

Late Cenozoic climate history of the Ross Embayment from the AND-1B drill hole: Culmination of three decades of Antarctic margin drilling.

T. R. Naish^{1,2}, R. D. Powell³, P.J. Barrett¹, R.H. Levy⁴, S. Henrys¹, G.S. Wilson⁵, L. A. Krissek⁶, F. Niessen⁷, M. Pompilio⁸, J. Ross⁹, R. Scherer³, F. Talarico¹⁰, A. Pyne¹ and the ANDRILL-MIS Science team¹¹

¹Antarctic Research Centre, Victoria University of Wellington, Wellington, New Zealand (t.naish@gns.cri.nz, peter.barrett@vuw.ac.nz, alex.pyne@vuw.ac.nz)

²Geological and Nuclear Sciences, Lower Hutt, New Zealand (t.naish@gns.cri.nz, s.henrys@gns.cri.nz)

³Department of Geology and Environmental Geosciences, Northern Illinois University, DeKalb, IL, USA (ross@geol.niu.edu)

⁴ANDRILL Science Management Office, University of Nebraska-Lincoln, 126 Bessey Hall, Lincoln, NE 68588-0341, United States (rlevy2@unl.edu)

⁵Department of Geology, University of Otago, PO Box 56, Dunedin, New Zealand (gary.wilson@otago.ac.nz)

⁶Department of Geosciences, The Ohio State University, Columbus, OH, USA (krissek@mps.ohio-state.edu)

⁷Department of Marine Geophysics, Alfred Wegener Institute, Postfach 12 01 61, Columbusstrasse, D-27515, Bremerhaven, Germany (fnriessen@awi-bremerhaven.de)

⁸Istituto Nazionale di Geofisica e Vulcanologia, Via della Faggiola, 32,

⁹I-56126 Pisa, Italy (pompilio@pi.ingv.it)
New Mexico Geochronology Research Laboratory, Socorro, NM 87801 (jirhiker@nmt.edu)

¹⁰Università di Siena, Dipartimento di Scienze delle Terra, Via Laterina 8, I-53100 Siena, Italy (talarico@unisi.it)

¹¹<http://www.andrill.org/support/references/appendixc.html>

Abstract Because of the paucity of exposed rock the direct physical record of Antarctic Cenozoic glacial history has become known only recently and then largely from off-shore shelf basins through seismic surveys and drilling. The number of holes has been small and largely confined to three areas (McMurdo Sound, Prydz Bay and Antarctic Peninsula), but even in McMurdo Sound, where Oligocene and early Miocene strata are well-cored, the Late Cenozoic is poorly known and dated. The latest Antarctic geological drilling program, ANDRILL, successfully cored a 1285m-long record of climate history spanning the last 13 m.y. from sub-sea floor sediment beneath the McMurdo Ice Shelf (MIS), using drilling systems specially developed for operating through ice shelves. The cores provide the most complete Antarctic record to date of ice sheet and climate fluctuations for this period of Earth's history. The >60 cycles of advance and retreat of the grounded ice margin preserved in the AND-1B record the evolution of the Antarctic ice sheet since a profound global cooling step in deep sea oxygen isotope records ~14 m.y. ago. A feature of particular interest is a ~90m-thick interval of diatomite deposited during the warm Pliocene, and representing an extended period (~200,000 years) of locally open water, high phytoplankton productivity and retreat of the glaciers on land.

Historical overview

The remarkable Late Cenozoic record of glacial history in the Ross Embayment recovered in late 2006 by the ANDRILL-MIS Project is the culmination of work begun over three decades ago to document and understand the more recent glacial history of Antarctica by drilling close to the margin. Ironically, although the last 3 million years of Earth's climate is often said to be the best studied interval of the Cenozoic, the contribution of Antarctic ice volume changes is the most poorly understood. Until this most recent hole was drilled the middle Cenozoic record was better known from several holes in both the McMurdo region and Prydz Bay (Table 1). This is because in general the older strata were exposed closer to the coast through basin uplift, where younger strata had been removed by Neogene erosion, but also because in offshore basins glacial debris from the last glacial advance prevented penetration and recovery of the older record from gravity coring (Anderson, 1999). Plainly deep drilling technology was required from an appropriate location. Here we summarise the history and rationale behind this successful outcome.

By the early 1970's the International Geophysical Year (1956-58) had already resulted in a vast increase in knowledge of rock types, ages and history of the continent itself (Bushnell and Craddock, 1970), but the history of its ice sheet was still a mystery. This changed in early 1973 with the voyages of the *Glomar Challenger* to the South Indian Ocean and Ross Sea (Leg 28; Hayes, Frakes et al., 1975) and the Tasman Sea and Southwest Pacific Ocean (Leg 29; Kennett, Houtz et al., 1974). The first of these cruises showed that the Antarctic ice sheet, which had previously been seen as a Quaternary feature, had been in existence since at least the Oligocene. The second yielded the first oxygen isotope measurements on deep-sea calcareous microfossils, providing the first evidence of dramatic ocean cooling and global ice volume increase at the Eocene-Oligocene boundary (Shackleton and Kennett, 1974).

