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The Global Stratotype Section and Point (GSSP) for the base of the Chattian Stage (Paleogene System, Oligocene Series) at Monte Cagnero, Italy

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*The GSSP for the base of the Chattian Stage (Paleogene System, Oligocene Series) is defined at meter level 197 in the Monte Cagnero section, which belongs to the pelagic succession of the Umbria–Marche basin (Urbania, central Italy: 43°38'47.81"N–12°28'03.83"E). This level with an astronomical age of 27.82 Ma coincides with the highest common occurrence of the planktonic foraminifer *Chiloguembelina cubensis* at the base of planktonic foraminiferal O5 Zone and falls in the upper part of calcareous nannofossil NP24 Zone, in the lower part of dinocyst Dbi Zone, and in the lower Chron C9n. The proposal was approved by the International Subcommission of Paleogene Stratigraphy in July 2015, approved by the International Commission of Stratigraphy in August 2016, and ratified by the International Union of Geological Sciences in September 2016.*

Introduction

The Rupelian and Chattian historical stratotypes were respectively located in northwestern Belgium and northwestern Germany, in the

southern part of the North Sea basin (e.g., Berggren, 1971; Hardenbol and Berggren, 1978; Van Simaey et al., 2004). Chronostratigraphically meaningful calcareous plankton is sparse in the generally marginal marine sedimentary environment of the historical reference sections, paleomagnetic signals are poor, and radiometrically datable volcanic horizons are absent. Therefore, the historical type localities do not represent suitable and continuous sections for the GSSP that comprises the Rupelian/Chattian boundary. Accordingly, a continuous, well-calibrated base Chattian GSSP section offering good correlation potential has to be sought outside the North Sea basin.

In search of a continuous section as GSSP of the Rupelian/Chattian boundary, the Oligocene Integrated Stratigraphy (OLIS) Working Group was formed to provide an integrated stratigraphy of the pelagic Scaglia Cinerea Formation in the Umbria–Marche (U–M) region of central Italy (Coccioni et al., 2008; Pross et al., 2010), which spans the uppermost Eocene through the entire Oligocene. Of the three sections identified and studied to verify regional correlations – the Contessa Barbetti Road (CBR) near Gubbio, and Pieve d'Accinelli (PAC) near Piobbico, and Monte Cagnero (MCA) near Urbania (Coccioni et al., 2008) – the latter was selected to define the basal Chattian Stage GSSP (Fig. 1).

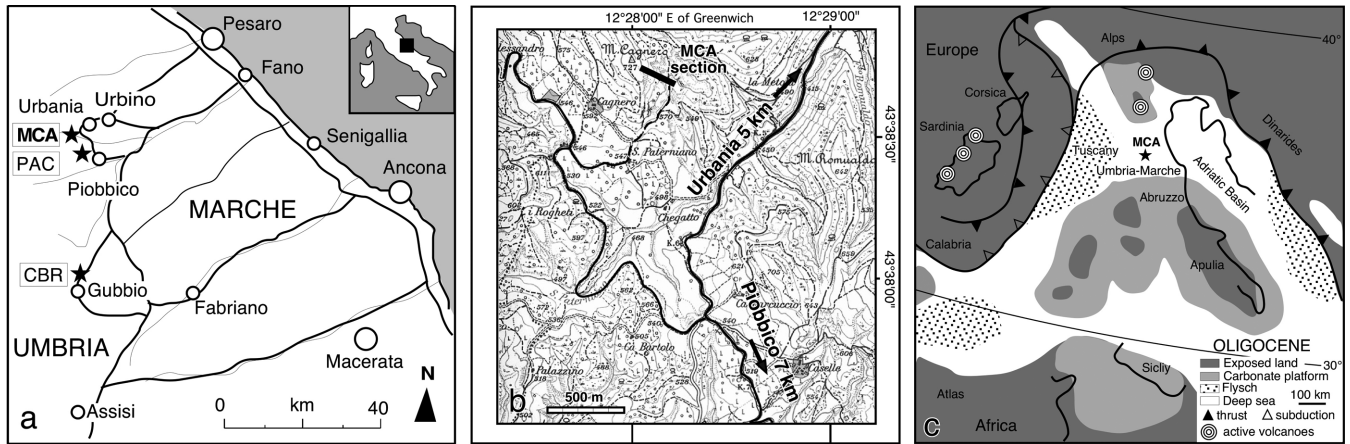


Figure 1. (a) Simplified roadmap of the Umbria–Marche region with the location of the Oligocene sections of Contessa Barbetti Road (CBR), Pieve d'Accinelli (PAC), and Monte Cagnero (MCA); (b) Precise location of the MCA section; (c) Location of the MCA section on an Oligocene paleogeographic map simplified and slightly amended from Dercourt et al. (1993).

Location and Geological Setting

The 86-m-thick MCA section is located on the southeastern slope of Monte Cagnero (727 m above sea level; coordinates: 43°38'47.81"N–12°28'03.83"E), a few kilometers to the SW of the town of Urbina (Fig. 1b) (Coccioni et al., 2008).

The MCA section consists of a rhythmic alternation of gray-bluish, pelagic marls, calcareous marls and marly limestones belonging to the Scaglia Cinerea Formation whose carbonate component is mainly made up of planktonic foraminiferal tests and calcareous nannofossils, and the siliciclastic component is represented by terrigenous clay and silt (Fig. 2). The MCA section is free of turbidites and any other manifestation of basinal instability. Rare intercalations of thin, biotite-rich volcano-sedimentary horizons are found and provide the means for direct radioisotopic dating of the magnetobiostratigraphy (Coccioni et al., 2008, and references therein) (Fig. 2).

Stratigraphy

The MCA section as GSSP for the base of the Chattian Stage (Fig. 2) is the continuation of the underlying upper Priabonian to middle Rupelian section, whose integrated stratigraphy and astrochronologic calibration were determined by Hyland et al. (2009). However, while working on the Rupelian to Chattian portion of the section, carefully re-checking the stratimetry of Hyland et al. (2009), we found an error of one meter in the measured thickness at about meter level 125, in a stretch of the section that crosses, for a few meters, a dirt road. Consequently, we had to add one meter to the section of Coccioni et al. (2008) (Fig. 2). A stratigraphic synthesis of the GSSP section of MCA is shown in Figure 3, and described in detail below.

Lithostratigraphy

The rhythmic bedding of the pelagic Scaglia Cinerea Formation is clearly manifested in the well-exposed southeastern slope of Monte Cagnero as a stepped topography, with prominent benches formed by

hard limestone layers protruding out of recessive, softer marls (Fig. 4a). These calcareous benches are organized in cycles with a mean wavelength of about 5.6 m, as revealed by a band pass filtered curve obtained from a high-resolution magnetic susceptibility series, where low susceptibility represents calcareous marls and marly limestones, and high susceptibility represents marls with a higher terrigenous (magnetic) content (Fig. 4a). The GSSP for the base of the Chattian Stage is located at the highest common occurrence (HCO) of the planktonic foraminifer *Chiloguembelina cubensis*, i.e., the O4/O5 planktonic foraminiferal zonal boundary of Wade et al. (2011), at meter level 197, thirty cm below the base of a prominent calcareous bed within an otherwise recessive, relatively thick marly interval (Figs. 2 and 4b).

Biostratigraphy and Paleodepth

The pelagic sediments of the MCA section contain abundant calcareous nannofossils and planktonic foraminifera, with common benthic foraminifera and rare ostracods. Palynomorphs are abundant as well. Macrofossils were not found. Ichnofossils are common and mainly represented by *Zoophycos* and *Planolites*.

Planktonic foraminifera

Planktonic foraminiferal analysis is based on 173 samples collected at an average sampling resolution of 50 cm (Fig. 2) that corresponds to ~37.5 kyr and processed to obtain washed residues greater than 32 μm using standard micropaleontological techniques. Planktonic foraminiferal assemblages are rich, diverse, and well preserved throughout the section. The planktonic foraminiferal standard zonation of Wade et al. (2011) was applied, with the taxonomy of *Paragloborotalia opima* and *Paragloborotalia nana* assessed following the taxonomic size-based concepts established by Wade et al. (2016). Primary and secondary events were recognized throughout O2 to O6 Zones (Figs. 2, 3, and 5; Table 1).

A progressive gigantism of *P. opima* with increasing abundance of the larger forms from the lower part of Chron C10n to the lower part of Chron C9n was recognized by Wade et al. (2016) through the mid-late Oligocene at Site U1334 (eastern equatorial Pacific Ocean). A

