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## The ShakeMap Atlas of historical earthquakes in Italy: configuration and validation --Manuscript Draft--

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<b>Corresponding Author:</b>	ILARIA OLIVETI, Ph.D Istituto Nazionale di Geofisica e Vulcanologia Roma, Italy ITALY
<b>Corresponding Author Secondary Information:</b>	
<b>Corresponding Author's Institution:</b>	Istituto Nazionale di Geofisica e Vulcanologia
<b>Corresponding Author's Secondary Institution:</b>	
<b>First Author:</b>	ILARIA OLIVETI, Ph.D
<b>First Author Secondary Information:</b>	
<b>Order of Authors:</b>	ILARIA OLIVETI, Ph.D Licia Faenza Andrea Antonucci Mario Locati Andrea Rovida Alberto Michellini
<b>Order of Authors Secondary Information:</b>	
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No changes have been made, except for the following corrections requested by the reviewer#3:

Line 35: "made since long" -> "made"

Line 55: "it has been .. Atlas" -> "version 4 of the Atlas has been released"

Line 180: "reveals to be" -> "is"

Line 302: "To this regard" -> "In this regard"

Line 383: "adjourned" -> "recalibrated"

1 The ShakeMap Atlas of historical earthquakes in Italy:  
2 configuration and validation

3 Ilaria Oliveti<sup>1\*</sup>, Licia Faenza<sup>2</sup>, Andrea Antonucci<sup>3</sup>, Mario Locati<sup>3</sup>, Andrea  
4 Rovida<sup>3</sup>, and Alberto Michelini<sup>1</sup>

5 <sup>1</sup>Istituto Nazionale di Geofisica e Vulcanologia, Sezione ONT, Rome, Italy.

6 <sup>2</sup>Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Bologna,  
7 Bologna, Italy.

8 <sup>3</sup>Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Milano, Milan,  
9 Italy.

10 \*Corresponding author. Email: [ilaria.oliveti@ingv.it](mailto:ilaria.oliveti@ingv.it)

11 **DECLARATION OF COMPETING INTERESTS**

12 The authors acknowledge there are no conflicts of interest recorded.

13 **ABSTRACT**

14 Italy has a long tradition of studies on the seismic history of the country and the neighbour-  
15 ing areas. Several archives and databases dealing with historical earthquake data – primarily  
16 intensity data points – have been published and are constantly updated. Macroseismic fields  
17 of significant events are of foremost importance in assessing earthquake effects and for the

18 evaluation of seismic hazard. Here we adopt the USGS-ShakeMap software to calculate the  
19 maps of strong ground shaking (shakemaps) of 79 historical earthquakes with magnitude  $\geq$   
20 6 that have occurred in Italy between 1117 and 1968 CE. We use the macroseismic data  
21 published in the Italian Macroseismic Database (DBMI15). The shakemaps have been de-  
22 termined using two different configurations. The first adopts the virtual intensity prediction  
23 equations approach (VIPE; i.e., a combination of ground motion models, GMMs, and ground  
24 motion intensity conversion equations, GMICEs; Bindi et al. (2011a); Oliveti et al. (2022b)).  
25 The second exploits the intensity prediction equations (IPE, Pasolini et al. (2008a); Lolli  
26 et al. (2019)). The VIPE configuration has been found to provide more accurate results  
27 after appraisal through a cross-validation analysis and has been applied for the generation of  
28 the ShakeMap Atlas. The resulting maps are published on the INGV ShakeMap (see Data  
29 and Resources; Oliveti et al., 2023), and on the ASMI (see Data and Resources; Rovida  
30 et al., 2017) platforms.

## 31 INTRODUCTION

32 Macroseismic intensity observations of past earthquakes can provide valuable constraints for  
33 reconstructing shaking distributions in the absence of instrumentally recorded data and are  
34 commonly used to estimate the location and magnitude of historical events (e.g., Teramo  
35 et al., 1996; Bakun and Wentworth, 1997; Gasperini et al., 2010; Beauval et al., 2010; Provost  
36 and Scotti, 2020, amongst others). Much effort has been made to aggregate the available  
37 data in comprehensive historical earthquake catalogs at both the national (e.g. Fäh et al.,  
38 2011; Manchuel et al., 2018; Rovida et al., 2020), and international scales (e.g., the European  
39 Preinstrumental Earthquake Catalogue EPICA; Rovida and Antonucci, 2021; Rovida et al.,  
40 2022a). These catalogs are fundamental for complementing and extending back in time  
41 instrumental earthquake catalogs for probabilistic seismic hazard assessment studies. In  
42 turn, it occurs that historical macroseismic intensities are the only long-term shaking data  
43 against which the outcomes of probabilistic seismic hazard studies can be tested, and sanity



44 checks be performed (Stirling and Petersen, 2006; Mucciarelli et al., 2008; Brooks et al.,  
45 2019).

46 However, depending on several historical, geographical and seismological factors, macro-  
47 seismic intensity distributions of past earthquakes as inferred from historical documentation  
48 may present temporal and spatial gaps. Several methods exist in the literature to reconstruct  
49 and/or integrate the spatial distribution of the ground shaking of historical events, based  
50 on different approaches and assumptions (for a review, see Antonucci et al., 2021, and refer-  
51 ences therein). Among these, the established ShakeMap methodology (Wald et al., 1999) has  
52 been used for defining the ground shaking of historical earthquakes at both the global (e.g.,  
53 Allen et al., 2008) and local scales (e.g., Schwarz et al., 2008; Faenza et al., 2013). In par-  
54 ticular, Allen et al. (2008) developed the so-called ShakeMap Atlas, a compilation of peak  
55 ground motions and intensity maps for  $\sim 14,100$  recent and historical earthquakes worldwide.  
56 More recently, Version 4 of the Atlas has been released (see Data and Resources), including  
57 a vastly expanded compilation of shakemaps for consequential and widely felt earthquakes  
58 using the updated ShakeMap (Version 4) software. For historical earthquakes in the global  
59 ShakeMap Atlas, macroseismic intensity values often represent the only available observa-  
60 tions, or provide valuable constraints, whereas strong-motion recordings are sparse (Allen  
61 et al., 2009b). The ShakeMap Atlas contributed to the development of fragility curves and  
62 loss model calibration (Luco and Karaca, 2007; García et al., 2012) and, to this end, pro-  
63 vides a fundamental resource for the USGS Prompt Assessment of Global Earthquakes for  
64 Response (PAGER) system (Earle et al., 2009; Allen et al., 2009a) and the Earthquake Con-  
65 sequences Database (Crowley et al., 2013) within the Global Earthquake Model (GEM)  
66 initiative.