The 1970's was also a period in which detailed paleontological and chronological studies of cores from continuously deposited deep-sea sediments displaced the long held view of four major Quaternary glaciations (Flint, 1971), showing much more frequent cycles of climate, ice volume and sea level in the late Quaternary (Hays et al. 1976) in response to variations in the Earth's orbital parameters calculated by Milankovitch. How far back in time these should be evident was not clear, but there seemed no reason why they should not have influenced the Antarctic ice sheet from the time of its inception.

While deep ocean sediments were useful for their continuous record of past ocean chemistry, providing an ice volume-temperature signal, they could not provide information on the extent of ice or regional climate in the

high latitudes. This could only come out of sediment cores from the Antarctic margin, where the direct influence of ice advance and retreat (and perhaps also sea level fall and rise) could be obtained.

In the early 1970's two drilling platforms were available – the *Glomar Challenger* operated by the Deep Sea Drilling Project and a land-based system put together for the Dry Valley Drilling Project, an initiative by the United States, Japan and New Zealand to explore the late Cenozoic history of the McMurdo Dry Valleys (Smith, 1981). Over the next three decades both of these platforms (and their successors) ran in parallel (Table 1, Figure 1) with varying degrees of success.

The ship-based system had the advantage that it could be deployed at a number of places around the Antarctic margin, but was limited by ice conditions and poor recovery of glacial sediments from near shore shelf basins. Nevertheless, the cores taken have provided useful constraints on the Cenozoic history of the Antarctic ice margin and climate from the Ross Sea, Prydz Bay and Antarctic Peninsula sectors (Table 1). The land-based system, once adapted for sea-ice, had the advantage that it could yield long and continuous core with near-complete recovery. However, it was rather cumbersome and required an ice platform that was firmly tied to land. Although the system has been improved through the MSSTS-1, CIROS and Cape Roberts Projects (Table 1), continuing to exploit the fast-ice rim around McMurdo Sound as a drilling platform, attractive drilling targets beneath fast sea ice elsewhere on the Antarctic margin have yet to be identified.

The results from both ship-based and sea-ice-based drilling have provided a framework for the Cenozoic history of the Antarctic ice sheet (Kennett and Warnke, 1992, 1993; Barrett, 1996), but more detail has become known for Oligocene and early Miocene times, especially from the 1500 m of strata cored off Cape Roberts (Barrett, 2007; Naish et al., 2001), than in subsequent times. The first results of the shift from sea ice- to shelf ice-based drilling reported below add a great deal to the late Cenozoic story.

The ANDRILL McMurdo Ice Shelf Project

The aim of the MIS Project was to obtain a continuous sediment core through approximately 1,200 meters (m) of Neogene (~0-10 Ma) glacial and volcanic sediment that had accumulated in the Windless Bight region of the McMurdo Ice Shelf (Fig. 2A). The present-day MIS forms the northwest part of Ross Ice Shelf where it has been pinned by Ross Island for the last ~10 ka (McKay et al., 2007), and is nourished by ice sourced locally and from the East Antarctic Ice Sheet (EAIS) outlet glaciers in the southern Transantarctic Mountains (TAM). The drillsite was situated above a flexural moat basin formed in response to Quaternary

volcanic loading of the crust by Ross Island, superimposed on regional subsidence associated with Neogene extension of the Terror Rift (Horgan et al., 2005; Naish et al., 2006; Fig. 2B).

Between 29th October and the 26th of December, 2006, a single 1284.87 m-deep, drill core (AND-1B) was recovered from the bathymetric and depocentral axis of the moat in 943 m of water from an ice shelf platform. The drilling technology employed a sea-riser system in a similar fashion to the Cape Roberts Project (CRP), but utilised a combination of soft sediment coring in upper soft sediments and continuous wire line diamond-bit coring. Innovative new technology, in the form of a hot-water drill and over-reamer, was used to make an access hole through 85 m of ice and to keep the riser free during drilling operations.

The MIS project has two key scientific objectives.

1. Provide new knowledge on the Late Neogene behaviour and variability of the Ross Ice Shelf/Ice Sheet (RIS) and the West Antarctic Ice Sheet (WAIS), and their influence on global climate, sea-level and ocean circulation.
2. Provide new knowledge on the Neogene tectonic evolution of the West Antarctic Rift System (WAR), Transantarctic Mountains (TAM) and associated volcanism.

