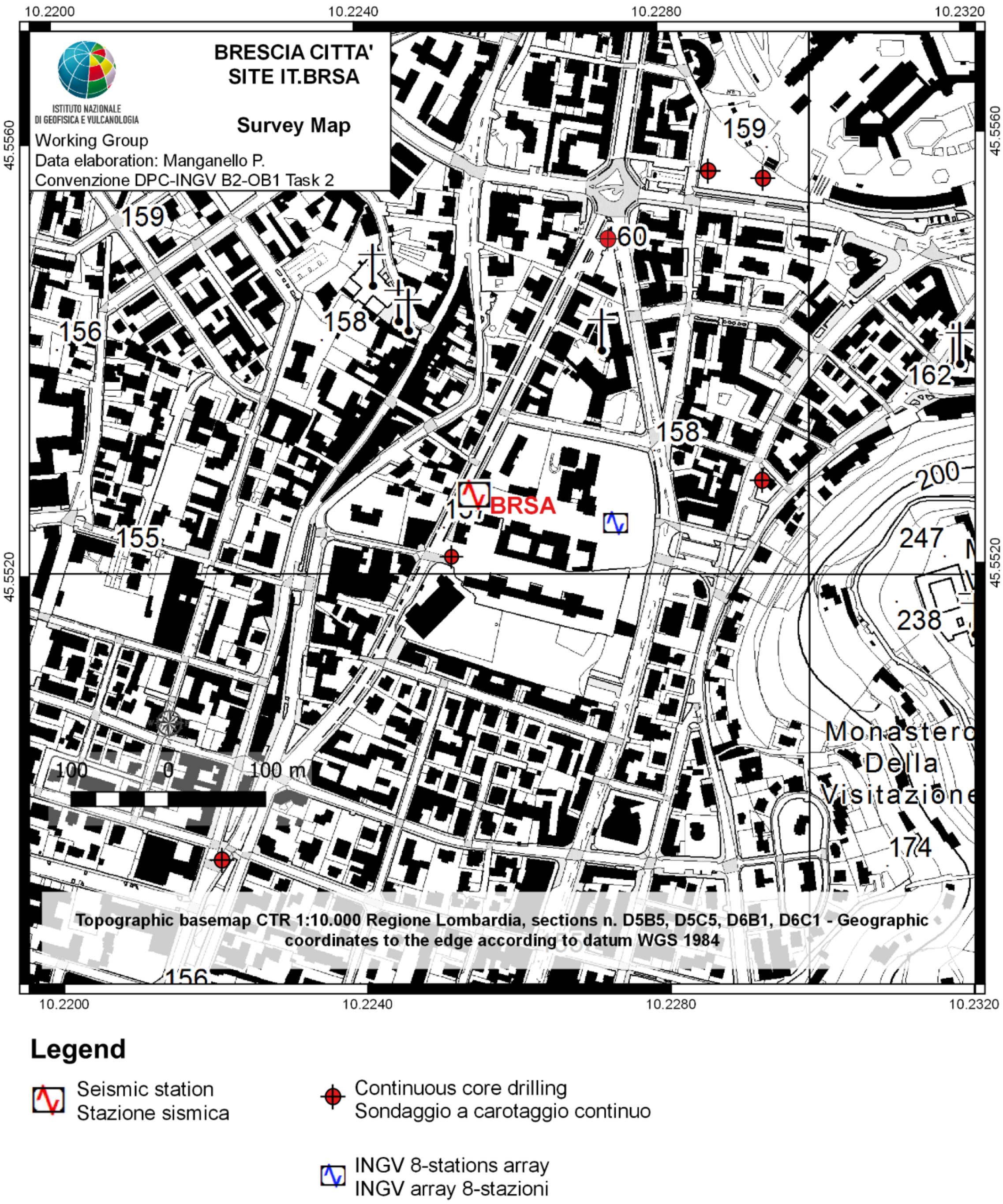

Figure 4: Map of the surveys in the surroundings of the station IT.BRSA. Scale 1:5.000.

A6. GEOLOGICAL MODEL

6.1 General description

The IT.BRSA seismic station is located in the town centre of Brescia municipality, which represents the capital of the homonymous province. Brescia municipality is located in the northern Po Plain close to the Trompia Valley. The territory of Brescia municipality is bounded to the North by the Brescia Prealps and to the East by the Garda Prealps.

The Mella River flows in NNE-SSW direction in the western part of Brescia municipality. This river has eroded since Upper Pleistocene the fluvio-glacial deposits of the plain. After that the deposition of fluvial and fluvio-glacial deposits of Würm, terraced alluvial deposits (Holocene) and recent alluvial deposits has occurred (Comune di Brescia, 2017).

The geological setting of the studied area is connected to the evolution of the Lombardian Basin, which represents a structurally complex area of the Mesozoic South-Alpine rifted margin, located between the Lake Maggiore fault and the Ballino - Garda fault. After the Liassic extension the Lombardian Basin consisted of several half-grabens delimited by normal faults. In the Lower Toarcian the tectonic activity ended and the turbiditic deposition was replaced by thick pelagic sedimentation across the entire Lombardian basin. The Cenozoic Alpine collisional history is thus primarily responsible for the geological and structural setting of the area (Bertotti *et al.*, 1993; Cassinis *et al.*, 2000).