67 From the technical point of view, the ShakeMap Atlas includes direct empirical equations  
68 that estimate site intensity for a given earthquake magnitude and distance (IPEs, intensity  
69 prediction equations) to incorporate macroseismic intensity as a native ground motion pa-  
70 rameter. According to the currently available peer-reviewed studies, IPEs strongly depend

71 on the selected data, and they differ in the approaches employed for the statistical analysis,  
72 such as the regression technique (e.g., Sørensen et al., 2009, amongst others) or the fully  
73 probabilistic method (e.g., Pasolini et al., 2008a). Due to the regional dependency of seis-  
74 mic characteristics, several studies developed regional or local intensity attenuation relations  
75 (e.g., Bakun, 2006; Bakun and Scotti, 2006; Stromeyer and Grünthal, 2009; Bindi et al.,  
76 2011b; Baumont et al., 2017; Oros et al., 2019, amongst others) for different regions of the  
77 world.

78 In the current study, we first produced shakemaps of strong ( $M \geq 6$ ) Italian histori-  
79 cal earthquakes using two different models implemented in the newly developed USGS-  
80 ShakeMap version 4 (Worden et al., 2020) by gathering intensity data from the Italian  
81 Macroseismic Database DBMI15 version 4.0 (see Data and Resources; Locati et al., 2022),  
82 and then selected the most appropriate configuration through the application of a ranking  
83 procedure consisting of statistical tests. In particular, we have considered for our configura-  
84 tion for the calculation of the intensity maps the Italian-derived IPE proposed by Pasolini  
85 et al. (2008a,b) recalibrated by Lolli et al. (2019), hereafter Pea08, and the default “vir-  
86 tual” IPE (Worden et al., 2017) available in ShakeMap, hereafter VIPE, as described in the  
87 ShakeMap Configuration section. The results obtained using the VIPE and Pea08 configu-  
88 rations have been appraised by analyzing the differences between the intensity predictions  
89 and the observed data using an iterative cross-validation procedure analysis (also known as  
90 leave-one-out analysis; Tomczak, 1998; Hofierka et al., 2007; Worden et al., 2010; Michelini  
91 et al., 2020). In addition, we have also investigated how the inclusion of finite faults affects  
92 the accuracy and the robustness of the ShakeMap prediction.

93 In summary, this work aims at presenting how we have developed the ShakeMap Atlas  
94 of historical earthquakes in Italy. To this goal, we have focused much attention to evaluate  
95 the accuracy of the selected configurations in order to provide a consistent and quantitative  
96 description of the distribution of shaking resulting from historical events in Italy.

## 97 DATA

98 Italy has a very long tradition of macroseismic investigation that produced a wealth of studies  
99 and data on the seismic history of the country and promoted the compilation of comprehen-  
100 sive historical macroseismic catalogues. The Italian Archive of Historical Earthquake Data  
101 ASMI (see Data and Resources, Rovida et al., 2017) collects more than 430 seismological  
102 studies, and grants access to a large number of intensity data from a variety of sources, such  
103 as macroseismic bulletins, online databases, and many scientific papers and reports. As a  
104 whole, it supplies data on more than 6500 Italian earthquakes in the period 461 B.C. to 2020  
105 CE. The current release of the Italian Macroscopic Database DBMI (DBMI15, see Data and  
106 Resources, here considered in its version 4.0; Locati et al., 2022) is obtained by selecting, for  
107 each earthquake, data that are collected in ASMI according to their content, reliability, and  
108 quality, and to the number and spatial distribution of intensity data. DBMI15 version 4.0  
109 contains 123981 Macroscopic Data Points (MDPs) related to 15343 populated places (from  
110 big towns to small villages) in Italy and 3229 earthquakes in the time-window 1000-2020  
111 CE. In order to provide a homogeneous set of intensity data, DBMI15 applies the following  
112 standardization procedures to the original input data: i) a consistent gazetteer related to the  
113 Italian territory was adopted in order to unambiguously match a pair of geographical coor-  
114 dinates of each locality with the intensity value provided by the original study, ii) a standard  
115 based on Arabic numerals (e.g., 6, 6-7, 7) was used to express the macroseismic intensity,  
116 and iii) a set of descriptive codes (e.g., “HF” for Highly Felt, “SD” for Slightly Damage,  
117 “D” for Damage, “HD” for Heavy Damage) was adopted when the original data source does  
118 not assess a proper numerical intensity value, e.g., because the available information is not  
119 sufficient.

120 To generate the shakemaps, we extracted from DBMI15 the MDPs related to earthquakes  
121 with magnitudes equal to or greater than 6 that occurred before 1972, for which no instrumen-  
122 tal ground motion recording exists according to the ITalian ACcelerometric Archive (ITACA;  
123 Russo et al., 2022). To this purpose, we improved the web service originally developed in the

124 framework of the EPOS Thematic Core Service for Seismology (Haslinger et al., 2022) for  
125 accessing macroseismic intensity data both in the European Archive of Historical Earthquake  
126 Data AHEAD (Locati et al., 2014; Rovida and Locati, 2015) and in the Italian Archive of  
127 Historical Earthquake Data ASMI (see Data and Resources). In particular, to the already  
128 supported XML, CSV and GeoJSON data encodings, we added an output format suitable  
129 for the ShakeMap software that wraps the required files into a zip package. The web-service  
130 recompiles the uncertain intensity values (e.g., 6-7 or 7-8) as half degrees (e.g., 6.5 or 7.5),  
131 according to the standard adopted in DBMI15 (see Rovida et al., 2020).

132 The dataset used for testing the ShakeMap configurations and generating the shaking  
133 maps includes 79 earthquakes that occurred between January 1117 and January 1968 with  
134  $6 \leq M \leq 7.3$  according to the Italian Parametric Earthquake Catalogue CPTI15 (Rovida  
135 et al., 2020, 2022b, Figure 1). The main characteristics of the selected events according to  
136 CPTI15 are reported in Table 1, in terms of origin time, epicentral location, magnitude,  
137 number of MDPs, name of the epicentral area and reference macroseismic study.