A key outcome of the project is to provide age control for, and determine the environmental significance of, seismic reflectors that have been mapped regionally within the Victoria Land Basin (VLB) (Henrys et al., 2007; Fielding et al., 2007) in order to assess the regional impact of global climatic and local tectonic events. A second key outcome of the project will be to use paleoclimatic proxies and boundary conditions to help constrain numerical climate and dynamical ice sheet models. This paper presents an overview of the MIS Project, the details are reported Initial Results volume (Naish et al., 2007).

Tectonic and stratigraphic setting

Ross Island lies at the southern end of the Victoria Land Basin (VLB), a ~ 350km-long, half-graben, hinged on its western side at the Transantarctic Mountain (TAM) front (Fig. 2). Major rifting in the VLB has occurred since the latest Eocene, perhaps having been initiated in the Cretaceous (Cooper and Davey, 1985; Brancolini et al., 1995), and has accommodated up to 10 km of sediment. A new rift history, based on the Cape Roberts Project drill cores linked to a new regional seismic stratigraphic framework (Fielding et al., 2007; Henrys et al., 2007) indicates that crustal stretching during the Oligocene syn-rift phase produced rapid subsidence, followed by thermally-controlled slower subsidence in the Early Miocene. Renewed rifting within the centre of the VLB beginning in the Late Miocene has continued through to present day. This forms the Terror Rift (Cooper et al.,

1987), and is associated with alkalic igneous intrusions and extrusive volcanism (e.g. Beaufort Island and Ross Island). Quaternary loading of the crust by the Ross Island volcanoes has added significantly to subsidence near Ross Island, and the development of an enclosing moat (Stern et al., 1991). The Terror Rift, has accommodated up to 3 km of Neogene sediment beneath Windless Bight. Here, the load-induced subsidence caused by Ross Island has contributed significantly to the generation of accommodation space, especially during the last 2 m.y. (Horgan et al., 2005).

Neogene strata have now been extensively mapped in southern McMurdo Sound from the Drygalski Ice Tongue south to Ross Island. These Neogene strata show a thickening and eastward-dipping succession extending under Ross Island in the vicinity of the MIS Project drill site. Analysis of these strata which have now been sampled by MIS project drilling will contribute significantly to the young tectonic history of the West Antarctic Rift System.

Chronostratigraphy of AND-1B

A preliminary age model for the upper 700m of drill core constructed from diatom biostratigraphy (Scherer et al., 2007), and radiometric ages on volcanic material (Ross et al., 2007), allows a unique correlation of ~70% of the magnetic polarity stratigraphy with the Geomagnetic Polarity Time Scale (GPTS; Wilson et al., 2007). The age model provides several well-constrained intervals displaying relatively rapid (<1m/k.y) and continuous accumulation of sediment punctuated by several 0.5 to 1 m.y. stratal hiatuses representing more than half of the last 7 Ma. Thus the AND-1B record provides several highly-resolved “windows” into the development of the Antarctic ice sheets during the Late Cenozoic. Strata below ~620mbsf are Late Miocene in age (5-13 m.y.). At the time of writing, the chronostratigraphic data available for this interval include three radiometric ages on volcanic clasts from near 1280mbsf constraining the age for the base of the AND-1B drill core to <13.5 m.y. Work continues to improve the age control on the lower part of the cored interval.

Relationship to regional seismic stratigraphy

Prior to drilling the MIS target, five distinctive reflectors marking regional stratal discontinuities had been mapped through a grid of seismic data in the vicinity of the drillsite (HPP and MIS lines, Naish et al., 2006; Fig. 2), and linked to marine seismic reflection data and reflector nomenclature in McMurdo Sound (Fielding et al., 2007). We have used whole-core velocity measurements and VSP first arrival travel-time picks (see Hansraj et al., 2007; Naish et al., 2007) to derive a time-depth conversion curve to convert the seismic reflection section to depth. We have summarised in Figure 3, and below, our correlation of the regional seismic stratigraphy with the AND-1B drill core.