6.2 Geological section

In the surroundings of IT.BRSA seismic station the executed surveys are represented by 6 continuous core drillings (depth between 20 and 52 *m*).

The WSW-ENE oriented geological section is reported and highlights the geological and structural setting of the IT.BRSA site. The trace with the location of the section is reported as a black line in the geological map (Fig. 5 upper left).

6.3 Subsoil model

The geological description reported from the surface to the bottom is described in the following part. A subsoil model is built up to a depth of 35 *m* on the basis of geological and stratigraphic information (Figure 5 bottom).

The stratigraphic succession starts with the fluvial deposits of Würm (f_g^w), which are characterized by mixed grain size sediments. At a depth of about 20-25 *m* the stratigraphic succession shows the presence of the Monte Domaro Limestone (G_2^3), represented by marly limestones with marly intercalations (Domerian). The Monte Domaro Limestone belongs to the Medolo Group (Southern Alps sedimentary succession).

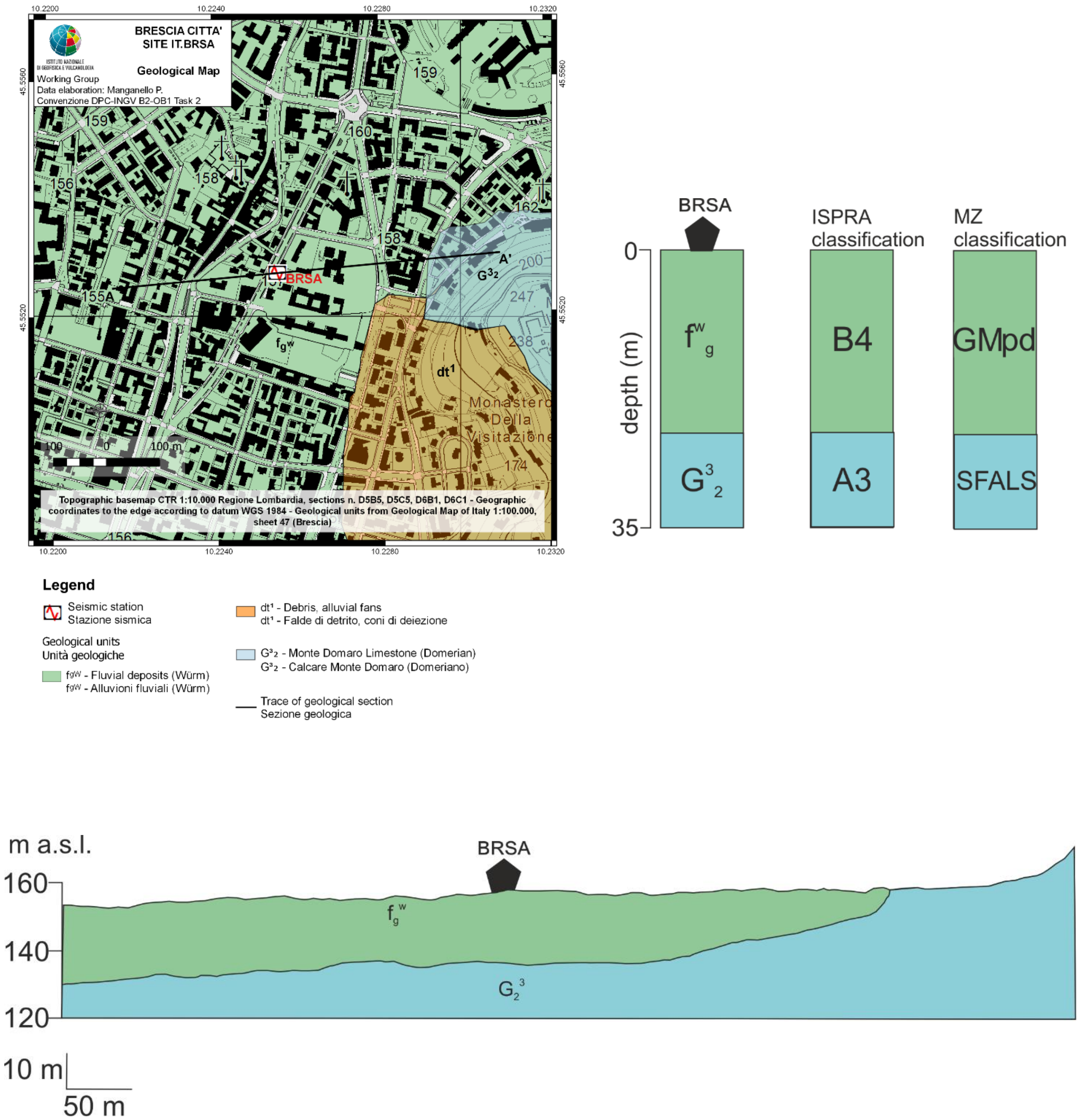

Figure 5: Upper left: Geological map of the study area where is installed IT.BRSA seismic station. Upper right: Geological section. Bottom: Subsoil model for the site.