138 The intensities provided by DBMI15 for the considered earthquakes are assessed in the  
139 Mercalli-Cancani-Sieberg scale (MCS; Sieberg, 1923), and MDPs with descriptive intensity  
140 codes were not included in the dataset because they represents data for which the available  
141 information is not considered sufficient for assessing any intensity value. As a result, our  
142 dataset consists of 12632 MDPs in total, and the number of available data per earthquake is  
143 extremely variable (Table 1), with a minimum of 2 MDPs for the July 17, 1361 earthquake  
144 (M 6.3) and a maximum of 1366 MDPs for the February 23, 1887 earthquake (M 6.3). As  
145 shown in Figure 2, the number of MDPs of the entire dataset (and of the single events)  
146 increases through time. This increase results from the low intensity values that start to be  
147 represented significantly only after  $\sim 1850$  (see earthquake number 63 and subsequent ones  
148 in Figure 2). In contrast, the number of the highest intensities in the dataset is rather  
149 uniform through time (Figure 2). This is in agreement with the historical analysis of the  
150 time series of significant earthquake effects of Stucchi et al. (2004) who concluded that the

151 completeness for intensities higher than degree 8 might start as far back as the 12th century,  
152 depending on the area. In addition, it is noteworthy to point out that our dataset consists of  
153 strong earthquakes whose intensities are homogeneously assessed with a-posteriori analyses  
154 of earthquake records provided by archival documentation (see Guidoboni and Stucchi, 1993;  
155 Guidoboni and Ferrari, 2000; Camassi, 2004). This implies that the macroseismic assessment  
156 of historical earthquakes is conducted by professional historians according to the methods of  
157 historiographic research, i.e., taking into account the specific temporal, cultural, social, and  
158 geo-political contexts in which the records were produced (see Guidoboni and Ebel, 2009).

159 In Figure 3 we make an attempt to verify whether any prominent bias affects the dataset  
160 used to determine our ShakeMap Atlas. All the panels graph the MDPs as distance from  
161 the earthquake versus intensity. In the panels on the left hand side (lhs) (Figure 3a, c,  
162 e), the MDPs are shown according to three time slots (1000-1399, 1400-1799, and 1800-  
163 2020). On the right hand side (rhs) (Figure 3b, d, f), the MDPs are grouped according to  
164 three magnitude ranges (6.0-6.5, 6.5-7.0, and 7-7.5). The panels to left and right appear  
165 to confirm that no significant bias (i.e., larger number of higher intensities) exists for the  
166 older events. The same panels evidence also that the intensities lower than or equal to 6 are  
167 poorly represented in the dataset, due to a possible incompleteness of the far field data of  
168 strong historical earthquakes (Antonucci et al., 2023). The panels in which the MDPs are  
169 grouped by magnitude (rhs), show that as the magnitude increases there is the expected shift  
170 towards larger distances of the higher intensities. The temporal color scale reflects the MDPs  
171 distribution of the lhs panels. In general, this figure shows that about a quarter (24.02%) of  
172 MDPs has distances from the earthquake larger than 100 km. As far as short distances are  
173 concerned, MDPs with intensity value greater than 7 are prevalent.

## 174 **SHAKEMAP CONFIGURATION**

175 ShakeMap is an interpolation algorithm that makes use of recorded data and seismological  
176 and geotechnical knowledge to produce maps of ground motion at local and regional scales.

177 Thus, in addition to the observations, the prediction equations expressed in terms of peak  
178 ground motion parameters, the so-called ground-motion models (GMMs), and the intensity  
179 prediction equations (IPEs) are indispensable in ShakeMap to supplement the generally  
180 sparse and incomplete available data. In addition,  $V_{s30}$ , defined as the average seismic  
181 shear-wave velocity from the surface to a depth of 30 meters, is important for estimating  
182 local site amplifications of the ground motion. Specifically, ShakeMap accounts for the  
183 local site amplifications using an equally spaced grid of  $V_{s30}$  values. When site classes  
184 are the only available information (e.g., Eurocode 8 [EC8] soil categories), they need to  
185 be converted into the corresponding  $V_{s30}$  values (see Michelini et al., 2020, for Italy). A  
186 comprehensive explanation of how site effects are integrated into ShakeMap can be found  
187 in the detailed description provided by Worden et al. (2017). Moreover, GMICEs (ground  
188 motion to intensity conversion equations) are adopted wherever macroseismic intensities have  
189 to be transformed into ground motion parameters as, for example, when macroseismic data  
190 are used as input for generating ground motion maps and vice-versa (i.e., when estimating the  
191 macroseismic intensity field from recorded instrumental peak ground motion parameters).  
192 Therefore, the selection of the proper set of equations plays a key role in accurately estimating  
193 the shaking.

194 The application of an updated method to the interpolation process (Worden et al., 2018;  
195 Engler et al., 2022) to generate the shakemaps allows for more rigorous estimates of ground  
196 shaking and proper accounting of associated uncertainties when conditioned on geograph-  
197 ically distributed strong-motion station data or macroseismic intensity observations. Ac-  
198 cording to this method, the interpolation in ShakeMap is performed by treating the ground  
199 motions (or the intensities) as a conditional multivariate normal distribution (MVN). This  
200 approach, in combination with a GMM and cross-correlation functions among the available  
201 data, provides a flexible framework for estimating the ground shaking at arbitrary locations.  
202 For quantifying the uncertainty in these estimates, this technique also preserves the separa-  
203 tion of the conditioned residuals into between-event (perfectly correlated) and within-event

204 (spatially correlated) spatial processes (Engler et al., 2022).

205 In this work, to the purpose of identifying the most accurate ground shaking field of past  
206 historical earthquakes, we have computed the shakemap set (Worden et al., 2020), i.e., maps  
207 of macroseismic intensity, and five intensity measures — peak ground acceleration (PGA),  
208 peak ground velocity (PGV), and spectral acceleration (SA) ordinates at 0.3, 1.0, and 3.0  
209 s, respectively, using two different configurations. For what concerns the generation of the  
210 “macroseismic intensity” maps, the first configuration adopts the default VIPE, whereas the  
211 second one implements the IPE proposed by Pasolini et al. (2008a,b), updated by Lolli et al.  
212 (2019). We remark, however, that for both configurations we have generated the maps of  
213 PGA, PGV, and SA adopting (1) the GMMs selected by Michelini et al. (2020) accounting  
214 for the subdivision of Italy in different tectonic regimes and (2) the GMICEs of Oliveti et al.  
215 (2022b) calibrated on the dataset by Oliveti et al. (2022a) for the conversion between ground  
216 motion and macroseismic intensity.

217 With regard to the first point, Michelini et al. (2020) identified the most suitable GMMs  
218 to be utilized in each region based on the GMM zonation proposed by Visini et al. (2022) for  
219 the Italian seismic hazard model MPS19 (Meletti et al., 2021). In particular, Michelini et al.  
220 (2020) validated this configuration within ShakeMap evidencing a substantial improvement  
221 in the accuracy of ground-motion estimates for Italy. Since all earthquakes in our validation  
222 dataset fall within the shallow active crustal region (SACR), both the VIPE and Pea08  
223 configurations use the Bindi et al. (2011a) GMM that is used for the SACR tectonic regime  
224 and shallow depth earthquakes.