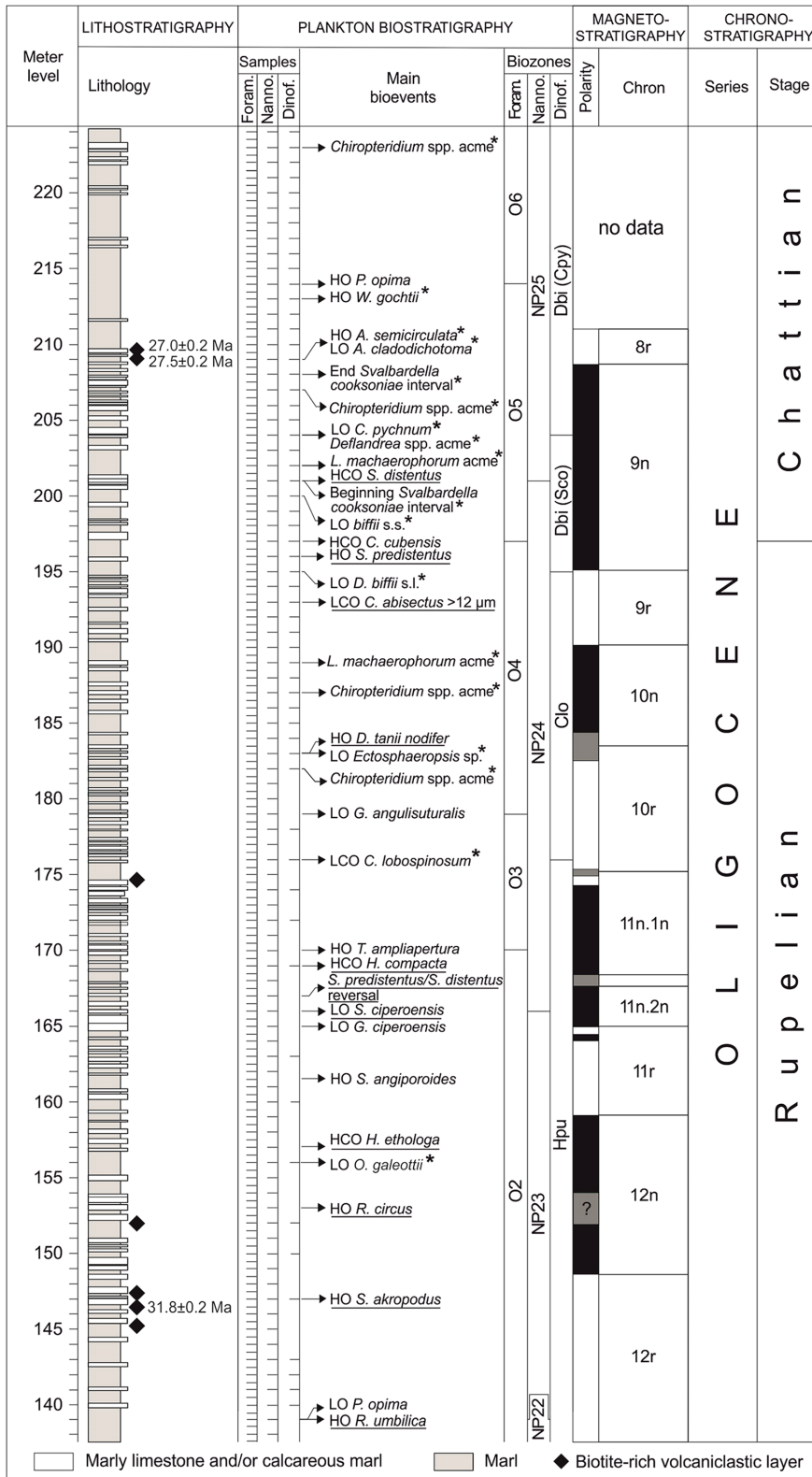


Figure 2. Litho-, magneto-, bio- and chronostratigraphy of the Monte Cagnero (MCA) section (modified after Coccioni et al., 2008; Hyland et al., 2009; Pross et al., 2010, and this work). Underlined – calcareous nannofossil events; asterisk – dinoflagellate cyst events; unmarked – planktonic foraminiferal events. Biozones after Wade et al. (2011) for planktonic foraminifera, Martini (1971) for calcareous nannofossils, and Pross et al. (2010) for dinoflagellate cysts. Ar/Ar ages from volcaniclastic biotite recalculated with the new age of the Fish Canyon Tuff standard by Kuiper et al. (2008). According to our astrochronology, the MCA section spans ~6 Myr with an average sedimentation rate of 1.45 cm/kyr.

peak size of 680 µm is recorded in the lower part of Chron C9n very close (~2 m apart) to the Top occurrence of *C. cubensis* that according to King and Wade (2017) represents the highest (last) occurrence of this species. Following this peak, the maximum size of *P. opima* decreases with its extinction event that marks the base of the late Oligocene Zone O6. Interestingly, a very similar pattern in abundance and size of *P. opima* is recognizable at MCA, with a peak size recorded in the lower part of Chron C9n very close (~1 m apart) to the HCO of *C. cubensis* and followed by rare and sporadic occurrences of *P. opima* up to its demise at meter level 214. After that only the smaller *P. nana* forms continue into Zone O6.

The degree of reliability of the HCO of C. cubensis as the global criterion for the Rupelian/Chattian boundary

The highest occurrence (HO) of *C. cubensis* has traditionally been used as the criterion for the recognition of the Rupelian/Chattian boundary worldwide and calibrated to the top of Subzone P21a associated in between Chrons C9n and C10n (e.g., Berggren et al., 1995; Luterbacher et al., 2004, and references therein), although sporadic occurrences in reduced numbers of this species have been reported through the upper Oligocene (Hess et al., 1989; Leckie et al., 1993). However, *C. cubensis* is a small species that may be easily reworked, as argued by Poore et al. (1982) and Poore (1984) for its occasional occurrences within the upper Oligocene.

Van Simaey et al. (2004) suggested that the HO of *C. cubensis* was globally diachronous and likely controlled by paleolatitude and paleobiogeography, therefore questioning its validity as a reliable global Rupelian/Chattian boundary criterion.

In view of these problems with the decline in abundance of *C. cubensis*, Berggren and Pearson (2005) used the HCO of this species rather than its HO to mark the Rupelian/Chattian boundary and placed this biohorizon at the boundary between the planktonic foraminiferal Zones O4 and O5 corresponding to the boundary between Subzones P21a and P21b of Berggren et al. (1995) and within Chron C10n.

Wade et al. (2007) documented that the HCO of *C. cubensis* is synchronous at a number of Drilling Deep Sea Drilling Project (DDSDP) and Ocean Drilling Program (ODP) sites and may be correlated directly or indirectly with Chron

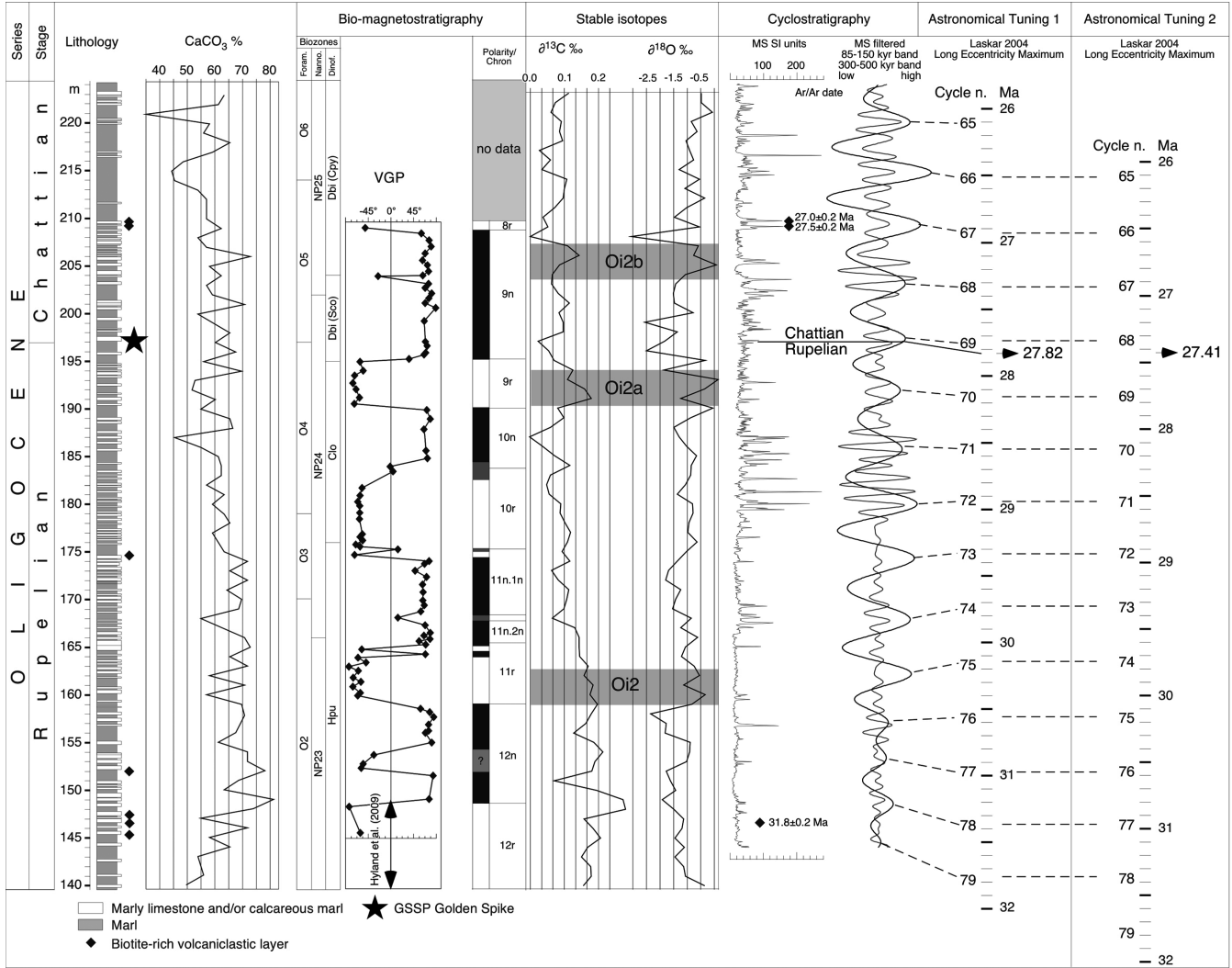


Figure 3. Integrated stratigraphic model of the Monte Cagnero (MCA) section. Litho- and magnetobiostratigraphy, stable isotope stratigraphy, glacial events (Oi2, Oi2a, and Oi2b), and radioisotope geochronology modified after Coccioni et al. (2008). Biozones after Wade et al. (2011) for planktonic foraminifera, Martini (1971) for calcareous nannofossils, and Pross et al. (2010) for dinoflagellate cysts. Astronomical tuning based on the long and short eccentricity cycles of Laskar et al. (2004) is from this work (see text for explanation).