B. V_S profile

B1. GEOPHYSICAL INVESTIGATIONS

Geophysical measurements executed nearby the station BRSA of the network IT (PCM-DPC, 1972) consist in ambient-vibration measurements in both single-station and 2D array configuration (Figure 6) that provide results in terms of resonance frequency of the soil deposits and in terms of dispersion curves of surface waves. These curves are inverted to obtain a shear-wave velocity (V_S) profile that, together with the geological study at section A, is suitable for assigning the soil class according to the current Italian seismic code (NTC18) and Eurocode (EC8). Figure 7 shows the location of the station IT.BRSA (Latitude 45.5527, Longitude 10.2255 WGS84) installed in Brescia (BS).

Seismic noise is acquired using 8 Reftek-130 24-bits recording systems equipped with short-period Lennartz LE-3D/5s sensors and GPS timing (Figure 7). The sampling rate is fixed to 200 *Hz*, while the gain is set as "high". Ambient noise recordings have a minimum duration of 1 hour. The array geometry (Figure 8) is chosen in order to have a good coverage of both azimuths and inter-station distances, the latter between the minimum (less than 10 *m*) and the maximum (about 30 *m*). These ranges allow the analysis of a range of wavelengths that guarantee sufficient shallow resolution (Okada, 2003) in order to estimate the $V_{S,30}$ and the site-class according to current building codes (i.e. NTC18 and EC8).

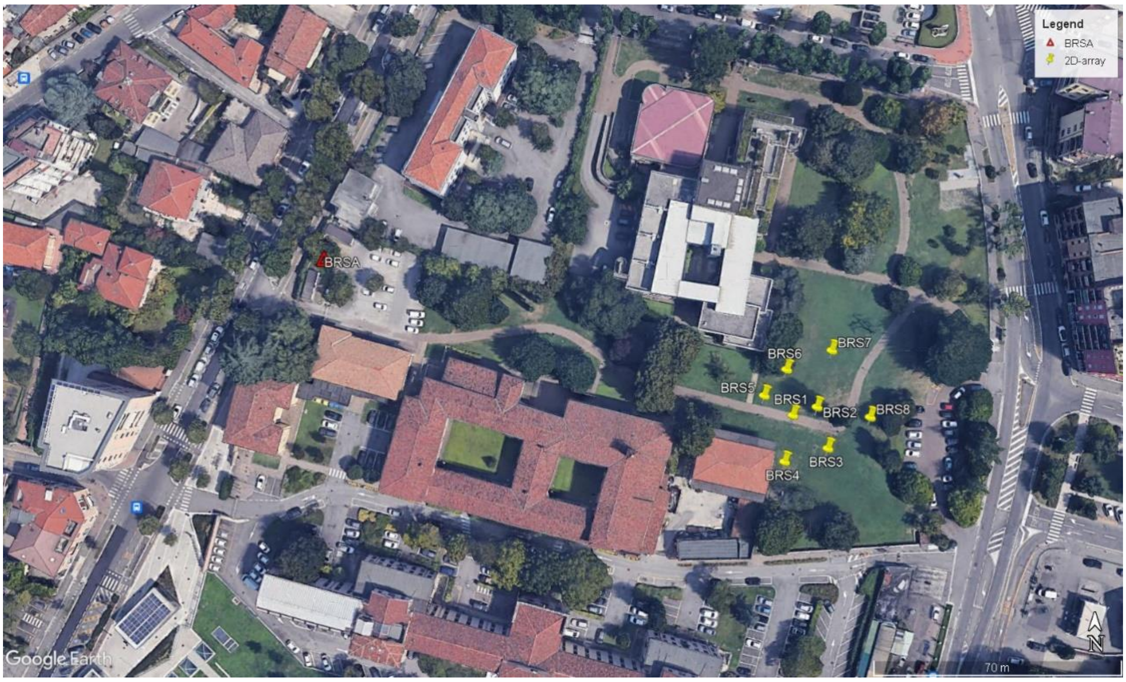

Figure 6: Map of the geophysical measurements performed at the IT.BRSA site. The yellow place-markers indicate the geometry used for 2D array in passive configuration. The red triangle indicates the IT.BRSA accelerometric station (image from Google Earth http://www.earth.google.com).

Figure 7: Left: Single station ambient noise measurement. Right: 2D passive ambient noise array installed close to the IT.BRSA accelerometric station.

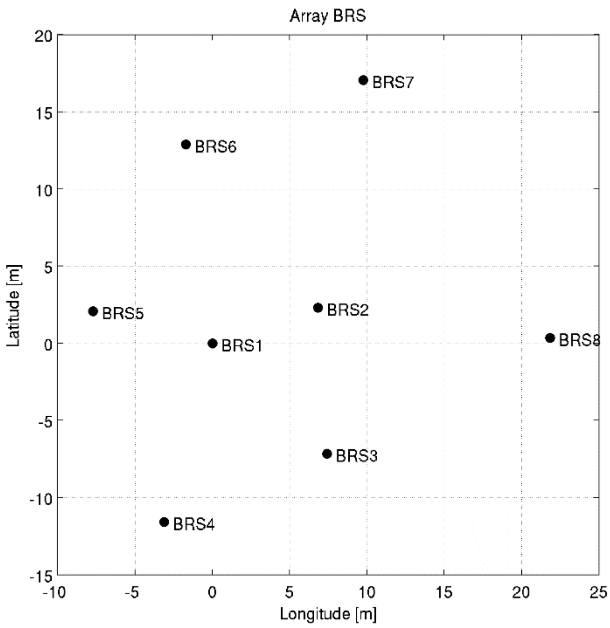

Figure 8: Array geometry.