1. Rg (Surface C, bilious green reflector): This regionally extensive discontinuity is correlated with top of a ~60m-thick interval Late Miocene volcanic sandstone (LSU 7) and the base of a 150m-thick, high-velocity (3000ms^{-1}) interval of diamictite cycles (LSU 6.4). $^{40}\text{Ar}/^{39}\text{Ar}$ dates on ashes beneath Rg indicate that this discontinuity is <13.8 Ma.
2. Rh (Surface B, dark green): This regionally extensive discontinuity is correlated with the base of a ~180m-thick interval of Late Miocene-Early Pliocene, pyrite-cemented, high-velocity volcanic sandstone and mudstone (LSU 5). Regionally the green reflector correlates with the base of volcanic bodies in the Victoria Land Basin north of Ross Island. It is also correlated with the base of White Island volcano (Fig. 2) dated at ~ 7.6 Ma (Alan Cooper, U. Otago, unpublished data).
3. Ri (Surface A2, b-clino, red reflector): This regionally extensive reflector marks the base of a ~100m-thick seismically-opaque interval, that separates high-amplitude reflections of the underlying unit. It corresponds to the base of prograding clinoforms north of Ross Island, and locally marks the base of flexure associated with Ross Island volcanic loading (Horgan et al., 2004). In AND-1B, Surface A2 correlates with the boundary between the ~90 m-thick, low density, low-velocity (1700ms^{-1}), Early Pliocene diatomite interval (LSU 4.1), and the higher-velocity (< 2500ms^{-1}) diamictites of LSU 4.2 beneath. Diatom assemblages indicate the age of this surface lies between 5.0 Ma and 4.0 Ma. Regionally this reflector has been traced into western VLB, where it is correlated biostratigraphically in MSSTS-1 to core at about 20 mbsf that yields a Pliocene age of 4.6-4.0 Ma based on diatom microfossils (Harwood et al., 2006).
4. Rj (Surface A1, Turquoise reflector): This regionally extensive reflector marks the base of a ~150m-thick unit of strongly alternating high- and low-amplitude reflections. These dramatic cycles in density and velocity reflect regular alternations between diatomite and diamictite in Late Pliocene (LSU 3). The turquoise reflector separates strata above that are younger than ~3.0 Ma and below greater than ~3.5 Ma.

Stratigraphic architecture

The 1285m-long AND-1B drill core provides the first high-resolution, Late Neogene record from the Antarctic margin (Fig. 4) as well as the first long geological record from under a major ice shelf. Details of the lithostratigraphic subdivision, facies analysis and sequence stratigraphy are presented in an Initial Results volume (Krissek et al., 2007).

Glacial-interglacial cyclostratigraphy

At the time of writing, 60 unconformity-bounded glacial marine sedimentary cycles, of probable

Milankovitch-duration have been identified, representing repeated advances and retreats of an ice sheet across the drill site during the Late Neogene. Bounding unconformities, *glacial surfaces of erosion* (GSEs), are typically sharp and planar and mark dislocations between enclosing facies. Facies immediately beneath display a range of intraformational deformation features, including physical mixing of lithologies, clastic intrusions, faulting and soft-sediment deformation. Facies above these surfaces are typically diamictites and conglomerates, and are interpreted as subglacial tillites or near grounding-line glacimarine deposits. In many cycles, the facies succession reflects retreat of the grounding line through ice shelf into open ocean environments at the interglacial minimum, followed by ice readvance characterized by progressively more glacially-influenced facies in the upper parts culminating in a glacial surface of erosion (GSE) at the glacial maximum.

Major chronostratigraphic intervals and their cyclic character

Here we summarise the sedimentary succession cored in AND-1B by subdividing it into nine chronostratigraphic intervals largely on the basis of characteristic facies cycles, with their glacial and climatic implications (Fig. 4). The latter are addressed in more detail by Powell et al., (2007). Three of the intervals lack cyclicity, interval 1 (a volcanic sandstone), interval 4 (largely volcanic ash and interval 6 (the ~90m-thick diatomite).

- (1) Late Miocene volcanic sandstone (1275.24-1220.15 mbsf), LSU 7.
- (2) Late Miocene diamictite-dominated sedimentary cycles (1220.15-1069.2 mbsf), LSU 6.4
- (3) Late Miocene, diamictite/mudstone and sandstone sedimentary cycles (1069.2-759.32 mbsf), LSU 6.1-6.3.
- (4) Late Miocene-Early Pliocene? lapilli tuff, lava flow, and volcanic sandstone and mudstone (759.32-586.45 mbsf), LSU 5.
- (5) Early Pliocene diamictite/diatomite sedimentary cycles (586.45-459.24 mbsf), LSU 4.2-4.4.
- (6) Early Pliocene diatomite (459.24-382.98mbsf), LSU 4.1
- (7) Late Pliocene diamictite/diatomite sedimentary cycles (459.25-146.79 mbsf), LSU 3.
- (8) Late Pliocene-Early Pleistocene diamictite/volcanic mudstone and sandstone cycles (146.79-82.72 mbsf), LSU 2.
- (9) Middle-Late Pleistocene, diamictite-dominated sedimentary cycles (82.72-0mbsf), LSU 1

Late-Miocene diamictite-dominated cycles of LSU 6.4 (1220.15-1069.2 mbsf).

This interval is dominated by thick massive diamictite with thin stratified mudstone interbeds, representing alternations between basal glacial and

glacimarine deposition. The diamictites which form over 90% of the interval contain medium to high-grade metamorphic basement clasts known from the Byrd Glacier region (Talarico et al., this volume). At this stage we have been conservative in our recognition of grounding-line features in the core that might be used to identify boundaries of the sedimentary cycles. More detailed study is likely to find features such as sets of alternating sharp-based massive and stratified diamictites, representing subglacial and pro-grounding-line environments. The dominance of subglacial deposition and the occurrence of granitoid and metamorphic rocks from the Byrd Glacier region imply the long-term existence of a grounded Ross Ice Sheet with short ice shelf phases, and ice streaming from the southern Transantarctic Mountains. Once a magnetostratigraphy is developed for this section of the core together with a new $^{40}\text{Ar}^{39}\text{Ar}$ ages, further sedimentological and petrographic work should provide an important history of Antarctic Ice Sheet behaviour during the “big” Late Miocene, Mi-glaciations of Miller et al. (1991).