225 As for the second point, the reversible GMICEs proposed by Oliveti et al. (2022b) corre-  
226 late the maximum horizontal component of recorded PGA, PGV, and SA at  $T = 0.3, 1.0$  and  
227  $3.0$  s to macroseismic intensity values for Italy. Specifically, Oliveti et al. (2022b) adopted  
228 the common current approach involving a regression for the intensity as a function of the  
229 PGM parameters and viceversa, resulting in magnitude-distance-independent conversions,  
230 showing no significant trend of the residuals for both magnitude and distance. Very recently,

231 two new different methodologies have been introduced by Gallahue and Abrahamson (2023)  
232 to develop GMICEs. The authors state that the GMICEs developed using their approaches  
233 lead to more accurate estimates of the intensities than currently adopted methodologies.  
234 Here we note, however, that the comparison made by Oliveti et al. (2022b) with similar re-  
235 gressions previously published for Italy (e.g., Faenza and Michelini, 2010, 2011; Zanini et al.,  
236 2019; Masi et al., 2020; Cataldi et al., 2021, amongst others) demonstrates that the proposed  
237 relationships provide significantly improved fits to the data regardless. Moreover, to further  
238 validate their effectiveness, these relations were tested within the ShakeMap system of the  
239 Italian configuration, showing very accurate estimates of shaking and minimal bias.

240 In the present study, VIPE is a combination of selected GMM and associated GMICEs,  
241 which, combined together, offer the same interface and behavior of an IPE. This makes  
242 VIPE inherently valid for a broader range of regional and tectonic environments but it also  
243 entails increased uncertainty in the estimated intensity values compared to the currently  
244 available IPEs. Generally, VIPE is used in ShakeMap when the operator does not specify  
245 an IPE. The related module predicts the ground motion through the GMM and converts it  
246 to intensity using the GMICEs. In detail, it first attempts to use PGV for the calculation  
247 of the intensities, and then tries PGA, and then SA(1.0).

248 By contrast, Pea08 is a specific Italian macroseismic intensity attenuation model cali-  
249 brated as a function of moment magnitude and epicentral distance (Pasolini et al., 2008a). A  
250 recalibration of Pea08 was done by Lolli et al. (2019) using the updated intensity data points  
251 from DBMI15 (Locati et al., 2022) and earthquake parameters provided by CPTI15 (Rovida  
252 et al., 2022b). As a result, the macroseismic intensity attenuation model proposed by Lolli  
253 et al. (2019) has the same functional form of the equation of Pasolini et al. (2008a) but  
254 different values of parameters.

255 Since ShakeMap requires as input the hypocentral depth, and CPTI15 does not provide  
256 this information when the instrumental epicentre is not available, we arbitrarily assigned  
257 a value of 10 km to all the analyzed historical earthquakes (see Table 1). However, our



258 results are not affected by this choice because the ground motion model adopted is based on  
259 epicentral distance and the IPEs are not depth-dependent relationships either.

## 260 **COMPARISON BETWEEN IPE AND VIPE IMPLEMENTATION**

261 In this section, we present the results of the tests on the two selected configurations (i.e. using  
262 VIPE and Pea08, respectively) to show their accuracy in predicting the intensity value at the  
263 macroseismic data points. To this end, we adopted an iterative cross-validation procedure  
264 that performs the following steps for each observed intensity. Select a target earthquake and,  
265 iteratively, for each MDP:

- 266 • remove the MDP from the dataset;
- 267 • use the ShakeMap procedure to predict the intensity at the removed MDP (i.e., while  
268 keeping all the others);
- 269 • compute the difference between the observed and predicted intensity value at the re-  
270 moved MDP.

271 This procedure has been repeated for all the earthquakes selected. For the validation  
272 analysis, we computed the shakemaps using Pea08 and VIPE separately as input. For both  
273 configurations, we used the same values for the source parameters (e.g., hypocenter and  
274 magnitude), and the GMM, site effects and GMICEs mentioned above.

275 It is important to note that the intensity predictions were not derived for the total  
276 number of observed data extracted from DBMI15, i.e. 12632. First, we removed from  
277 further processing all the data points located outside the ShakeMap regular grid (i.e., a finely-  
278 sampled grid - nominally 1km spacing - of latitude and longitude pairs, whose dimensions  
279 depend on the earthquake magnitude). As a result, only 12299 MDPs were used initially  
280 for the cross-validation analysis. Then, ShakeMap's automatic removal of outliers (i.e.,  
281 observations that exceed two standard deviations above or below the prediction) reduced  
282 the number of data points ensuring the calculation of robust maps of ground shaking. In

283 our case, this quality assurance protocol found 115 and 377 outliers using VIPE and Pea08,  
284 respectively, showing a better performance in predicting intensity of the former over the  
285 latter. Finally, we also removed MDPs not common to both datasets obtained using the  
286 two models. This step is needed to ensure that the comparison occurs only among the same  
287 observed points. In summary, the entire data processing resulted in a final validation dataset  
288 with a total of 11885 MDPs.

289 The results of the cross-validation analysis for all the earthquakes are presented as dif-  
290 ferences between observed and predicted intensity values (i.e., residuals) through the violin  
291 plot representations in Figure 4.

292 Figure 4a shows that the median value for both models is close to zero, whereas the  
293 standard deviation calculated using VIPE is smaller than that obtained adopting Pea08.  
294 This indicates that both configurations do not suffer from significant systematic bias, but  
295 VIPE shows a smaller scatter in the residuals than Pea08.

296 When the data are grouped according to the EC8 site classes (Fig. 4b-d), we observe  
297 overall the same behavior described for the entire dataset (Fig. 4a). If we focus our attention  
298 to the disaggregated results in Figure 4c-d, we note that median values close to zero are found  
299 for the EC8 B–C soil site classes, which extend over a significant part of the Italian territory  
300 (CEN, 2004) and, consequently, over the great part of the selected localities. Conversely, in  
301 Figure 4b positive median values for the EC8 site class A indicate a slight underprediction  
302 of the level of intensity predicted by VIPE. This latter trend is likely due to VIPE using  
303 the configured GMM for predicting ground motion. In this regard, Michelini et al. (2020)  
304 explained it as due to inappropriate attribution of the EC8 soil site class A to stations  
305 effectively sited on softer and more amplifying soils. In contrast, Pea08 does not include  
306 site-amplification factors for implementing site effects, and the residuals do not suffer from  
307 the same underprediction.