The first step of the analysis consists in a visual inspection of the recordings at each station of the array. In particular, in order to identify malfunctioning and to select signal windows suitable for the surface wave analysis, the quality of the recording is evaluated analyzing the signal stationarity in the time domain, the relevant unfiltered Fourier spectra, and the H/V variation over time. Figures 9 and 10 provide graphical results about station BRS5.

It is common practice during surface wave investigation to verify the reliability of the one-dimensional site structure assumption (Aki, 1957; Okada, 2003). For this reason, we estimated the HVSR at each station of the array and the stability of HVSR among the array stations has been verified. Figure 11 depicts the HVSR assumed as representative for the array.

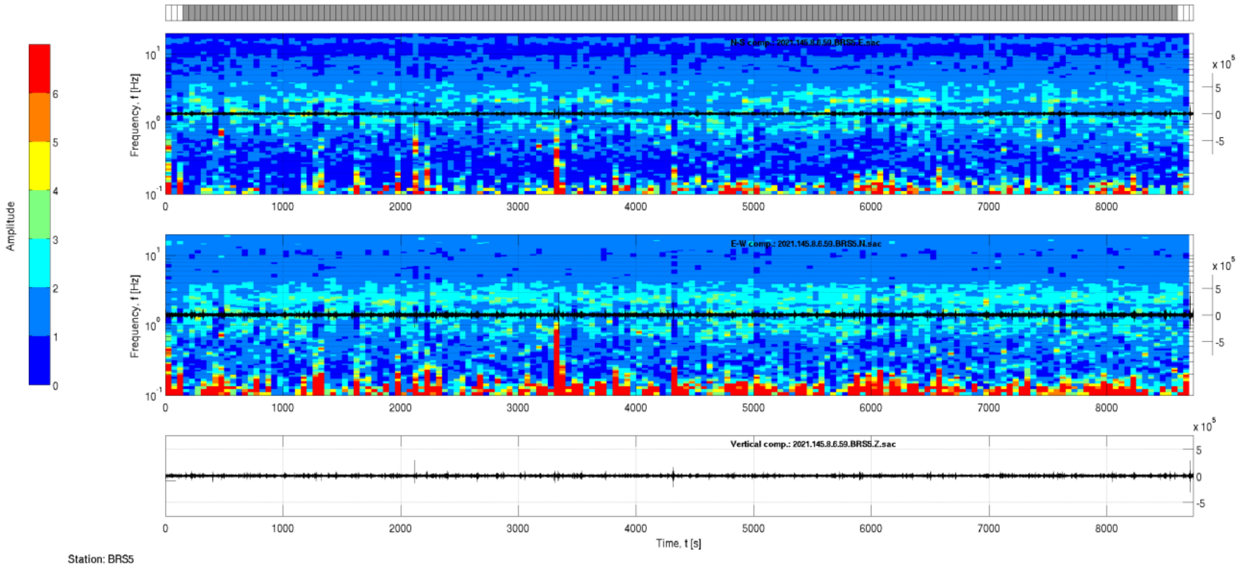

Figure 9: HVSR versus time (top and central panel for the NS and EW component, respectively) and corresponding time-histories.

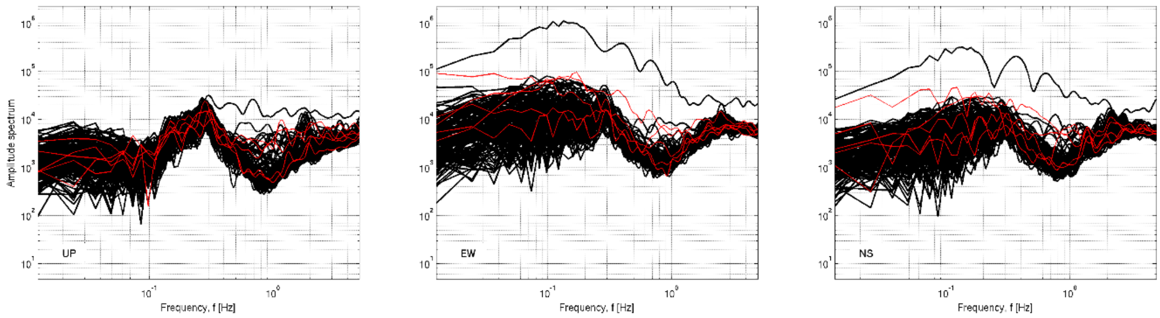

Figure 10: Fourier spectra for each noise window (left: Vertical, center: EW, right: NS). Red spectra are excluded from HVSR analysis.