Late Miocene, diamictite/mudstone-sandstone sedimentary cycles of LSU 6.1-6.3(1069.2-759.32 mbsf).

This section is characterized by cycles of subglacial and grounding-line diamictites that pass upwards into a glacimarine retreat succession of redeposited conglomerate, sandstone and mudstone. These units are overlain by a more distal hemipelagic terrigenous mudstone with out-sized clasts and limestones. The retreat facies succession is then followed by a proglacial advance facies assemblage in which clast abundance increases together with the occurrence of submarine outwash facies, which is truncated by a glacial surface of erosion. In contrast to the underlying diamictite-dominated interval, over 60% of this unit comprises strata representing periods of both of ice-shelf cover and open ocean, though still with significant ice-rafting, over the drillsite. These observations imply a warmer climate for this whole period, and the association increased subglacial melt-water.

Late Miocene-Early Pliocene volcanic sandstone and mudstone of LSU 5 (759.32-586.45mbsf).

This interval is dominated by subaqueously-redeposited volcanic sediments, many with near-primary volcanic characteristics. The volcanoclastic sediments are organized into sediment gravity flow deposits (mostly proximal turbidites) indicating a nearby active volcanic centre, delivering primary volcanic material into a deep basin (several hundred meters water depth). A series of fining upwards, altered and degraded pumice lapilli tuffs occur at 623, 603 and 590 mbsf, and a pure fine volcanic glass sand occurs at 577 mbsf. All of these have been targeted for argon geochronology. A plagioclase Hawaiiite lava with chilled margins and “baked sediments” under its lower contact interrupts the volcanic sediment gravity flows between 646.49 - 649.30 mbsf, and has been dated

by $^{40}\text{Ar}/^{39}\text{Ar}$ at 6.38 Ma. The source of this lava and the redeposited volcanic units is thought to be nearby because of the freshness of the glass and angular nature of clasts in the breccias. The unit is virtually devoid of out-sized clasts or other features indicative of iceberg rafting and/or grounding-line proximity, suggesting that much of the 120 m-thick interval was deposited rapidly in open-water. At this stage we consider this volcanic interval to represent a single eruptive sequence with very high rates of accumulation.

Early Pliocene diamictite/diatomite sedimentary cycles of LSU 4.2-4.4 (586.45-459.24 mbsf).

These cycles can be summarized as comprising, in ascending stratigraphic order: (1) a sharp-based massive diamictite with variable amounts of volcanic glass (subglacial to grounding zone) passing up into (2) stratified diamictite, sandstone and mudstone with dispersed clasts (grounding-line to distal glacimarine) followed by (3) stratified diatomite (open ocean) which in turn commonly passes up into glacimarine mudstone and sandstone with dispersed clasts (perhaps an indication of an approaching grounding line). The interval that includes the lower part of the diamictite and the underlying glacimarine lithologies is often physically intermixed and displays deformation associated with glacial overriding and shearing, but interestingly, appears to represent little significant erosion by the advance. These cycles show a dramatic change from hemipelagic open water sedimentation to open water biogenic sediments— diatomite at interglacial minima.

Early Pliocene diatomite of LSU 4.1 (459.24-382.98 mbsf).

This interval is almost solely diatomite, indicating an extended period of high-productivity open water over the site. It is of Early Pliocene age (~4.2 Ma), and is likely to span a period of about 4.2-4.0 m.y., involving a number of glacial-interglacial cycles. The onset of an inferred warmer period is captured in the 10m transition beginning at 460mbsf from a clast-rich to clast-poor muddy diamictite through terrigenous mud to diatomite, with a decline in mud component, out-sized clasts and frequency of cm- to dm-thick gravity flow deposits. Further research is planned in order to quantify the contemporaneous climate, and will involve examination of the terrestrial microfossils and estimation of sea-surface temperatures from geochemical and biological proxies. This diatomite appears to coincide the widely recognised Early Pliocene global warming and accompanying higher sea levels (Dowsett et al., 1999; Ravelo et al., 2004), for which there is evidence elsewhere on the Antarctic margin (e.g. Whitehead et al., 2005). However, well-dated deposits and geomorphological evidence from the adjacent McMurdo Dry Valleys indicate a polar climate persisted there throughout this period (Marchant et al., 1996).

Middle to late Pliocene diamictite/diatomite sedimentary cycles of LSU 3 (382.98-146.79 mbsf).