308 Table 2 shows the mean, median, standard deviation, first and third quartiles of the  
309 distribution of the residuals for the entire dataset and for the EC8 A-C classes disaggregated

310 subsets. The standard deviations of the residuals vary between 0.15 and 0.22 for VIPE, and  
311 between 0.23 and 0.24 for Pea08. The first and third quartiles range between -0.11 and 0.06  
312 for VIPE, except for Q3 which equals 0.23 for EC8 class A, whereas the same parameters  
313 for Pea08 feature larger values, i.e. between -0.13 and 0.14.

314 While VIPE offers better predictions at the MDPs, it comes at the cost of higher uncer-  
315 tainty in these predicted intensity values than Pea08. In the case of Pea08, the uncertainty  
316 at the observation site is assumed to be zero, whereas, at the predicted points, it is assumed  
317 to have a non-zero uncertainty due to the spatially averaged nature of intensity assignments.  
318 More specifically, the uncertainty for estimates from Pea08 is the stated uncertainty given  
319 in Lolli et al. (2019) conditioned on geographically distributed macroseismic intensity ob-  
320 servations, as described in the ShakeMap Configuration section. By contrast, for VIPE,  
321 an additional uncertainty associated with the conversion itself (i.e., the uncertainty of the  
322 GMICES) results in the predictions. This is due to the three-step procedure adopted in  
323 ShakeMap when using VIPE, that first converts intensities to peak ground motions (PGMs)  
324 using the GMICES, then uses the GMM to supplement sparse data in its interpolation and  
325 estimation of ground motions, and finally converts the PGMs back to intensities using the  
326 GMICES. In particular, the standard deviation of the predicted intensity calculated using  
327 VIPE is given by the rules of error propagation (Ku et al., 1966). In practice, the uncer-  
328 tainty in the predicted intensity values is computed by combining the uncertainty of the  
329 GMM with the uncertainty of the GMICES. This is confirmed by the results illustrated in  
330 Figure 5. The histograms of Figure 5 show the distribution of the standard deviation of the  
331 predicted intensity values, for both VIPE and Pea08, as calculated by ShakeMap.

332 Additional tests were carried out to explore the behavior of the calculated residuals with  
333 distance from the earthquake (Fig. 6a) and intensity value (Fig. 6b), respectively. The  
334 residuals obtained using either VIPE or Pea08 are very close to zero when considering all  
335 the distances and the intensity classes, except for intensities lower than 4-5. This leads us  
336 to state that ShakeMap slightly overestimates the intensity values when compared to the

337 observed data. A possible explanation for this behavior comes from the magnitude range  
338 of the earthquakes in our dataset ( $M \geq 6.0$ ) that feature few low value intensities (less  
339 than 4-5), at long earthquake distances as shown in Figure 3. Other factors, however, more  
340 related to the calibration of Pea08 and the adopted GMM (Bindi et al., 2011a) can account  
341 for the low shaking and they cannot be excluded. Furthermore, we observed that, in all  
342 cases shown in Figure 6a-b, the residuals calculated using VIPE perform better than those  
343 obtained adopting Pea08, which show many outliers and much scattered data.

344 In order to verify if the accuracy of the intensity predictions improves when the finite-  
345 ness of the source dimensions are considered, the leave-one-out cross-validation analysis was  
346 applied to a subset of the original dataset using VIPE. We selected the faults and the focal  
347 mechanism parameters from the Database of Individual Seismogenic Sources (DISS, see Data  
348 and Resources; DISS Working Group, 2021) matching each selected earthquake with a fault  
349 whenever possible. As a result, we found that for only 16 earthquakes we could identify the  
350 appropriate fault (Table 3). We then used the leave-one-out cross-validation technique as  
351 above to test the goodness of the selected configuration considering a total of 4799 MDPs.  
352 The violin plot diagrams of Figure 7 show the distribution of the differences between the  
353 configuration with and without the fault geometry at all the intensity points. Figure 7 shows  
354 no significant improvement in the prediction performance when including the finite faults.  
355 The same comparison cannot be made for the configuration that adopts Pea08 because this  
356 IPE depends only on epicentral distance. In fact, one feature of ShakeMap is that it consid-  
357 ers the actual rupture plane (or its surface projection) rather than the epicenter, when the  
358 fault is included in the processing.

359 Overall, our results evidence the goodness of VIPE in predicting the intensity data within  
360 the ShakeMap algorithm. Since VIPE is computed by combining the GMM by Bindi et al.  
361 (2011a) and the GMICEs by Oliveti et al. (2022b), our tests are crucial to cross-verify the  
362 consistency of these relationships. This relevant observation confirms the ShakeMap accuracy  
363 in estimating the shaking when adopting proper GMMs and GMICEs.

364 The intensity maps shown in Figure 8 have been obtained with the investigated configu-  
365 rations (i.e. using VIPE and Pea08) and were drawn as examples from the entire ShakeMap  
366 Atlas of historical earthquakes in Italy (see Data and Resources). These include the 5 Decem-  
367 ber 1456 M 7.2, the 27 March 1638 M 7.1, and the 13 January 1915 M 7.1 earthquakes. They  
368 have been selected since they all resulted in significant fatalities and damage. More specif-  
369 ically, the December 1456 M 7.2 earthquake resulted in over 10,000 deaths (Meletti et al.,  
370 1988), whereas the other two events caused the deaths of nearly 30,000 people (Guidoboni  
371 et al., 2007; Molin et al., 1999, respectively). With regard to the comparison between the  
372 maps obtained using the two different configurations, we note that the selected earthquakes  
373 seem well suited to this end since they all have a large number of MDPs. In general, the  
374 VIPE configuration appears to generate slightly larger intensity values when compared to  
375 Pea08 at large earthquake distances.

## 376 CONCLUSIONS

377 In this work, we appraised two different USGS-ShakeMap configurations to compute the  
378 ShakeMap Atlas of large historical earthquakes in Italy using the available macroseismic  
379 data. To this end, we produced a shakemap set, in terms of macroseismic intensity, PGA,  
380 PGV, and SA at 0.3s, 1.0s, and 0.3s, for 79 earthquakes with magnitude  $\geq 6$  between 1117  
381 and 1968 CE.