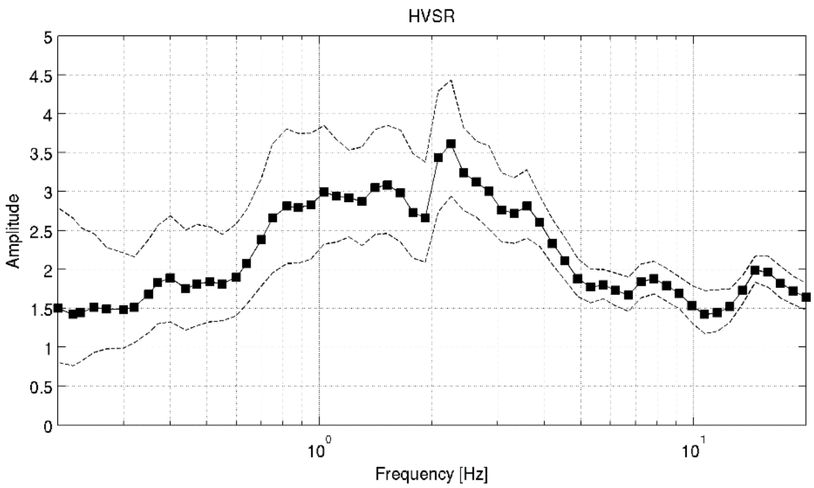

Figure 11: HVSR representative for the array. Dashed lines represent +/- one standard deviation.

The Rayleigh-wave dispersion curve is estimated by analyzing the vertical component of the recorded seismic noise. In particular, the Extended Spatial Auto-Correlation (ESAC; Ohori *et al.*, 2002; Okada, 2003) and the frequency-wavenumber (F-K; Lacoss *et al.*, 1969; Capon, 1969) methods are adopted. Further details about the combined use of ESAC and F-K approaches can be found in Parolai *et al.* (2006).

Both analyses use 50 synchronized signal windows of 60 *s* each, extracted from recordings within the UTC date-time interval 2021-05-25 09:10:00 – 2021-05-25T10:00:00, avoiding time periods affected by local disturbance.

The ESAC Rayleigh-wave dispersion curve is obtained by minimizing the root-mean-square (RMS) of the differences between experimental and theoretical Bessel functions (Figure 12). Values differing by more than two standard deviations from those estimated by the best fitting functions are automatically discarded (red circles in Figure 12) and the procedure is repeated iteratively. For this data set, data are also discarded whenever the inter-station distance is 2 times longer than the relevant wavelength. Figure 13 shows the Rayleigh-wave dispersion curve estimated using the ESAC approach.

The F-K analysis allows checking on the noise source distribution. One of the basic assumptions for the application of the ESAC method is indeed that the seismic noise wavefield is nearly isotropic. Figures 14 and 15 show results of the F-K analysis in terms of power density function for several frequencies using the Maximum Likelihood Method (MLM) and the Beam-Forming (BF) respectively. Figure 16 shows the good agreement above 9 *Hz* between the Rayleigh wave dispersion curves estimated by both ESAC and F-K approaches. As expected, due to the array geometry, below this threshold the F-K analysis provides larger phase velocities.