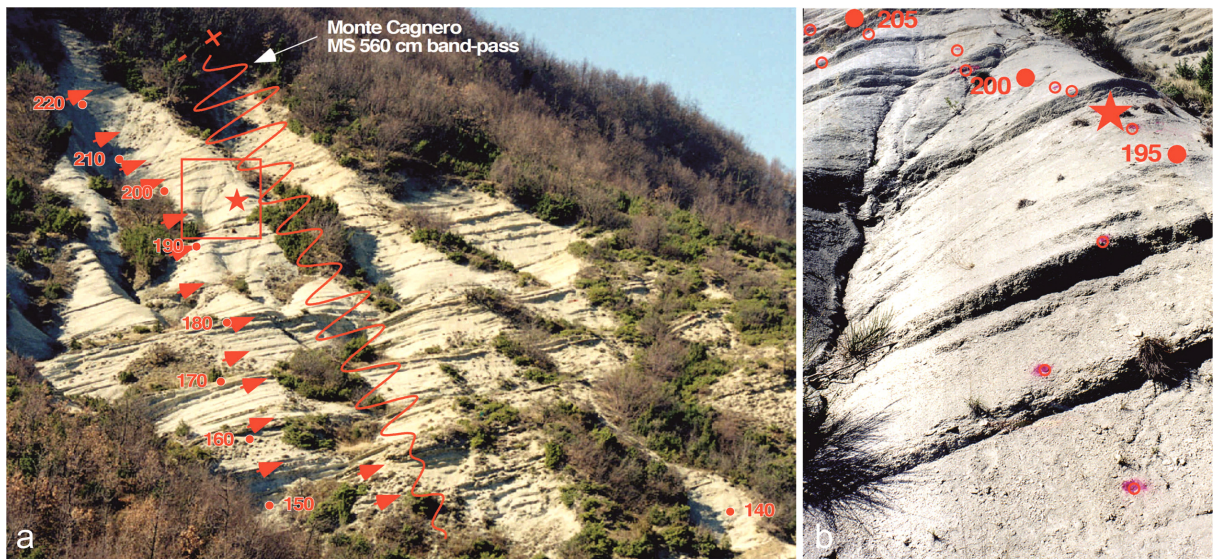


Figure 4. (a) Panoramic view of the Monte Cagnero section with superimposed band-pass curve representing the long eccentricity cycle, and arrows indicating corresponding limestone layers bundles; (b) Location of the Rupelian/Chatthian boundary at meter level 197, as indicated by the star.

C10n. Contrary to Van Simaëys et al. (2004), these results indicate that the HCO of *C. cubensis* is a robust bioevent for the O4/O5 zonal boundary and correlates to Chron C10n, consistent with many other deep-sea sections.

King and Wade (2017) documented contrasting results in the Pacific Ocean, with the abrupt termination of *C. cubensis* close to the top of Subchron C10n.1n at ODP Site 1237 and two peaks in abundance of this species recorded just before its demise in the lower part of Chron C9n at Integrated Ocean Drilling Program (IODP) Site U1334. Moreover, King and Wade (2017) evidenced no latitudinal control on the extinction level of *C. cubensis* that was possibly controlled by ecological preferences. These findings would therefore question the legitimacy of the biostratigraphic utility of the extinction of *C. cubensis* as a reliable boundary marker for the Rupelian/Chattian boundary.

From all the above, it seems appropriate to provide a precise definition of the HCO of *C. cubensis*. We regard the HCO of an index species as the biohorizon that, on the basis of discrete changes in the abundance record, marks the end of the consistent (continuous and/or common) distribution of this species prior to its extinction. Our abundance record of *C. cubensis* from the MCA section is well suited for this purpose because it is based on high-resolution sampling and covers a long time interval (~4.2 Myr) (Figs. 3 and 5). At MCA section, *C. cubensis* is consistently common in most of the samples from the base of the section up to meter level 201, with a highest value of 30.5% and an average value of 10.3% (Fig. 5). From meter level 198, however, this species becomes consistently rare and above the meter level 202 it occasionally occurs with a very few specimens. Similar decline has been observed into the upper Oligocene of the western subtropical North Atlantic and the equatorial western tropical Pacific (Leckie et al., 1993), and at the low latitudes of Syria (Kucenjak et al., 2014).

Accordingly, we adopted the unambiguous criterion of using the consistent and continuous reduction in abundance of *C. cubensis* to less than one third of the average value of the abundance of this species (counted on at least 300 specimens and the average value should be not less than 3%) as threshold value to define its HCO. Following this, the HCO of *C. cubensis*, is constrained in between the upper half of Chron C10n and the lower part of Chron C9n, where its abundance decreases from 19% to 3.6% (Fig. 5). We placed the HCO of *C. cubensis* at meter level 197 in the lower part of Chron C9n (Fig. 5). Interestingly, although with different percentage abundances, similar records have been reported from the southern Spain (up to 22% with an average value of ~3.5%; Alegret et al., 2008), the North Adriatic (up to 19.5% with an average value of ~5%; Kucenjak et al., 2014), and the south-east (up to 14% with an average value of ~5%) and equatorial Pacific Ocean (up to 13% with an average value of ~4%; King and Wade, 2017).

More importantly, as mentioned above, the abundance and size distribution of *P. opima-nana* plexus specimens at MCA closely mirrors that recognized at Site U1334. This provides evidence that the Top occurrence of *C. cubensis* at Site U1334 and the HCO of *C. cubensis* at MCA are likely the same bioevent, which is well constrained in the lower part of Chron C9n. This also shows that there are no substantial differences in extinction horizon of this species and strongly supports the reliability of the HCO of *C. cubensis* as the global criterion for the Rupelian/Chattian boundary.

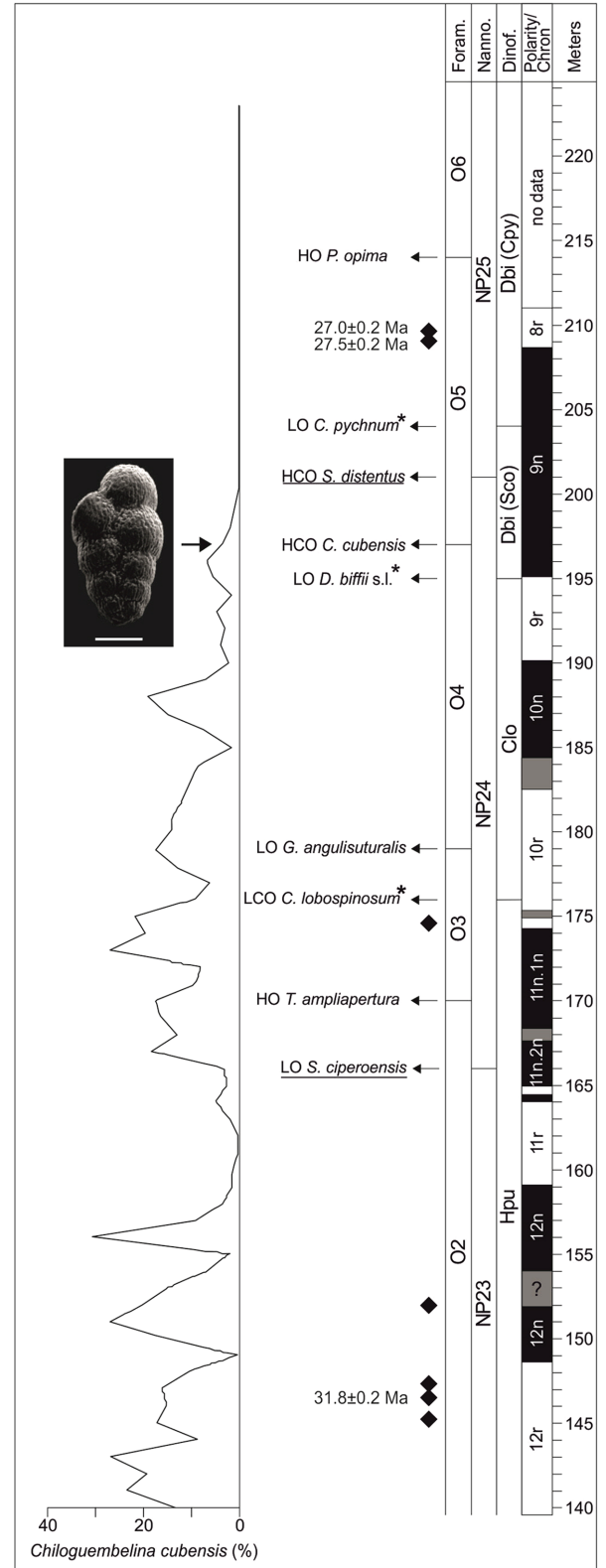


Figure 5. Abundance record of the planktonic foraminifer *Chiloguembelina cubensis* at the Monte Cagnero section through counting at least 300 specimens of planktonic foraminifera from the fraction larger than 63 µm. The highest common occurrence (HCO) of this species, which marks the base of Zone O5 of Wade et al. (2011), occurs at meter level 197. Magnetobiostratigraphy, biozonal markers, and Ar/Ar ages from volcanoclastic biotite as in Figure 4. Scale bar 50 µm.

Benthic foraminifera

The benthic foraminiferal analyses were performed on the fraction greater than 63 μm of 34 selected samples among those used for the planktonic foraminiferal investigation (Fig. 2). The benthic foraminiferal assemblages of the MCA section are highly diverse and dominated by calcareous-hyaline taxa, indicating deposition well above the calcite compensation depth. The occurrence of benthic foraminiferal paleobathymetric index taxa (e.g., *Bulimina truncana*, previously reported as *Bulimina alazanensis*, *Bulimina jarvisi*, *Cibicidoides mexicanus*, *Cibicidoides mundulus*, *Eggerella bradyi*, *Globocassidulina subglobosa*, *Hansenisca soldanii*, *Hanzawaia ammophila*, *Heterolepa bradyi*, *Karrerella bradyi*, *Oridorsalis umbonatus*, *Planulina costata*, *Planulina renzi*, *Pullenia bulloides*, and *Pullenia quinqueloba*) was evaluated (e.g., Van Morkhoven et al., 1986; Katz et al., 2003). The plankton/benthos ratio (> 95%), the presence of paleobathymetric diagnostic taxa and the low percentages of agglutinans constrain the deposition of the MCA section to a well oxygenated, stable, bathyal environment, characterized by a depth very likely ranging between 500 and 1000 m (upper-middle bathyal) or somewhat deeper of 1000 m (upper part of lower bathyal) following the paleobathymetric assignment of van Morkhoven et al. (1986) and Katz et al. (2003), respectively.

Calcareous nannofossils

The calcareous nannofossil analysis is based on 85 samples that were collected at an average sampling resolution of ~ 1 m corresponding to ~ 75 kyr (Fig. 2). Light microscope techniques at 1000x magnification were used for examination of smear slides prepared using standard methods. Calcareous nannofossils assemblages are generally abundant and moderately etched and overgrown. The calcareous nannofossil zonal boundaries through NP23–NP25 are traced according to the standard zonation of Martini (1971) with slight modification for the definition of the NP24/NP25 zonal boundary, for which the HCO of *Sphenolithus distentus* is used here instead of its HO (Fig. 2 and Table 1). In fact, this taxon is rare and scattered in the upper part of its range (Olafsson and Villa, 1992; Maiorano and Monechi, 2006), while the final decrease in abundance enables an easy detection of the HCO of *S. distentus* that seems to be a more reliable biostratigraphic signal than the HO. In addition to a few standard and well-known primary events, several secondary events were recognized from the uppermost part of NP22 to NP25 Zones (Figs. 2, 3, and 5; Table 1).