382 We identified the most appropriate configuration between VIPE (i.e., the combination of  
383 the selected GMM and GMICEs) and the (direct) IPE proposed by Pasolini et al. (2008a)  
384 and recalibrated by Lolli et al. (2019), through the adoption of an iterative cross-validation  
385 procedure within ShakeMap. To convert from macroseismic intensities to peak ground mo-  
386 tion (and viceversa), we adopted the configuration of ShakeMap described by Michelini et al.  
387 (2020) and the GMICEs proposed by Oliveti et al. (2022b). To assess the accuracy of the  
388 results obtained using the two configurations, we used the leave-one-out cross-validation  
389 analysis applied to the macroseismic intensity points within ShakeMap. Our analysis of the

390 residuals (i.e., the differences between the observed and predicted intensity data) obtained  
391 with VIPE and Pea08 shows that, overall, the former predicts more accurately the intensity  
392 values for all the MDPs irrespective of distance and EC8 soil class type. The only exception  
393 is represented by the EC8 A class (hard rock) which shows some slight underestimation  
394 of the predicted intensities. The leave-one-out cross-validation analysis was also applied to  
395 estimate the intensity prediction capabilities when the finite fault is available to find that it  
396 does not improve significantly the accuracy of the intensity estimates.

397 In conclusion, the proposed configuration using VIPE appears to provide accurate macro-  
398 seismic intensity estimates for historical earthquakes in Italy. The resulting shakemaps are  
399 available on the INGV ShakeMap (see Data and Resources; Oliveti et al., 2023), and on the  
400 ASMI platforms (see Data and Resources; Rovida et al., 2017).

## 401 **DATA AND RESOURCES**

402 The earthquakes have been selected from the Parametric Catalogue of Italian Earthquakes  
403 CPTI15 Version 4.0 (<https://doi.org/10.13127/CPTI/CPTI15.4>) and a tabular list is pro-  
404 vided in Table 1. The intensity observations are all accessible on the Version 4.0 of the Italian  
405 Macroseismic Database DBMI15 through the ShakeMap webservices ([https://emidius.mi.  
406 ingv.it/services/macroseismic/](https://emidius.mi.ingv.it/services/macroseismic/)). The U.S. Geological Survey (USGS)-ShakeMap open-  
407 source software is available on the GitHub development platform ([https://github.com/  
408 usgs/shakemap](https://github.com/usgs/shakemap)). The shakemaps presented in this paper ([https://doi.org/10.13127/  
409 shakemaps/historical](https://doi.org/10.13127/shakemaps/historical)) are available at <http://shakemap.ingv.it/shake4/>, through the  
410 web portal of the Italian Archive of Historical Earthquake Data ASMI ([https://doi.  
411 org/10.13127/asmi](https://doi.org/10.13127/asmi)). The USGS ShakeMap Atlas is available at [https://earthquake.  
412 usgs.gov/data/shakemap/atlas/](https://earthquake.usgs.gov/data/shakemap/atlas/). Version 3.3.0 of the Database of Individual Seismo-  
413 genic Sources (DISS) is available at <https://doi.org/10.13127/diss3.3.0>. OpenQuake  
414 library of modern ground-shaking intensity models is available at [https://docs.  
415 openquake.org/oq-engine/3.13/\\_modules/openquake/hazardlib/gsim/](https://docs.openquake.org/oq-engine/3.13/_modules/openquake/hazardlib/gsim/). Some analy-

416 ses and plots are made using ObsPy (Beyreuther et al., 2010; Megies et al., 2011; Krischer  
417 et al., 2015) and the Python pandas software (<https://pandas.pydata.org>).

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693 Full physical mailing address for each author:

- 694 • Ilaria Oliveti: Istituto Nazionale di Geofisica e Vulcanologia, Sezione ONT, via di  
695 Vigna Murata 605, 00143, Rome, Italy.
- 696 • Licia Faenza: Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Bologna, viale  
697 Berti Pichat, 6/2, 40127, Bologna, Italy.
- 698 • Andrea Antonucci: Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Milano,  
699 via Alfonso Corti 12, 20133, Milan, Italy.
- 700 • Mario Locati: Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Milano, via  
701 Alfonso Corti 12, 20133, Milan, Italy.
- 702 • Andrea Rovida: Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Milano, via  
703 Alfonso Corti 12, 20133, Milan, Italy.
- 704 • Alberto Michelini: Istituto Nazionale di Geofisica e Vulcanologia, Sezione ONT, via di  
705 Vigna Murata 605, 00143, Rome, Italy.

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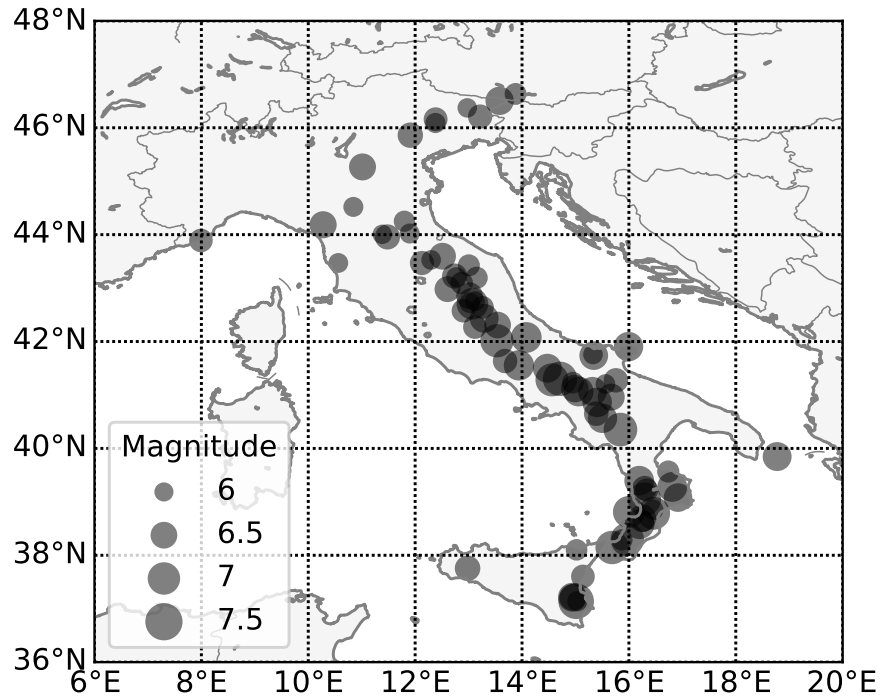


Figure 1: Spatial distribution of the selected seismic events (grey circles). Circle sizes are plotted relative to their magnitude value.