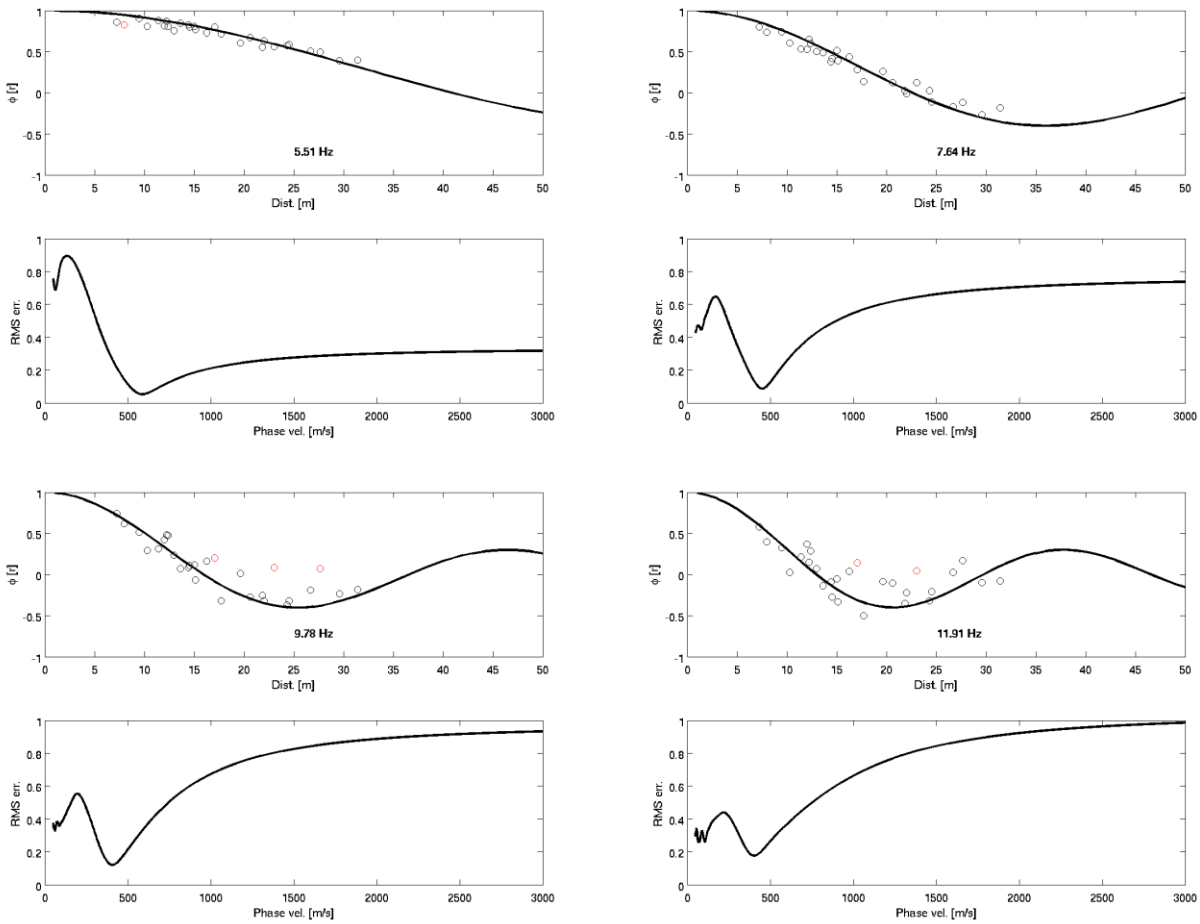

Figure 12: Experimental space-correlation function values versus distance (circles) for different frequencies. The red circles indicate values that are discarded. The black lines depict the estimated space-correlation function values for the phase velocity that furnishes the best fit to the data. The bottom panels show the relevant root-mean-square errors (RMS) versus phase velocity tested.

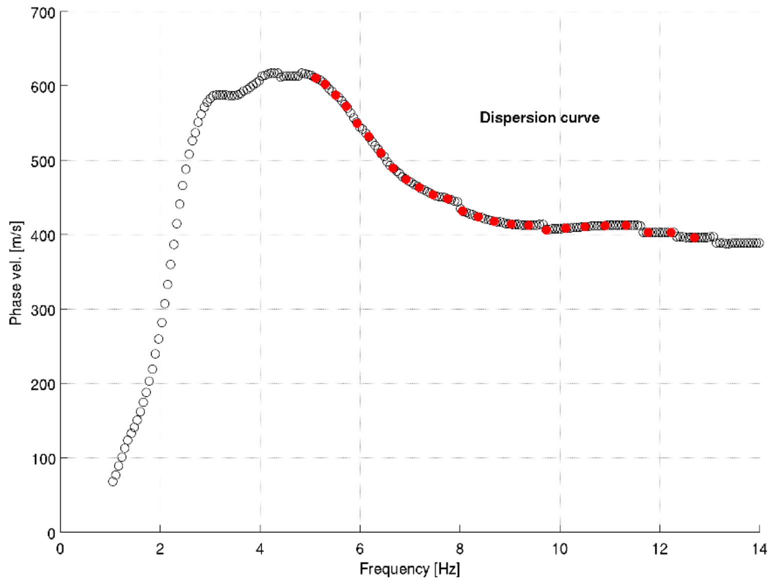

Figure 13: Rayleigh-wave dispersion curve from ESAC. Red-filled circles represent values potentially used for inversions.

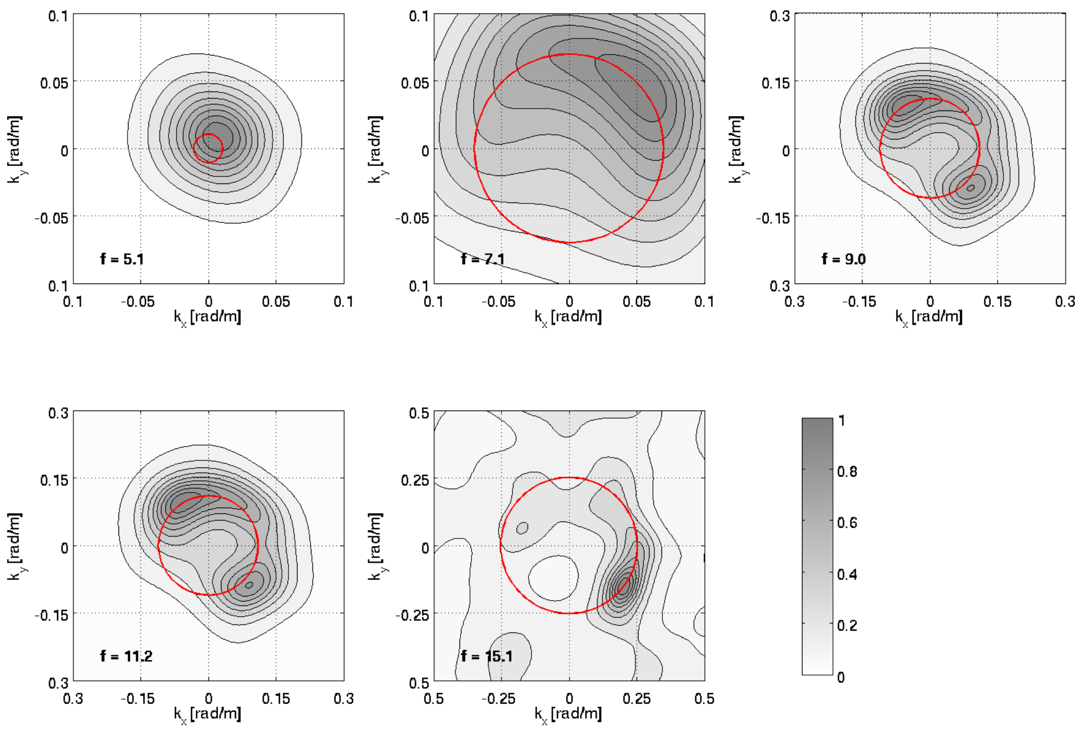

Figure 14: F-K power density function (Maximum-Likelihood Method) at selected frequencies.