Dinoflagellate cysts

Dinoflagellate cyst (dinocysts) analysis is based on 58 samples (Fig. 2) that were collected at an average sampling resolution of 1.45 m (Fig. 2) corresponding to ~ 110 kyr and processed using standard palynological techniques. All samples contain rich, diverse and well-preserved palynological assemblages dominated by dinocysts and bisaccate pollen. The dinocyst zonation of Pross et al. (2010) was applied. Primary and secondary events were recognized through Hpu to Dbi Zones (Figs. 2, 3, and 5; Table 1). Acmes of *Chiropteridium* spp., *Lingulodinium machaerophorum* and *Deflandrea* spp. occur at different stratigraphic levels, with a marked increase in abundance of

Svalbardella cooksoniae between meter levels 201 and 208 (Fig. 2 and Table 1).

Magnetostratigraphy

The magnetic properties of the Scaglia Cinerea of the MCA section were described and documented in detail by Coccioni et al. (2008). The magnetostratigraphic interpretation of this section is relatively straightforward and allows the recognition of all the polarity reversals from the top of Chron C12r to the base of Chron C8r, although some uncertainty remains in interpreting short polarity subzones within Chrons C12n, C11n, and C9n and the exact reversal level of Chrons C10r–C10n (Fig. 3). The succession of Chrons C13r up to C8r at the MCA section is fairly consistent with that of the classic Gubbio section at Contessa (Fig. 6). However, the polarity reversal pattern and the relative thicknesses of the Chrons of the U–M pelagic sections exhibit some differences with the patterns of marine magnetic anomalies across the South Atlantic and North Pacific Oceans and the ODP Sites 1218 and 1219 (equatorial Pacific) (Fig. 6). Despite these uncertainties, the GSSP for the base of the Chattian at meter level 197 undoubtedly falls within the lower part of a normal polarity zone, which we interpret as Chron C9n (Fig. 2).

Chemostratigraphy

The Oligocene stretch of the numerical strontium isotope time scale is characterized by a steep, monotonic increase of the $^{87}\text{Sr}/^{86}\text{Sr}$ from about 0.70782 at the base of the Rupelian to about 0.70828 at the top of the Chattian (Fig. 7a). From the combined data of MCA and PAC sections (Coccioni et al., 2008), a precise $^{87}\text{Sr}/^{86}\text{Sr}$ value of 0.708105 ± 0.00001 can be given to the base of the Chattian at MCA section (Fig. 7b).

Carbon ($\delta^{13}\text{C}$) and oxygen ($\delta^{18}\text{O}$) stable isotopes were measured through the MCA section by Coccioni et al. (2008) on bulk rock samples. A sample from meter level 197, the Rupelian/Chattian boundary, records a relatively low $\delta^{13}\text{C}$ value of 0.05‰ and a relatively low $\delta^{18}\text{O}$ value of -2.0 ‰. These isotopic minima fall within an interval of relatively low $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values, which is bracketed by two intervals, centered at meter levels 192 and 206, of relatively high isotopic values, which were respectively interpreted by Coccioni et al. (2008) as the Oligocene glacial events Oi2a and Oi2b of Pekar and Miller (1996) and Miller et al. (1991), respectively (Fig. 3). Accordingly, the base of the Chattian would precede the Oi2b event, i.e., the Oligocene Glacial Maximum (“OGM”) of Van Simaëys (2004), by some 500 kyr (Fig. 3).

Cyclostratigraphic Analysis

Magnetic susceptibility (MS) was measured on 1600 ground rock samples collected at 5-cm resolution from meter level 144 to meter level 224, with a Bartington MS2 dual frequency sensor on low frequency (0.465 kHz) and $\times 0.1$ sensitivity, as a proxy for our cyclostratigraphic analysis. Three measurements were taken on each sample and averaged to a single value in SI magnetic susceptibility unit. We conducted spectral analyses of the MCA MS proxy series in Matlab using algorithms modified from Muller and MacDonald (2000) in agreement with Hyland et al. (2009).

Table 1. Planktonic foraminiferal, calcareous nannofossil, and dinocyst events, and Chron boundaries, with their numerical ages for the Monte Cagnero section

Biostratigraphic event and Chron	Meter level	Martini (1971)	Okada and Bukry (1980)	Berggren et al. (1995)	Wade et al. (2011)	Pross et al. (2010)	Gradstein et al. (2012) GTS age (Ma)	MCA Solution 1 Astrochronological age (Ma)	MCA Solution 2 Astrochronological age (Ma)
<i>Chiropteridium</i> spp. acme	223							25.90	25.49
HO <i>Paragloborotalia opima</i>	214			P21b/P22	O5/O6		26.91	26.60	26.17
HO <i>Wetzeliella gochtii</i>	213							26.71	26.23
HO <i>Areoligera semicirculata</i>	209							26.97	26.54
LO <i>Artemisiocysta cladodichotoma</i>	209							26.97	26.54
C9n/C8r	208.7						26.41	26.95	26.53
End <i>Svalbardella cooksoniae</i> interval	208							27.03	26.73
<i>Chiropteridium</i> spp. acme	207							27.09	26.67
<i>Deflandrea</i> spp. acme	204							27.32	26.93
LO <i>Caligodinium pychnum</i>	204					Dbi (Sco)/ Dbi (Cpy)		27.32	26.93
<i>Lingulodinium machaerophorum</i> acme	202							27.46	27.08
Beginning <i>Svalbardella cooksoniae</i> interval	201							27.52	27.12
HCO <i>Sphenolithus distentus</i>	201	NP24/ NP25	CP19a/ CP19b				26.80	27.52	27.12
LO <i>Distatodinium biffii</i> s.s.	200							27.55	27.15
HCO <i>Chiloguembelina cubensis</i> (Rupelian– Chattian boundary)	197			P21a/ P21b	O4/O5		28.10	27.82	27.41
HO <i>Sphenolithus predistentus</i>	196							27.90	27.52
C9r/C9n	195.1						27.41	27.94	27.56
LO <i>Distatodinium biffii</i> s.l.	195					Clo/Dbi (Sco)		27.96	27.58
LCO <i>Cyclicargolithus abisectus</i> > 12 µm	193							28.08	27.72
C10n/C9r	190.2						27.83	28.26	27.88
<i>Lingulodinium machaerophorum</i> acme	189							28.35	27.94
<i>Chiropteridium</i> spp. acme	187							28.47	28.09
C10r/C10n	183.5						28.40	28.69	28.24
HO <i>Discoaster tanii nodifer</i>	183							28.74	28.31
LO <i>Ectosphaeropsis</i> sp.	183							28.74	28.31
<i>Chiropteridium</i> spp. acme	182							28.78	28.36
LO <i>Globigerina angulisurealis</i>	179			P20/P21a	O3/O4		29.19	29.00	28.62
LCO <i>Chiropteridium lobospinosum</i>	176					Hpu/Clo		29.23	28.85
C11n.1n/C10r	175.2						29.18	29.32	28.92
HO <i>Turborotalia ampliapertura</i>	170			P19/P20	O2/O3		30.28	29.58	29.18
HCO <i>Helicosphaera compacta</i>	169							29.73	29.33
C11n.1r/C11n.1n	168.4						29.45	29.77	29.28
C11n.2n/C11n.1r	167.6						29.50	29.81	29.30
<i>Sphenolithus predistentus/Sphenolithus distentus</i> reversal	167							29.85	29.37
LO <i>Sphenolithus ciproensis</i>	166	NP23/ NP24	CP18/ CP19a				29.60	29.88	29.47
C11r/C11n.2n	165						29.99	29.95	29.54
LO <i>Globigerina ciproensis</i>	165							29.95	29.54
HO <i>Subbotina angiporoides</i>	161.5							30.20	29.83
C12n/C11r	159.1						30.59	30.43	30.03
HCO <i>Helicosphaera ethologa</i>	157							30.58	30.20
LO <i>Oligokolpoma galeottii</i>	156							30.62	30.27
HO <i>Reticulofenestra circus</i>	153							31.00	30.63
C12r/C12n	148.6						31.00	31.37	30.97
HO <i>Sphenolithus akropodus</i>	147							31.48	31.06
HO <i>Reticulofenestra umbilica</i>	139	NP22/ NP23	CP16c/ CP17				32.02	32.30	
LO <i>Paragloborotalia opima</i>	139							32.30	

The ages for the HO of *Reticulofenestra umbilica* and the LO of *Paragloborotalia opima* are after Hyland et al. (2009).