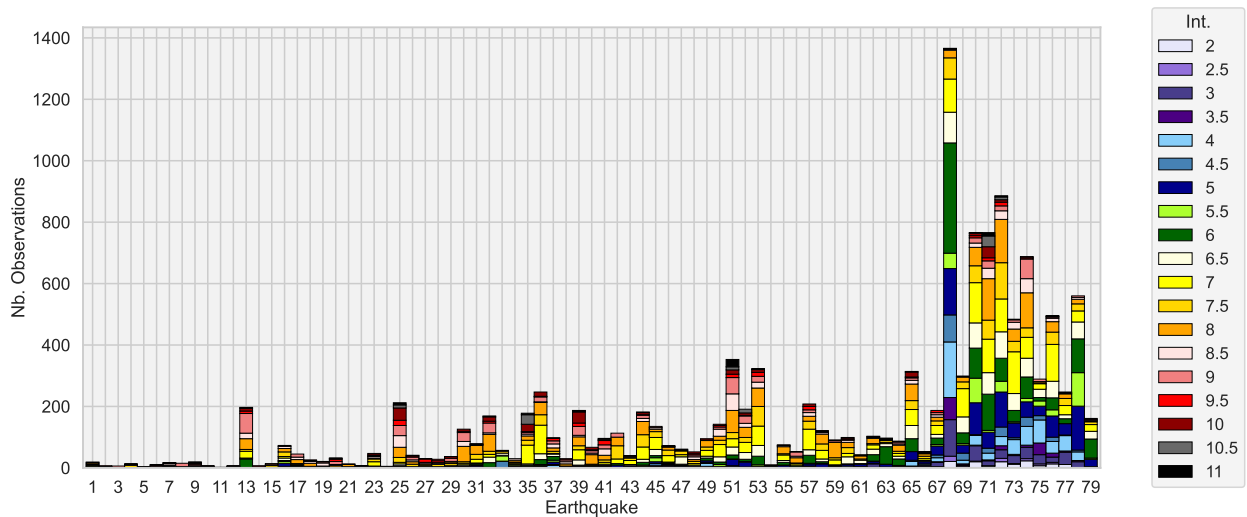


Figure 2: Number of MDPs extracted from DBMI15 per earthquake for different macroseismic intensity values. The dataset includes 79 earthquakes sorted in chronological order (from 1117 to 1968) following the numbering provided in Table 1