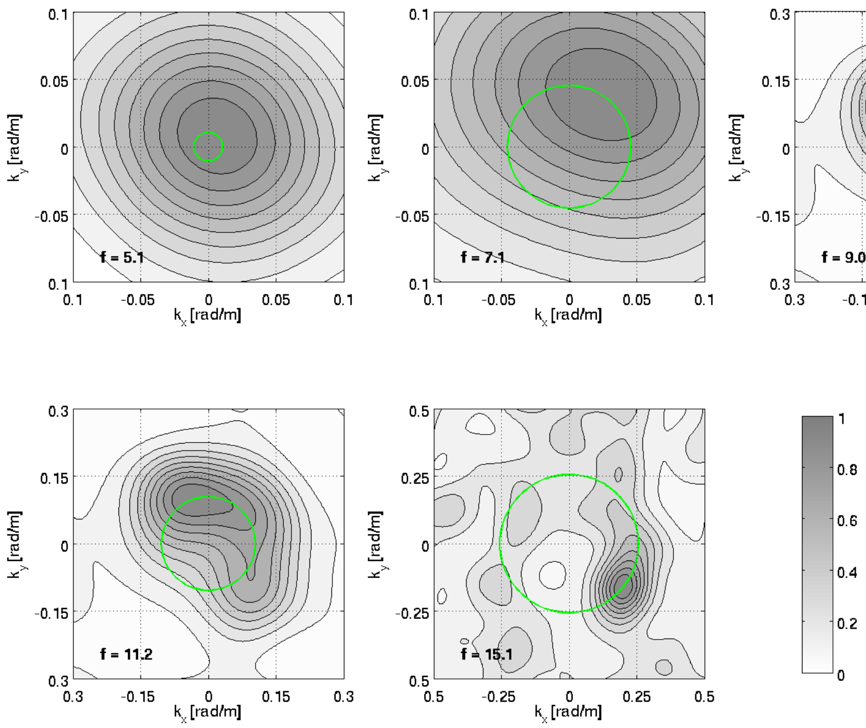

Figure 15: F-K power density function (Beam-Forming) at selected frequencies.

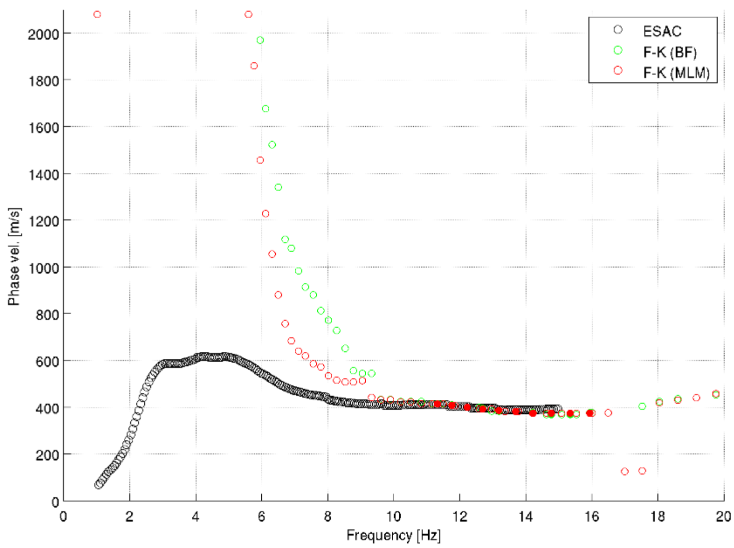

Figure 16: Comparison of experimental phase velocity estimated by the ESAC and the F-K (for both Beam-Forming and Maximum-Likelihood Method) methods. Filled circles represent values potentially used for inversions.

B2. SEISMIC VELOCITY MODEL

The non-linear inversions are performed using the software *joinv6* (Parolai *et al.*, 2005; Giustiniani *et al.*, 2020), which adopt a genetic algorithm (Yamanaka and Ishida, 1996). The forward modelling of Rayleigh wave phase velocities and HVSR curves is performed under the assumption of a vertically heterogeneous 1D Earth model using the modified Thomson-Haskell method proposed by Wang (1999) and following the suggestions of Arai and Tokimatsu (2004) and Tokimatsu *et al.* (1992). The modelling is not restricted to the fundamental mode, preserving the possibility that higher modes participate in simulating the observed dispersion and HVSR curves.

The experimental dispersion curve used as input for inversions is the one estimated from the ESAC analysis in the frequency interval 5-11 *Hz*, in combination with the one from F-K in the interval 11-16 *Hz*. The experimental HVSR is used between about 0.7 and 5 *Hz*. In the left panel of Figure 17 tested models are shown in different colors according to their cost value: the more reliable model (minimum cost) is in white, the models lying inside the 10% range of the minimum cost are in black and the other tested models are shown in grey. In the right-central and right-bottom panels of Figure 17 agreement between experimental and theoretical (grey and open circles, respectively) Rayleigh-wave dispersion curves and HVSR are shown. The agreement is good and, considering the wavelengths related to the dispersion curve frequency range, the V_S profile between about 5-60 *m* is very well constrained. Table 5 reports the minimum-cost shear-wave velocity model.