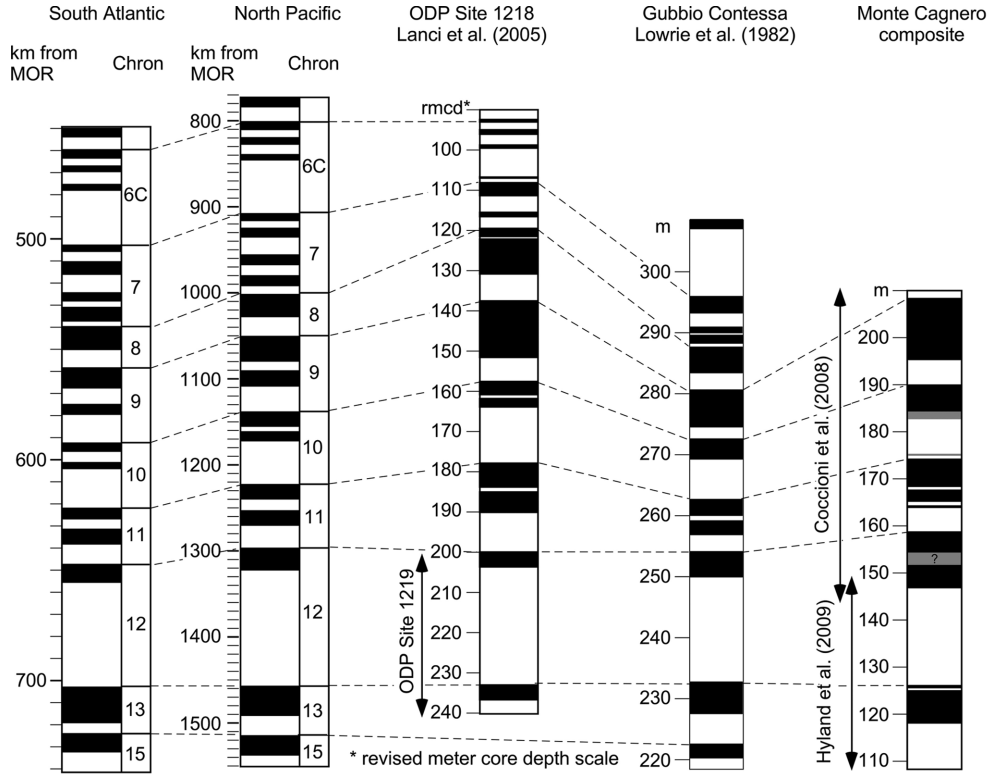


Figure 6. Correlation of the Eocene–Oligocene magnetic polarity pattern among the transects across the South Atlantic and North Pacific oceans, the cores from ODP Sites 1218 and 1219, the classic Gubbio section at Contessa, and the composite section of Monte Cagnero.

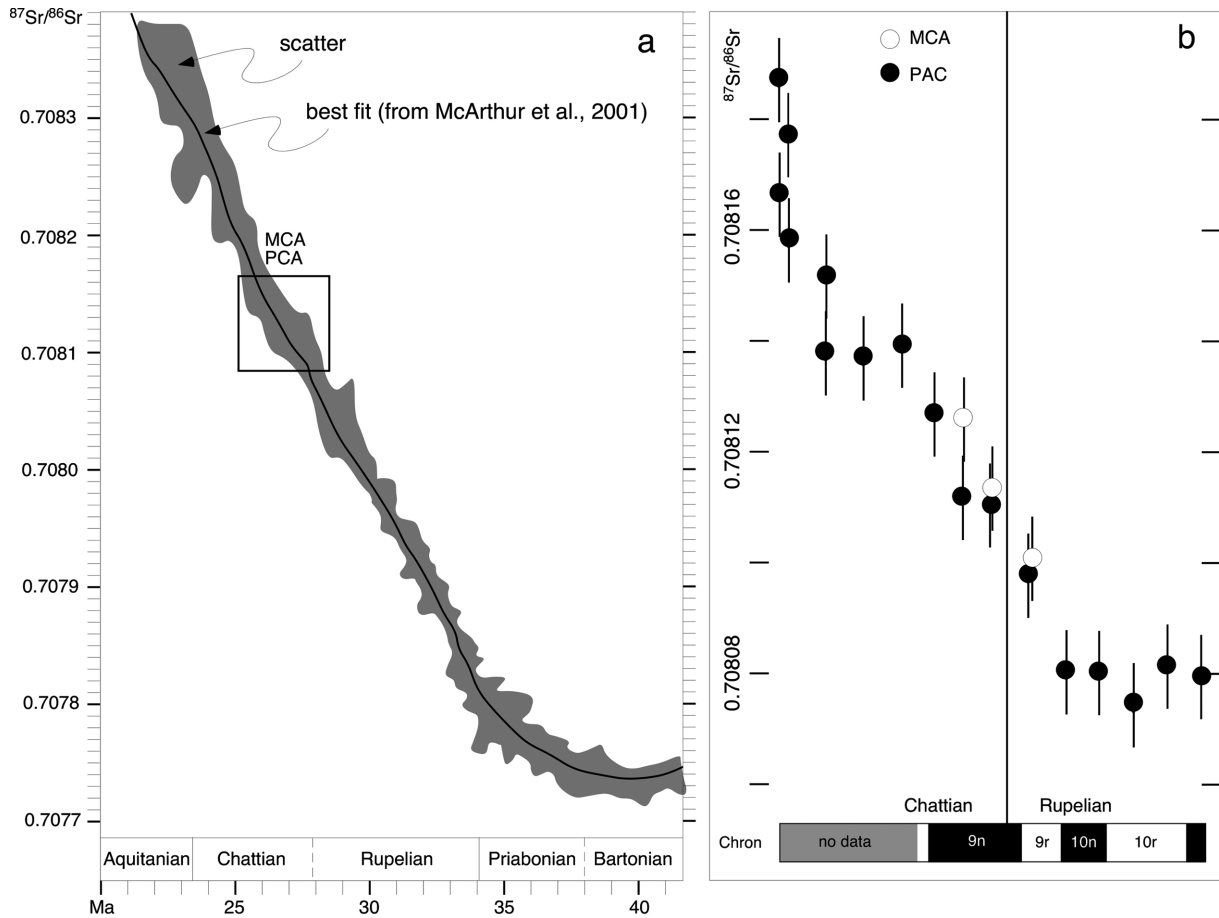


Figure 7. (a) Location of the Monte Cagnero (MCA) and Pieve d'Accinelli (PAC) sections in the strontium isotope curve of McArthur et al. (2001). (b) Strontium isotope composition across the Rupelian/Chattian boundary in the MCA and PAC sections (from Coccioni et al., 2008).

The MS data contain an overall up-section trend toward higher values, reflecting a general increase in terrigenous influx through time, which is consistent with a general decrease in CaCO_3 content (Fig. 3).

We evaluated the statistical significance of the spectral peaks by generating a 95% confidence level from a Monte Carlo noise simulation (Muller and MacDonald, 2000). Prominent, statistically significant peaks related to one another by the same ratios found between Milankovitch cycles, are assumed to represent eccentricity, obliquity (tilt), and precession, making it possible to correlate meter level to age through the calculation of an average sedimentation rate of 1.45 cm/kyr for the MCA section (Figs. 2 and 8a). The persistence of the spectral peaks through the stratigraphic section was evaluated through the use of an evolutionary, or sliding window, fast Fourier transforms (FFT). In this approach, a subset or window of the data is analyzed using the same FFT; the window then moves up through the time series. The window size was chosen to be ~26 m, which is large enough to enable the detection of short and long eccentricity signals given the preexisting sedimentation rate from the overall FFT power spectrum (Fig. 8b). The linearity of the signals in the evolutionary FFT plot is an indica-

tion of a fairly constant sedimentation rate throughout the section. We then applied a Hilbert transform to the presumed orbital signals of precession and eccentricity, which resulted in an enveloping curve that graphically illustrates the amplitude modulation. We then applied a FFT to this enveloping curve to identify the spectral character of the amplitude modulation. A similar amplitude modulation analysis of the Laskar2004 ETP orbital parameters calculated by Laskar et al. (2004 – Laskar2004 hereafter) allowed us to compare our presumed orbital signals with the mathematically predicted orbital cycles. The frequency signals in the MCA section contained in a 17–27 kyr band (i.e., precessional cycle) do show modulations at lower frequencies compatible with Laskar2004 eccentricity cycles (Fig. 8c). Similarly, frequency signals in a 85–150 kyr band (short eccentricity) show the expected modulation frequency of the ~400 kyr long eccentricity cycle compatible with Laskar2004 (Fig. 8d). We did not run a Hilbert transform for the presumed obliquity signal because the power spectrum appears too noisy with multiple peaks in the band of 35–45 kyr.

Astronomical tuning and ultimate dating of the MCA section were accomplished through the use of broad band-pass filtering. Variations