Within this interval the core is strongly cyclic in nature, and is characterized by 12 glacial-interglacial cycles that typically comprise a sharp-based lower interval of diamictite in the upper few meters that passes upwards into a ~5-10m-thick unit of biosiliceous ooze and/or biosiliceous-bearing mudstone (diatomite). The lower part of diamictites and the upper few meters of diatomites are sheared and deformed, presumably from grounded ice. The transitions from diamictite (glacial and glacimarine) to diatomite (open ocean) are dramatic, in many cases occurring in less than a meter of core. Powell et al. (2007) discuss these abrupt facies transitions in terms of rapid retreat of an ice sheet or collapse of an ice shelf. Our preliminary age model suggests that the cycles in this interval may have formed in response to orbital forcing of the ice sheet. This interval will be the focus of further work as we evaluate its significance in terms of the global climatic deterioration coincident with the development of large ice sheets on the northern hemisphere continents (e.g. Shackleton et al., 1984; Maslin et al., 1999).

Late Pliocene-Early Pleistocene diamictite/volcanic mudstone-sandstone cycles of LSU2 (146.79-82.72 mbsf)

This interval displays lithologic alternations between ice-proximal facies and open-water volcanoclastic facies. A GSE at 150.90 mbsf is correlated with the Rk-pink reflector and spans as much as 0.7 m.y between 2.4 and 1.7 Ma. Above this, the interval from 146.79-134.48 mbsf represents another glacial-interglacial cycle, that above the diamictite is almost entirely made up of redeposited, volcanic (basaltic) sandstone. The sandstone is organized into a series of normally-graded turbidites. The interval from 134.48-120 mbsf is characterized by an open water assemblage of interstratified mudstone and volcanic sandstone that lies stratigraphically above, weakly stratified diamictite alternating with sparsely fossiliferous claystone and mudstone, more typical of grounding line and iceberg zone systems. A muddy-sandy volcanic breccia with near-primary volcanic material at 117 mbsf has yielded a preliminary age $^{40}\text{Ar}/^{39}\text{Ar}$ age of ~1.6 Ma on basaltic volcanic glass.

A dramatic change upwards into massive diamictite from 109.7-97.2 mbsf indicates the return of the ice sheet to the proximity of the drill site. Our preliminary age model suggests that up to ~0.4 m.y., between 1.5-1.1 Ma is missing between GSE at 109.7 mbsf and at least 2 GSEs in the overlying interval of amalgamated diamictite above. From 97.2-92.5 mbsf diamictite and volcanic sandstone pass downwards into interstratified bioturbated, volcanic sandstone and mudstone. The interval from 92.5-86.6mbsf contains the youngest biosiliceous-bearing sediments in the AND-1B core. The presence of numerous volcanoclastic units and biosiliceous sediments

in this interval indicate an extended period of open water conditions with no sea ice beyond the calving line. An immediately overlying, fining upwards lapilli tuff between 85.86-85.27 mbsf has yielded a high-precision $^{40}\text{Ar}/^{39}\text{Ar}$ age on sanidine phenocrysts of 1.01 Ma within a short normal polarity interval interpreted as the Jaramillo Subchron. Thus, the underlying biosiliceous interglacial sediments are tentatively correlated with warming during the “super-interglacial” associated with Marine Isotope Stage 31. This interval will be the focus of a future concentrated effort to better characterize the impact of this warm period on ice sheet behaviour.

Middle-Late Pleistocene, diamictite-dominated sedimentary cycles of LSU 1 (82-0 mbsf).

Between 83 and 27 mbsf there are at least 8 cycles of diamictite and thinner bioturbated and stratified intervals of mudstone, volcanic sandstone and muddy conglomerate. Like the Late Miocene diamictite-dominated cycles, the high proportion of subglacial and grounding-line proximal deposits, together with the occurrence of granitoid and metamorphic rocks known from the Byrd Glacier region (Talarico et al., 2007) implies existence of a grounded Ross Ice Sheet in the Middle and Late Pleistocene. The implied presence of a grounded ice sheet for much of this time is intriguing as it also corresponds to a period of Earth History dominated by 80-120ka fluctuations in large northern hemisphere ice sheets. Further work will focus on: (1) How this ice sheet responded to the Late Pleistocene interglacials, which in Antarctic ice core records indicated polar temperatures warmer than today (e.g. Petit et al., 1998; EPICA Community, 2004), and (2) the role of orbital forcing on Antarctic Ice Sheets (e.g. Huybers and Wunsch, 2005; Raymo et al., 2006) in regulating Pleistocene climate.