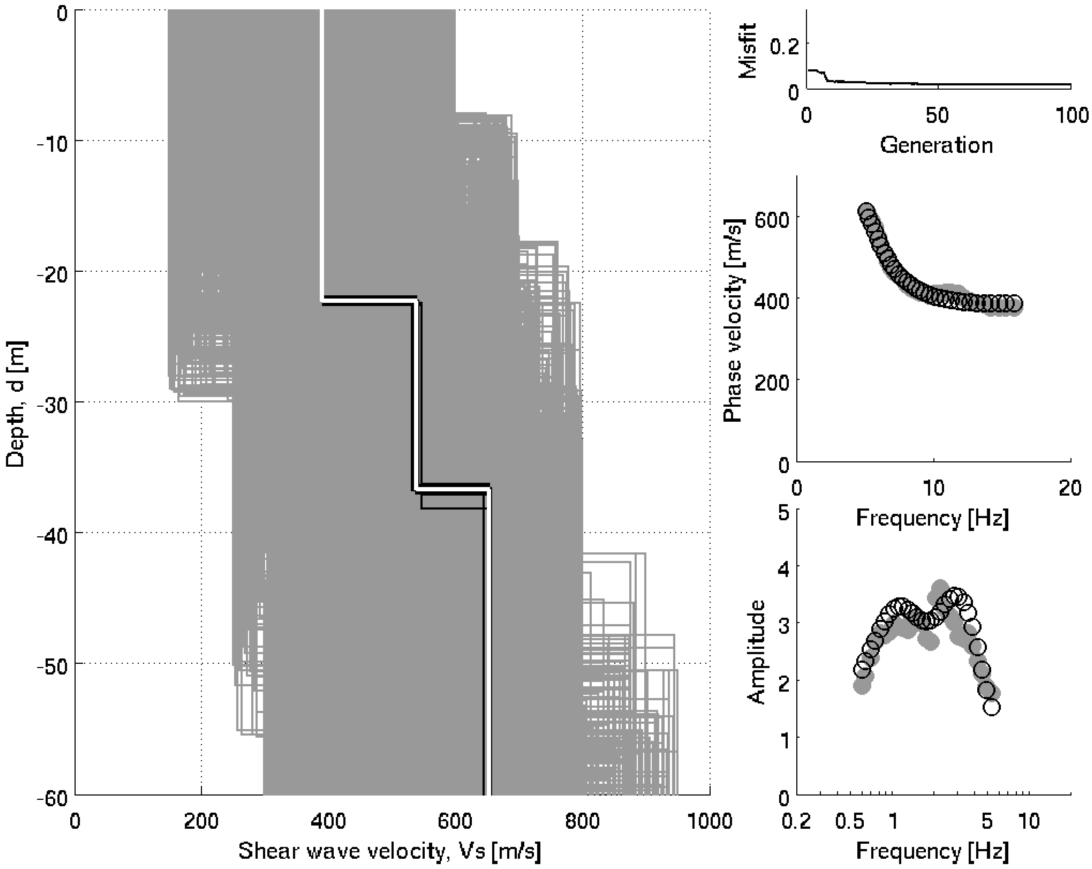

Figure 17: Shear-wave velocity models modeled during the inversion procedure (left panel): tested models (grey lines), the minimum cost model (white line) and models lying inside the minimum cost + 10% range (black lines). The generation values versus misfit (right-upper panel). The fitting of experimental data (grey circles) and empirical values relative to the minimum cost model (white circles) relevant to the dispersion curve (right-central panel) and to HVSR (right-bottom panel).

Table 5: Best-fit shear-wave velocity model

From [*m*]	To [*m*]	Thickness [*m*]	V_S [*m/s*]
0	22	22	390
22	37	15	539
37	-	-	653

B3. CONCLUSIONS

As evinced from results of geophysical investigations carried out by INGV Working Group, we can attribute to the fluvial deposits of Würm (f_g^w) a V_S value of 390 m/s and to the Monte Domaro Limestone a V_S value of 539 m/s, compatible with EC8 class assigned at the site according to geological evidences.

According to the current Italian seismic code (NTC18), if the bedrock (V_S > 800 m/s) is more than 30 m in depth, the equivalent velocity ($V_{S,eq}$) is equal to the $V_{S,30}$. From Figure 17, the velocity of 800 m/s is reached for an unknown depth, well below the depth of 30 m.

Therefore, in this case, both $V_{S,eq}$ and $V_{S,30}$ are equal to 421 m/s. Of consequence, IT.BRSA site is classified in the soil category B, for both the NTC18 and EC8 seismic codes (Table 6).

Table 6: $V_{S,eq}$, $V_{S,30}$ and soil classes

$V_{S,eq} = V_{S,30}$ [m/s]	Soil class (NTC18)	Soil class (EC8)
421	B	B

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[1]This document is level 3 as defined in the "Principi della politica dei dati dell'INGV (D.P. n. 200 del 26.04.2016)"