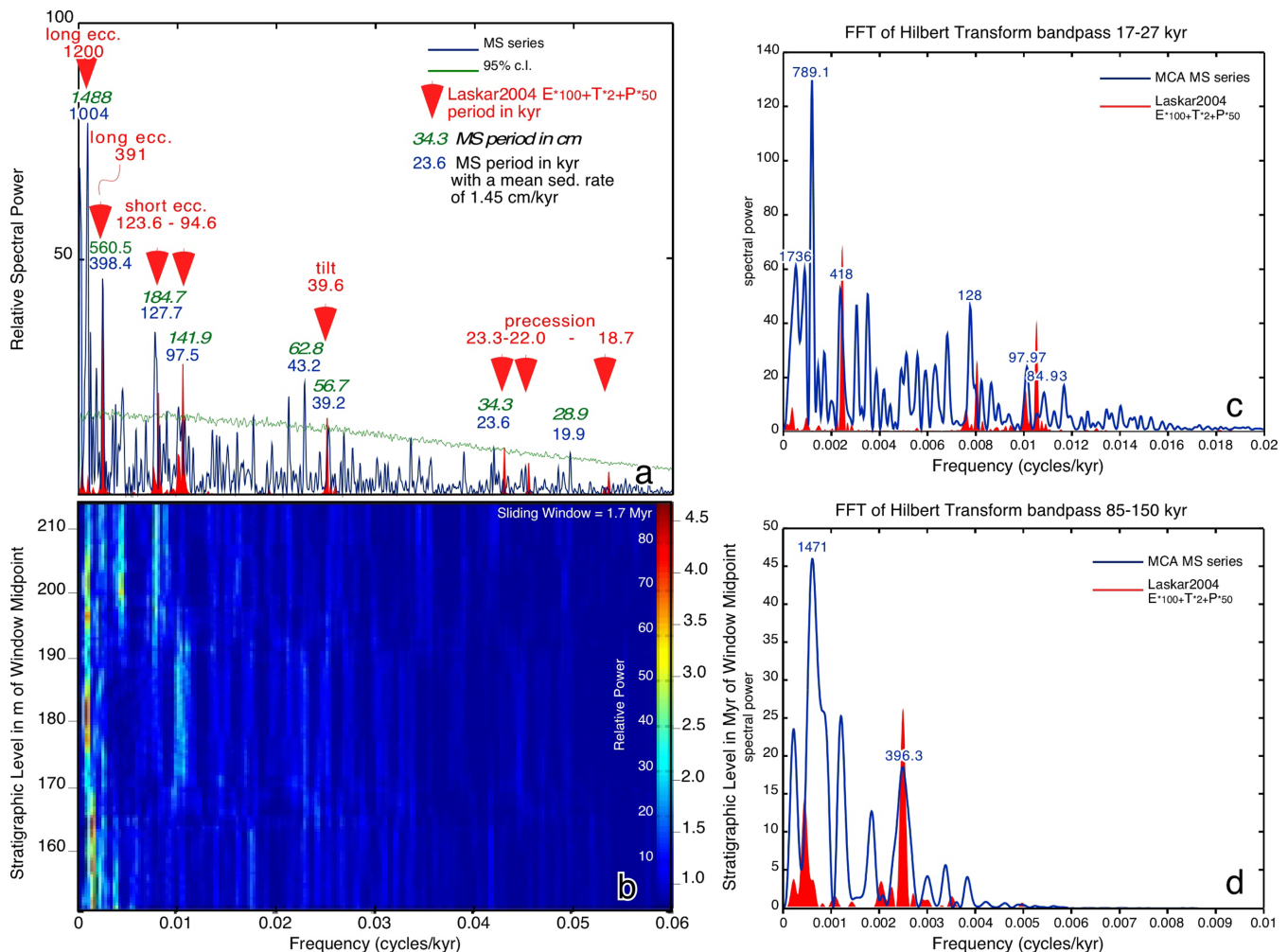


Figure 8. Spectral analysis synthesis of the magnetic susceptibility (MS) time series from the Monte Cagnero section: (a) Fast Fourier periodogram of the Monte Cagnero MS series compared to the Laskar et al. (2004) ETP series; (b) Evolutionary or Sliding Window FFT of the Monte Cagnero MS series in the time domain; (c) Hilbert transform periodogram of the precession frequency band of the Monte Cagnero MS series compared to the same frequency band from Laskar et al. (2004) ETP series; (d) Hilbert transform periodogram of the short eccentricity frequency band of the Monte Cagnero MS series compared to the same frequency band from Laskar et al. (2004) ETP series.

in MS proxy series between 85 and 150 kyr were isolated so as to emphasize the variance in the short eccentricity band, which, in our MCA series, includes the strongest and most consistent frequency peaks. These filtered data were then used for correlation via constrained pattern matching with the theoretical eccentricity curve of Laskar et al. (2004). The uncertainties in astronomical dating depend on (1) the astronomical solution applied; (2) the assumption of an overall constant sedimentation rate; and (3) the assumption that there is no lag between orbital forcing and its sedimentary expression (Kuiper et al., 2008). Laskar et al. (2004) estimated a maximum chrono-

logical error in their calculations of 30 kyr over the last 50 Myr. If we assume our tuning to be correct, then the combined error in astronomical solution and stratigraphical calibration on MCA section will be about ± 50 kyr.

Astronomical Tuning and Age Model

We performed the actual tuning by correlating the maxima of a broad 4.4–7.3 m band pass curve of the MS proxy series with the maxima of the long eccentricity curve of Laskar2004, as shown in

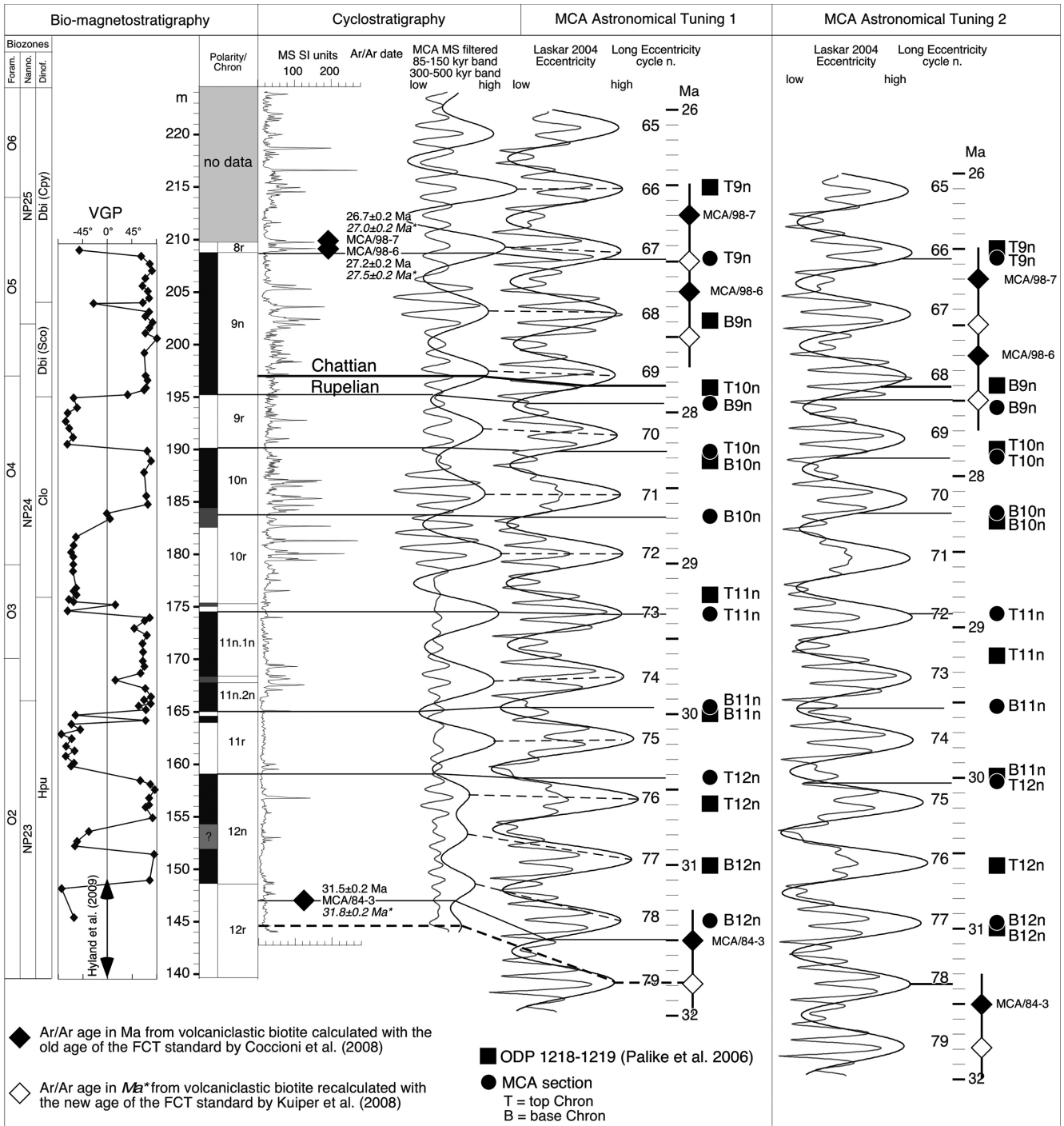


Figure 9. Two alternative astrochronological calibrations of the Monte Cagnero MS series based on the long eccentricity cycle tuned with the solution of Laskar et al. (2004). FCT: Fish Canyon Tuff.

Figure 9. The phase relation was assumed to be such that maxima in MS (i.e., maxima in terrigenous, magnetic content) corresponded to maxima in eccentricity, which would imply a correspondence between maxima in temperature and humidity in a glacio-eustatic cyclic pattern. Once a fairly clear long and short eccentricity cyclicity has been determined through the MCA section via FFT spectral analysis of the MS proxy series, the problem arose of how to anchor our ~400 kyr (long eccentricity) cycle curve to the long eccentricity model calculated by Laskar et al. (2004), which has to be considered astrochronologically precise (to the 1 kyr) and accurate.

One way to perform this anchorage is to consider the radioisotopic ages obtained from volcanoclastic layers found at the bottom and the top of the MCA section (Fig. 9). Ar/Ar radioisotopic ages from two biotite-rich volcanoclastic layers located at meter level 147 (MCA/84-3) and at meter level 209.65 (MCA/98-7), which were respectively recalculated at 31.8 Ma and 27.0 Ma by Gradstein et al. (2012) using the new age of the Fish Canyon Tuff (FCT) sanidine standard of Kuiper et al. (2008), are fairly consistent, within the 2σ analytical error, with their astrochronologic ages resulting from the tuning 1 (Fig. 9). However, these dates, regardless of their analytical precision, bear an unquantifiable geochronological uncertainty common to all radioisotopic dates and the intrinsic method of geochronological age determination. In the case of the MCA section, a volcanoclastic biotite sample collected at 209.05 m (sample MCA/98-6 in Coccioni et al., 2008) yielded a $^{40}\text{Ar}/^{39}\text{Ar}$ plateau age of 27.2 ± 0.2 Ma (or 27.5 ± 0.2 Ma using the new age of the FCT sanidine standard of Kuiper et al., 2008), which would be 500 kyr older than the age of sample MCA/98-7 found just 60 cm above it. Given a mean sedimentation rate for the upper part of the Scaglia Cinerea limestone at MCA of 1.45 cm/kyr, the difference in age between these two volcanoclastic layers would be about 40 kyr and not 500 kyr as otherwise derived from radioisotopic dating, suggesting that one or both biotite ages bear an accuracy error. This discrepancy, along with the intrinsic accuracy uncertainty in these radioisotopic ages, constitutes a problem in the attempt of tuning the MCA section with the astrochronologic model of Laskar et al. (2004). Aware of it, we anchored the continuous ~400 kyr cycle of the interval spanning the uppermost Chron C12r to the lowermost Chron C8r to the long eccentricity curve obtained by Hyland et al. (2009) continuously from the mid Chron C13r to the lower Chron 12n, practically matching a minimum of the long eccentricity cycle found, in both intervals, at about meter level 147. The long eccentricity curve of Hyland et al. (2009) was in turn anchored to the GSSP of Massignano (mid Chron C16n to mid Chron C13n), which was tuned by Brown et al. (2009). In the MCA section, the Eocene–Oligocene boundary gave an astrochronologic age of 33.925 ± 0.025 Ma, which closely corresponded to a maximum of the short eccentricity cycle placed in the old-side shoulder of the long eccentricity maximum n. 84 (where long eccentricity maximum n. 1 is located, in the astrochronologic scale of Laskar et al. (2004), at 0.22 Ma). Moreover, Hilgen and Kuiper (2009) found consistency between the astronomically tuned ages for the Eocene–Oligocene boundary of 33.9 Ma as proposed by Brown et al. (2009) and Hyland et al. (2009), the astronomically tuned age of 33.8 Ma in the ODP core at the equatorial Pacific Site 1218 as proposed by Pälike et al. (2006), and the $^{40}\text{Ar}/^{39}\text{Ar}$ sanidine ages for the boundary as derived from the volcanic ignimbrite complex in New Mexico, and for the President White Layer (PWL) ash bed in North

America (Hilgen and Kuiper, 2009), which suggested to them that the tunings were correct at the scale of the ~400 kyr cycle. In summary, the matched minimum of the ~400 kyr cycle found at about 147 m in the upper Chron C12r of the MCA section corresponds to the minimum of the long eccentricity cycle n. 78, whereas the maximum found at about 197 m would correspond to long eccentricity cycle n. 69. On the basis of fine tuning using the short eccentricity cycle (see Astronomical Tuning 1 in Fig. 9), the Rupelian/Chattian boundary at exactly meter level 197 in the MCA section yields an astrochronologic age of 27.82 Ma.