Implications for Antarctic glacial and climate history

The AND-1B core has the potential to contribute significant new knowledge on the dynamics of the WAIS and Ross Ice Shelf/Sheet system, as well as contributing to understanding the behaviour of the EAIS outlet glaciers during the Late Cenozoic. New chronological data will allow certain intervals of the core to be correlated with other proxy climate records (e.g. ice and marine isotope records) and hence the sensitivity of the ice sheet to a range of past global climate changes to be evaluated. The significant results thus far are that the Ross Ice Sheet, fed by WAIS and EAIS outlet glaciers, has undergone significant cyclic variations in extent and timing during the Late Neogene. A relatively colder and more stable ice sheet dominated the Ross Embayment in the early Late Miocene between 13-10 Ma, becoming more dynamic in the latest Miocene(9-7 Ma) with subglacial water discharge, still with periodic grounded ice in the Ross Embayment. These conditions were followed by a period around 4 Ma when the Ross Embayment was relatively ice-free, with highly productive, warmer oceanic conditions, followed by a

return to cycles of advance and retreat of grounded ice. From Middle Pleistocene to Recent the ice sheet is characterised by a change back to more stable colder conditions. Our preliminary analysis of the more than 25 Pliocene sedimentary cycles indicates significant glacial-interglacial variability, with regular oscillations between subglacial/ice proximal and open ocean ice distal environments, including extended periods of interglacial warmth when the ice was not calving into the ocean. Our environmental reconstructions, to date, imply changes in ice sheet volume that must have contributed significantly to eustasy (e.g. 10-20m).

Cold Early Miocene and Pleistocene till-dominated intervals with clasts originating in the TAM to the south imply that grounded ice from the big outlet glaciers to the south was reaching McMurdo Sound during these times. Today Byrd Glacier-sourced ice in the shelf flows east of Ross Island. Glaciological reconstructions [e.g. Denton and Hughes, 2002] require significant ice volume from WAIS to direct the flow lines of the southern outlet glaciers into the McMurdo region during glacial periods, and also to maintain an ice shelf during ensuing interglacial retreats. Thus, we view the sedimentary cycles representing primarily the expansion and contraction of WAIS in concert with fluctuations in the flow of TAM outlet glaciers. The TAM outlet glaciers alone do not provide enough ice volume to maintain an ice sheet or ice shelf in the Ross Embayment.

Work is planned in the near future to further understand and constrain ice sheet dynamics, especially during periods of hypothesised global warmth that will provide useful analogues for constraining the behaviour of Antarctic ice sheets in the context of future climate change. To achieve this, we will integrate our core data, such as clast provenance, biological and geochemical proxies, with the new generation of ice sheet and climate models.

Summary

Repetitive vertical successions of facies imply at least 60 fluctuations, of probable Milankovitch-duration, between subglacial, ice-proximal and ice-distal open marine environments. These have been grouped into 3 main facies associations that correspond to glacial-interglacial variability during climatically-distinct periods of the Late Neogene:

- i) Cold polar climate with MIS site dominated by grounded ice, but some retreat to ice shelf conditions (Late Miocene, ~13-10 Ma and Pleistocene, ~1-0 Ma),
- ii) Warmer climate with MIS site dominated by ice-shelf and open water conditions (hemipelagites), with occasional periods of grounded ice. (early-Late Miocene, ~9-6 Ma).
- iii) Warmer climate with extended periods experiencing open ocean conditions (pelagic diatomites) with periods of sub-ice shelf and grounded ice deposition (Pliocene, ~5-2 Ma).

The ~90m-thick Early Pliocene (~4.2 Ma) interval of diatomite shows no apparent glacial cyclicity and represents an extended period of ice-free conditions indicative of a reduced WAIS. Late Pliocene (~2.6-2.2 Ma) glacial-interglacial cycles characterised by abrupt alternations between subglacial/ice-proximal facies and open marine diatomites imply significant WAIS volume fluctuations around the time of the early Northern Hemisphere glaciations. A ~4m-thick interval of diatomaceous mudstone in the Middle Pleistocene also represents similar open-ocean interglacial conditions. The last million years is dominated by deposition from grounded ice with periods of sub-iceshelf sedimentation like the present day.

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Figure 1. Locations of geological drill sites on land and on the Antarctic continental shelf. Details and references are given in Table 1. McMurdo Sound drill sites are shown in Fig. 2A.

Figure 2. [A] Location of key geographical, geological and tectonic features in southern McMurdo Sound. Volcanic centres of the Erebus Volcanic Province include Mt Erebus (E), Mt Terror (T), Mt Bird (B), White Island (W), Black Island (B), Mt Discovery (D), Mt Morning (M) and Minna Bluff (MB). Boundary faults of the southern extension of Terror Rift are also shown. Location of ANDRILL Program drill sites and previous programs (DVDP, CIROS, MSST) are shown. [B] Schematic structural-stratigraphic cross-section across the VLB (located in [A] as A-A') shows the stratigraphic context of the MIS and SMS drill sites with respect to previous drilling in Southern McMurdo Sound. The cross-section is compiled from interpreted seismic reflection data, previous drill core data from MSSTS-1 and CIROS-1 (Barrett 1986; 1989) and models for the evolution of the VLB.