However, Tuning 1 in Figure 9 bears some major astrochronologic inconsistencies with the tuning of the magnetic polarity sequence from Chron C13n (the Eocene/Oligocene boundary) to Chron C6An.2n (uppermost Aquitanian) performed by Pälike et al. (2006) using an apparently continuous and complete, high-resolution stable isotopes proxy series from benthic foraminifera through the composite ODP Sites 1218 and 1219. These series were also anchored to the Eocene/Oligocene boundary in the upper part of Chron C13r, with an age of about 33.9 Ma. Yet, with the exception of the close astrochronologic age match of the base and top of Chron C11n between the MCA section and ODP series, all the other Chron boundaries have age differences between 200 and 600 kyr (see Tuning 1 in Fig. 9). Although we do not have an explanation for such a discrepancy, we tested an alternative tuning by sliding down the Laskar 2004 astrochronologic scale by one long eccentricity cycle with respect to the ~400 kyr cycle of the MCA section (see Tuning 2 in Fig. 9). By doing this, a good astrochronologic match better than 100 kyr between the MCA and ODP series is obtained for top and bottom of Chron C9n, top and bottom of Chron C10n, and the bottom of Chron C12n, whereas a mismatch of about 300 kyr remains for the top of C11n, and about 600 kyr for the bottom of Chron C11n and the top of Chron C12n. If Tuning 2 was considered less troubled than Tuning 1 (i.e., with less mismatches), it would imply that the age of the Eocene/Oligocene boundary would be that of the old-side shoulder of long eccentricity maximum cycle n. 83 with an age of about 33.5 Ma and not of the old-side shoulder of long eccentricity maximum cycle n. 84 with an age of 33.9 Ma, as it was determined by Brown et al. (2009) and Hyland et al. (2009), and confirmed and corroborated by $^{40}\text{Ar}/^{39}\text{Ar}$ sanidine ages by Hilgen and Kuiper (2009). In the impossibility of solving this conundrum, in Table 1 we propose astrochronologic ages for the various biostratigraphic and paleomagnetic events recorded at the MCA section according to these two alternative tunings. It follows that the age of the Rupelian/Chattian boundary found in the lower part of Chron C9n at 197 m is 27.82 Ma according to Tuning 1, and 27.41 Ma according to Tuning 2.

Historical Background on the Rupelian/Chattian Boundary and Implications for the Establishing the GSSP for the Base of the Chattian Stage at MCA

The Rupelian and Chattian historical stages, introduced respectively by Dumont in 1849 and by Fuchs in 1894, are located in the southern North Sea basin, where the Rupelian Boom Clay Formation underlies the Chattian Voort Sand Formation (Fig. 10). The Rupelian and Chattian stratotypes have been the topic of many sedimentologi-

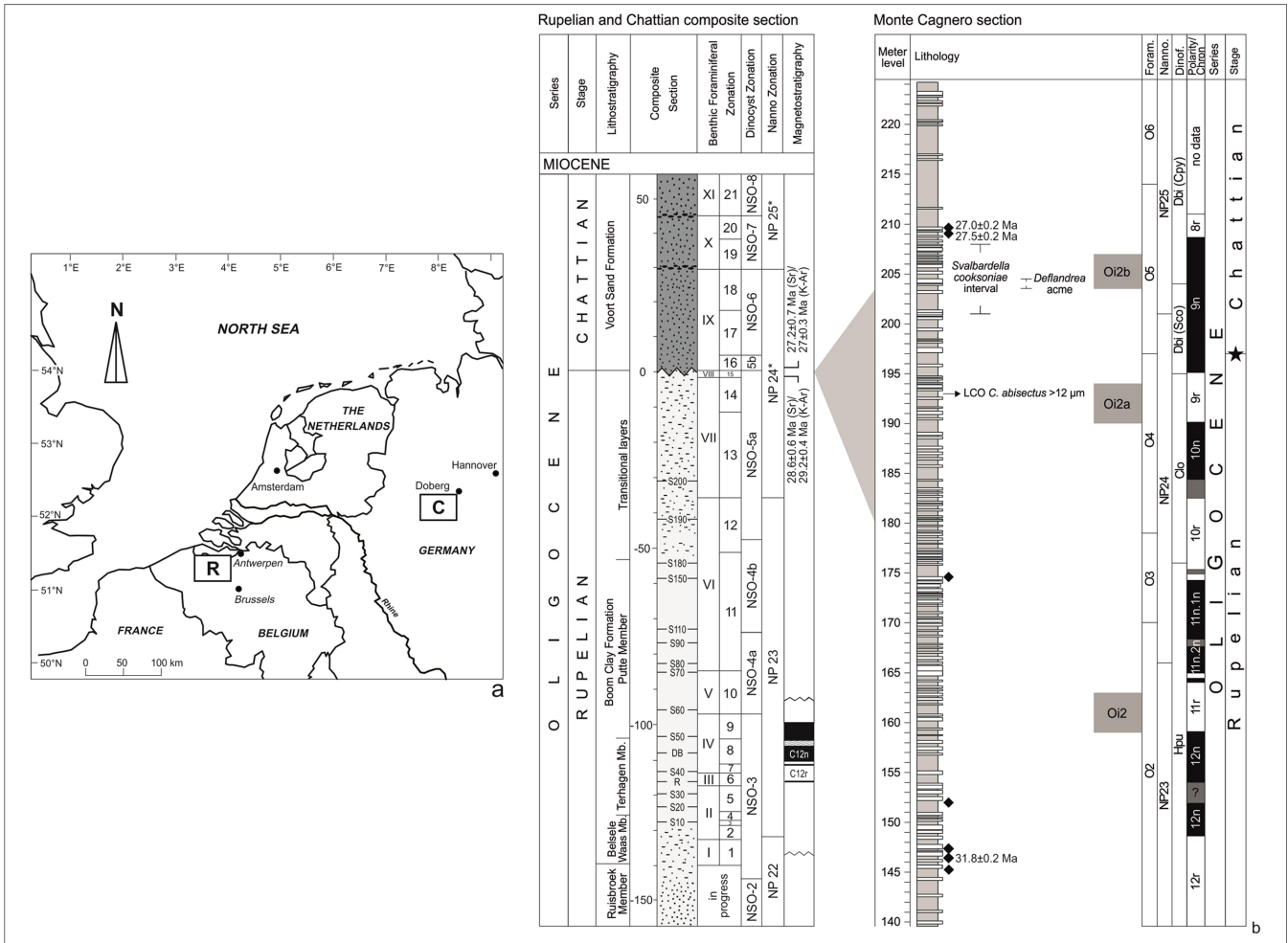


Figure 10. (a) Location of the historical Rupelian and Chattian stratotypes; (b) Correlation of the Rupelian and Chattian composite section in their type region (southern North Sea basin) (after De Man and Van Simaey, 2004; De Man et al., 2004) with the Monte Cagnero (MCA) section in central Italy (after this work). The interval IX of the benthic foraminiferal zonation marks the range of the “Asterigerina Horizon”, which defines the base of the Chattian in its original stratotype area. The star marks the location of the GSSP for the Chattian. According to the Sr and K-Ar dating provided by De Man et al. (2010), in the classical area of the Rupelian/Chattian boundary the top of the Rupelian should not be older than 29 Ma and the basal Chattian should not be younger than about 27 Ma. The grey zone indicates the hiatus that spans the classic Rupelian/Chattian boundary unconformity. Following our astrochronologic time scale it should extend through ~1.7 Myr from the middle part of Chron C10r to the middle part of Chron C9n.

cal, paleontological, paleoenvironmental, paleoclimatic, cyclostratigraphic, geochemical and geochronological studies (e.g., Vandenberghe et al., 1998, 2001, 2012; Van Simaey, 2004; De Man and Van Simaey, 2004; De Man et al., 2004, 2010; Van Simaey et al., 2004, 2005a, 2005b; Van Simaey and Vandenberghe, 2006 and references therein).

The basal Chattian deposits in the North Sea basin represent a major transgressive systems tract unconformably overlying the Rupelian stratotype (Vandenberghe et al., 1998). This unconformity, which was driven by the glacio-eustatic sea-level fall linked to the Oi2b cooling event, corresponds to a stratigraphic hiatus between the top of the Rupelian and the base of the Chattian transgression and represents a third-order sequence boundary (Van Simaey, 2004; Van Simaey et al., 2004).

The most distinct biostratigraphic marker characterizing the base of the Chattian in the North Sea Basin is the record of the *Asterigerina guerichi guerichi* benthic foraminifer acme event, known as the “Asterigerina Horizon” (Fig. 10b) (De Man et al., 2004; Van Simaey

et al., 2004, 2005b; Van Simaey and Vandenberghe, 2006).

However, accurate calibration of this bioevent and, therefore, the establishment of the age of the Rupelian–Chattian unconformity and the duration of the boundary hiatus in its type region, to the Geomagnetic Polarity Time Scale (GPTS) has been notoriously hampered by 1) the endemic nature of the North Sea Oligocene benthic foraminifera, 2) the scarcity of chronostratigraphic important calcareous plankton in the Rupelian and Chattian marginal-marine type and reference sections and 3) limited availability of reliable paleomagnetic signals and radiometric datable layers in both stratotype section and regions.

In the Boom Clay Formation, no planktonic foraminiferal zonal boundaries could be identified since some of the zonal markers were not encountered or had a peculiar range (Van Simaey and Vandenberghe, 2006; and references therein). Some magnetostratigraphic data were provided for the Rupelian stratotype by Vandenberghe et al. (2001) (Fig. 10b).

Through the Rupelian and Chattian stratotypes, the calcareous nan-

nofossils NP23 to NP25 Zones were recognized, adopting, however, the LO of *Helicosphaera recta* as the substitute marker of the base of the standard Zone NP24 and the LO of *Pontosphaera enormis* as the substitute marker of the base of the standard Zone NP25 (Van Simaey et al., 2004) (Fig. 10b). In addition, a benthic foraminiferal and a dinocyst zonation have been also established (De Man et al., 2004; Van Simaey et al., 2004, 2005b; Van Simaey and Vandenberghe, 2006) (Fig. 10b).

A striking feature coinciding with the Rupelian–Chattian unconformity in the southern North Sea Basin is the major change in paleotemperature and paleobathymetry as testified by the benthic foraminiferal assemblages, with the shallow marine to restricted marine tropical to subtropical assemblages at the base of the Chattian strongly contrasting with the deeper marine and cold to cold-temperate upper Rupelian assemblages (De Man and Van Simaey, 2004).

The nature and chronostratigraphic position of the Rupelian/Chattian boundary unconformity in its historical type region has been examined by De Man et al. (2010) by using biostratigraphy, strontium isotope dating of benthic foraminifera and K-Ar dating of glauconites (Fig. 10b). Although there is some inconsistency in the K-Ar glauconite dates, those from the upper part of the Boom Clay Formation and from lower part of the Voort Sand Formation suggest that the Rupelian/Chattian boundary lies between 29.2 ± 0.4 Ma and 27 ± 0.3 Ma (De Man et al., 2010). Sr and K-Ar dating indicate the top of the Rupelian to be approximately 29 Ma and the basal Chattian about 27 Ma (Fig. 10b).

The observations of Śliwińska et al. (2012, 2014) demonstrated that in the North Sea basin an interval with the dinocyst *Svalbardella* occurs within the “*Asterigerina* Horizon” and is synchronous with the global Oi2b isotope cooling event as recognized in the Italian sections of Coccioni et al. (2008). Interestingly, the correlations presented by Śliwińska et al. (2012, 2014) also suggest that the few isolated specimens of *Svalbardella* recognized by Van Simaey (2004) within the “*Asterigerina* Horizon” and interpreted by the same author as an effect of reworking, are on the contrary in situ. All this points to place both the *Svalbardella* interval and the Oi2b event within the lower Chattian deposits in the original type area and contradicts previous assumptions (Van Simaey, 2004; Van Simaey et al., 2004, 2005a) that both the *Svalbardella* interval and the Oi2b event correlates with the unconformity below the type Chattian.

From all the above, the hiatus that spans the classic Rupelian–Chattian unconformity should extend from the middle part of Chron C10r to the middle part of Chron C9n (Fig. 10b) and according to our astrochronologic time scale (Figs. 3 and 9) it appears to have a duration of ~ 1.7 Myr.

The GSSP at MCA at the HCO of the planktonic foraminifer *C. cubensis* at meter level 197 in the lower part of Chron C9n and with an astronomical age of 27.82 Ma is therefore in agreement with the position of the historically defined Rupelian/Chattian boundary.

The Criterion for the Base of the Chattian and the Sequence of Events across the Rupelian–Chattian Transition at MCA Section

The best criterion for defining the GSSP of the base of the Chattian Stage is the HCO of *C. cubensis*. In the MCA section, this biostrati-

graphic event lies at meter level 197, in the upper part of the calcareous nannofossil NP24 Zone, in the lower part of dinocyst Dbi Zone, and in the lower part of Chron C9n, i.e., about 14% distance up from the base of this 14-m thick Chron. Meter level 197 falls 30 cm below the base of a prominent, 40-cm-thick limestone bed in the middle of a recessive, 7-m-thick marly interval (Fig. 4b). This distinctive lithostratigraphic package is easily recognizable in the MCA section, and useful for recognizing the other lithostratigraphic intervals in this section, as well as in other coeval sections for regional correlation.

The location of this “golden spike” is fully consistent with the original position of the historical stratotypes (Fig. 10b). Moreover, because the HCO of *C. cubensis* is placed within a lithologic horizon easily recognized in the field, it also serves to locate the “golden spike” precisely in the horizon where a specific, arguably widely correlatable, biostratigraphic or magnetostratigraphic event occurs.

Our high-resolution and detailed bio-, magneto-, and chemostratigraphic study of the proposed Chattian GSSP at the MCA section, which accurately records calcareous plankton and dinoflagellate cyst biozones, magnetic reversals, seawater Sr, and C and O isotopic variations along with its astrochronologic calibration allows accurate and prompt correlation with the most recent mammalian zonation, biostratigraphy and events of the Oligocene according to the current state of knowledge provided by Vandenberghe et al. (2012).

Furthermore, the results discussed above enable recognition of the sequence of events across the Rupelian–Chattian transition at the MCA section (Figs. 2, 3, and 5). These are, in stratigraphic order from oldest (1) to youngest (6):

- (1) the top of Chron C10n at meter level 190.2 that would allow correlation with marine and continental records;
- (2) the oxygen isotope excursion possibly related to the Oi2a glaciation event, which falls at meter level 193 in the middle part of Chron C9r;
- (3) the base of Chron C9n at meter level 195.1 that would serve as correlation with marine and continental records;
- (4) the HO of the calcareous nannofossil *Sphenolithus predistentus* at meter level 196;
- (5) the HCO of the calcareous nannofossil *S. distentus*, i.e., the base of Zone NP25, which occurs between meter levels 201 and 202 in the middle part of Chron C9n;
- (6) the abundance interval of the Arctic cold-water *S. cooksoniae* dinocyst from meter level 201 to meter level 208, which is consistently found within Chron C9n.

Summary

The MCA section in central Italy is continuously and spectacularly exposed, well accessible, unaffected by structural deformations, rich in well-preserved planktonic foraminifera, calcareous nannofossils, and dinoflagellate cysts, and contains several biotite-rich volcanoclastic layers, some of which provided valuable radiometric dates. It fulfills all the sixteen requisites demanded by the International Commission of Stratigraphy for holding the Chattian GSSP (Table 2). The International Union of Geological Sciences has ratified the level that contains the highest common occurrence (HCO) of the planktonic foraminifer *Chiloguembelina cubensis* as the Global Stratotype Section and Point

Table 2. Summary of evaluation of the Monte Cagnero section for holding the Chattian GSSP

Requisites to be fulfilled by a GSSP	The GSSP Chattian Stage at Monte Cagnero
Exposure over an adequate thickness of sediment	Yes – The mid Rupelian to upper Chattian succession is extensively exposed over 86 m of pelagic carbonate rhythmites.
Continuous sedimentation	Yes – No hiatuses are detected in the Rupelian–Chattian interval.
High sedimentation rate	Yes – The mean sedimentation rate across the Rupelian–Chattian interval is about 1.45 cm/kyr, a relatively high sedimentation rate for a deep marine, pelagic sediment.
Absence of synsedimentary and tectonic disturbances	Yes – The Rupelian–Chattian transition is absolutely devoid of turbidites, slumps, or faults.
Absence of metamorphism and strong diagenetic alteration	Yes – Other than diagenetic cementation and recent superficial weathering, no metamorphism nor strong alteration have affected these pelagic carbonates.
Abundance and diversity of well preserved fossils	Yes – The carbonate pelagites bear abundant and well diversified microfossils. Up to date, planktonic foraminifera, calcareous nannofossils, dinoflagellate cysts have been studied in detail. Ostracods, smaller benthic foraminifera and ichnofossils will be the potential subjects for paleontological studies in the future.
Absence of vertical facies changes	Yes – No facies changes occur.
Favorable facies for long-range biostratigraphical correlations	Yes – The whole section is made up of marine, middle bathyal, fossiliferous pelagic carbonates deposited in a western Tethyan epeiric basin at a subtropical paleolatitude of about 35°N. Correlations with Boreal and Tethyan pelagic records are therefore possible. The combination of magnetostratigraphy, chemostratigraphy and cyclostratigraphy allows correlation with shallow-marine and terrestrial deposits.
Amenability to radiometric dating	Yes – The section contains several thin biotite-rich volcanoclastic layers. Two of these, which bracket the Rupelian–Chattian interval, were dated on biotite with the ⁴⁰ Ar/ ³⁹ Ar radioisotopic technique.
Amenability to magnetostratigraphy	Yes – The section yielded a complete magnetostratigraphic succession of polarity zones from the middle-upper Chron C12r to the lowermost Chron C8r.
Amenability to chemostratigraphy	Yes – Calcium carbonate and oxygen, carbon and strontium isotopes were analyzed through the section.
Amenability to cyclostratigraphy and astronomical tuning	Yes – The whole section has been tuned to Laskar 2004 long and short eccentricity via FFT spectral analysis using magnetic susceptibility as a proxy from closely spaced (5 cm) bulk rock samples.
Accessibility	Yes – The section is accessible thanks to an easy mountain trail, which can be navigated for less than a kilometer from an off road parking.
Free access	Yes – Free access to the section is guaranteed.
Permanent protection of the site	Yes – The area around the section is attributed to agriculture and wilderness, with no planning for any kind of urban development. Local institutions (Urbania town, Province of Pesaro-Urbino, Marche Region, University of Urbino, the Geological Observatory of Coldigioco and Osservatorio Piante terra) are willing to cooperate.
Possibility to fix a permanent marker	Yes – A metal plate will be fixed in the calcareous marly layer at meter level 197. Local institutions are willing to cooperate.

(GSSP) for the Chattian Stage, the standard chronostratigraphic unit of the upper Oligocene. This robust stratigraphic marker for the Rupelian/Chattian boundary occurs at meter level 197 of the MCA section, at the base of the planktonic foraminiferal O5 Zone and in the upper part of calcareous nannofossil NP24 Zone, in the lower part of dinocyst Dbi Zone, and in the lower Chron C9n. The location of this “golden spike” is consistent with the original position of the historical stratotypes. An astrochronologic calibration of the MCA section provides an astronomical age of 27.82 Ma for the Chattian GSSP.

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