Figure 3. Integrated plot correlating seismic reflectors with the lithologic log. We have used the whole-core velocity (c) to derive a time-depth conversion curve (b) together with VSP arrival times to map the seismic reflection profile from MIS-1 (a) to drill hole depth and correlate with core lithologies and lithostratigraphic units (e). Seismic stratigraphic units identified are from the MIS Science Logistics and Implementation Plan (SLIP; Naish et al., 2006), and have been mapped regionally (Fielding et al., 2007).

Figure 4. Lithostratigraphy, chronostratigraphy and sequence stratigraphy of the AND-1B drillcore. Cyclic variations in lithologies reflect periodic fluctuations of the ice margin in western Ross Embayment during the last 13 Ma. Notes: *After Ross et al. (2007). ^ After Wilson et al. (2007).

Table 1. Summary of stratigraphic drill holes on the Antarctic continent and shelf from 1973 to 2006.

Project	Year	Site	Lat	Long	Elev/Depth (+ / -)	Depth cored	% rec	Oldest core	Reference			
<i>Ross Sea</i>												
DSDP 28	1973	270	77°26'S	178°30'W	-634 m	423 m	62%	gneiss - E Paleozoic	Hays, Frakes et al., 1975			
		271	77°26'S	178°30'W	-562 m	233 m	7%	diamict clasts - E Pliocene				
		272	77°26'S	178°30'W	-629 m	439 m	37%	diamict - E Miocene				
		273	77°26'S	178°30'W	-491 m	333 m	25%	diamict - E Miocene				
<i>McMurdo Sound area - onshore</i>												
DVDP	1973	1	77°50'S	166°40'E	67 m	201 m	98%	basalt - L Quat	Kyle, 1981			
		2	77°51'S	166°40'E	47 m	179 m	96%	basalt - L Quat				
		3	77°51'S	166°40'E	48 m	381 m	90%	basalt - L Quat				
		10	77°35'S	163°31'E	3 m	182 m	83%	diamict - L Miocene				
		11	77°35'S	163°25'E	80.2 m	328 m	94%	diamict - L Miocene				
		12	77°38'S	162°51'E	75.1 m	185 m	98%	migmatite - E Paleozoic				
<i>McMurdo Sound area - offshore</i>												
DVDP	1975	15	77°26'S	164°23'E	-122 m	62 m	52%	black sand - E Pleist	Barrett & Treves, 1981			
MSSTS	1979	1	77°34'S	163°23'E	-195 m	230 m	62%	mudstone - L Oligocene	Barrett, 1986			
CIROS	1986	1	77°05'S	164°30'E	-197 m	702 m	98%	boulder congl - L Eocene	Barrett, 1989			
CIROS	1984	2	77°41'S	163°32'E	-211 m	168 m	67%	gneiss - E Paleozoic	Barrett & Hambrey, 1992			
		CRP	1997	1	77°00'S	163°45'E	-154 m	148 m		86%	diamict - E Miocene	CRST, 1998
		2		77°00'S	163°43'E	-178 m	624 m	95%		mudstone - Oligocene	CRST, 1999	
CRP	1998	3	77°00'S	163°43'E	-295 m	939 m	97%	sandstone - Devonian	CRST, 2000			
		ANDRILL	2006	1B	77°55'S	167°01'E	-840 m	1285 m	98%	basalt - E Miocene	Naish et al., 2007	
<i>Prydz Bay</i>												
ODP 119	1988	739	67°17'S	75°05'E	-412 m	487 m	34%	diamict - L Eo-E Oligocene	Barron, Larsen et al., 1988			
		740	68°41'S	76°43'E	-808 m	226 m	32%	red beds - ?Triassic				
		741	68°23'S	76°23'E	-551 m	128 m	26%	sandst, siltst - ?E Cretaceous				
		742	67°33'S	75°24'E	-416 m	316 m	53%	mudst, diamict - ?Eo-Olig				
		743	66°55'S	74°42'E	-989 m	97 m	22%	diamict - Pleistocene				
ODP 188	2000	1166	67°42'S	74°47'E	-475 m	381 m	19%	claystone - L Cretaceous	O'Brien, Cooper, Richter et al., 2001			
<i>Antarctic Peninsula</i>												
ODP 178	1998	1097	66°24'S	70°45'W	-563 m	437 m	14%	diamict - E Pliocene	Barker, Camerlenghi, Acton et al., 1999			
		1098	64°52'S	64°12'W	-1010 m	47 m	99%	mud - Holocene				
		1099	64°57'S	64°19'W	-1400 m	108 m	102%	mud - Holocene				
		1100	66°53'S	65°42'W	-459 m	111 m	5%	diamict - Pleistocene				
		1102	66°48'S	65°51'W	-431 m	15 m	6%	diamict - Pleistocene?				
		1103	64°00'S	65°28'W	-494 m	363 m	12%	diamict - L Miocene				