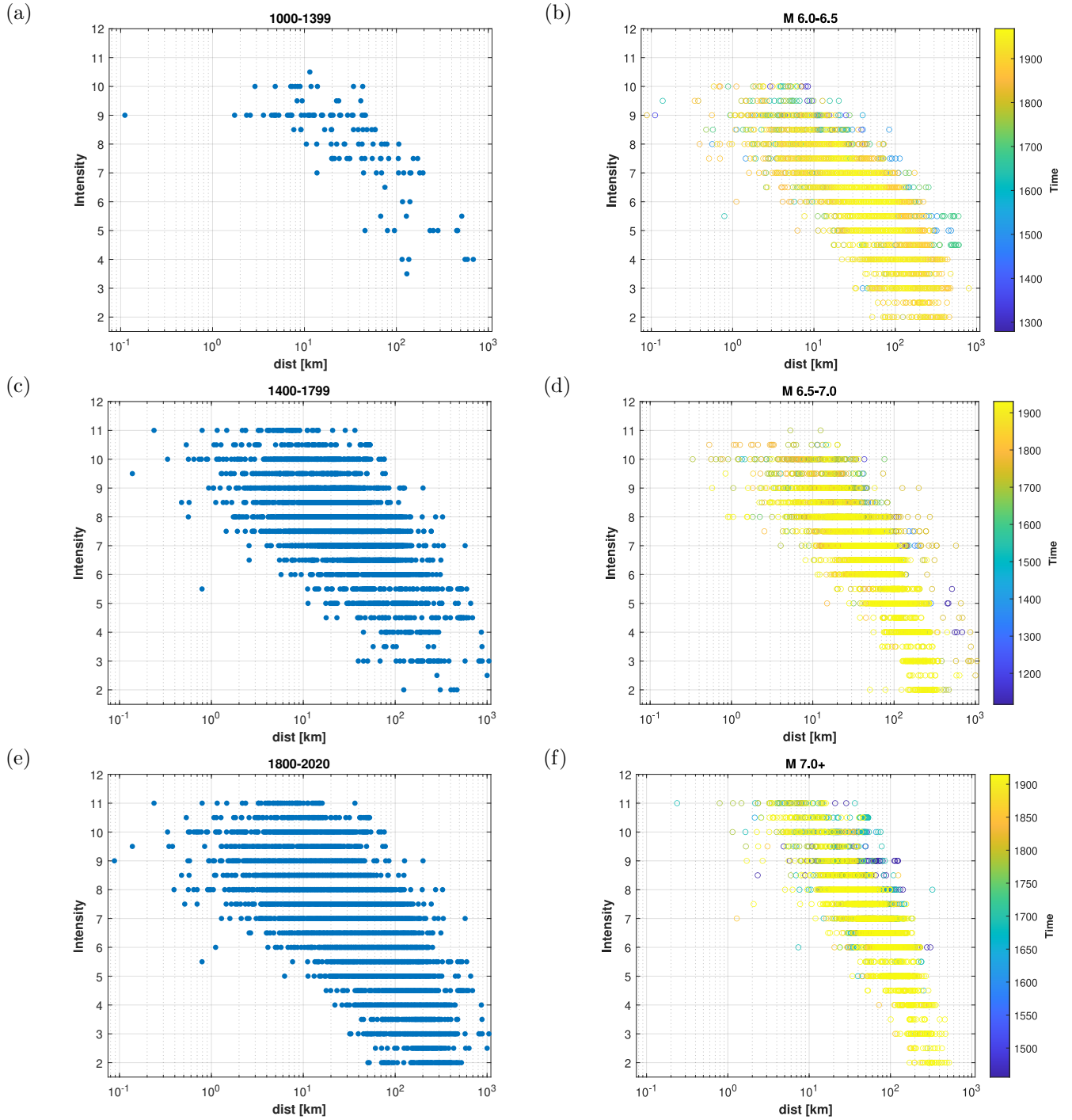


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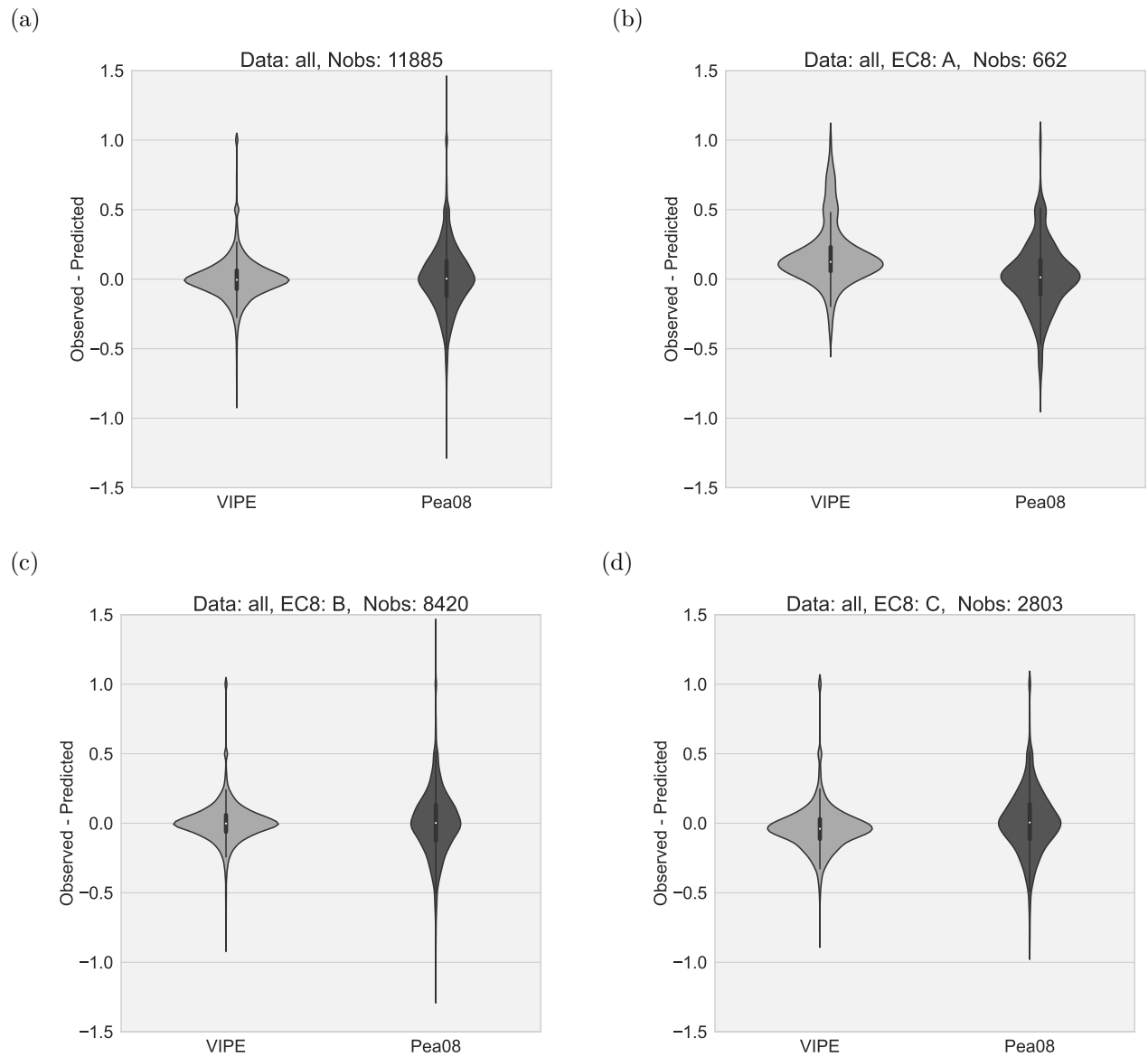


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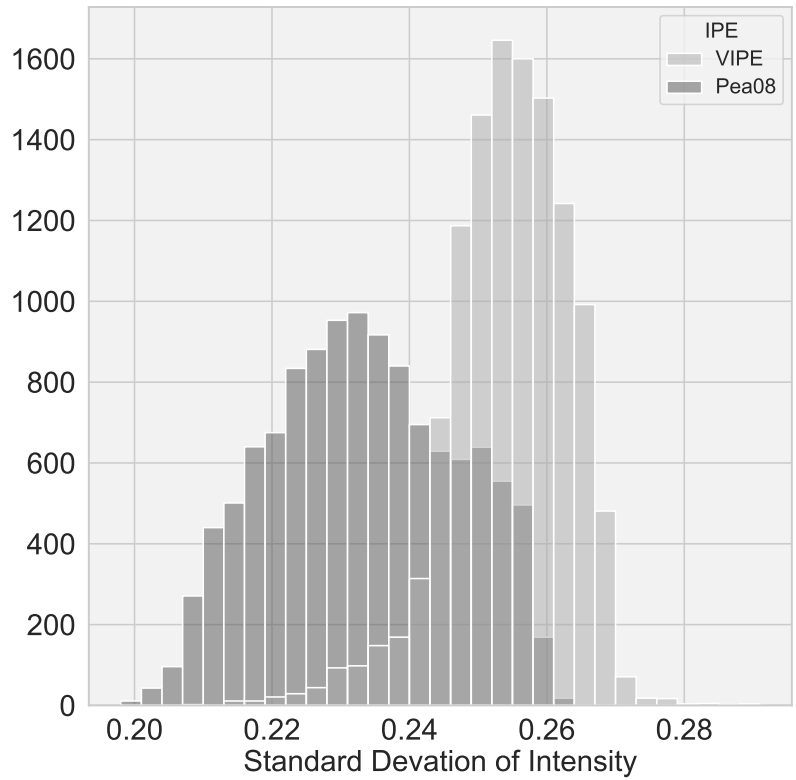
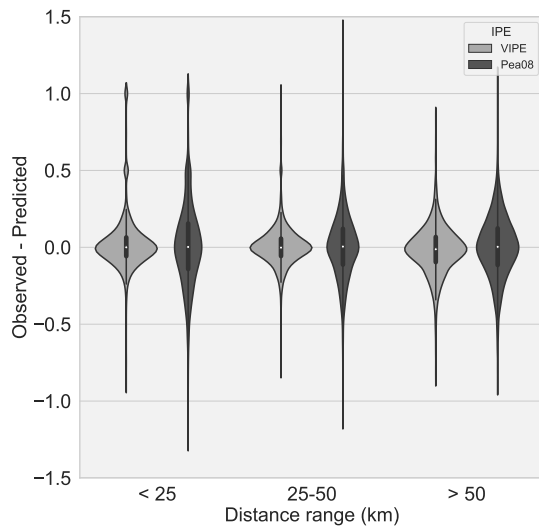


Figure 5: Histograms of the standard deviation distribution of the predicted intensity values at the 11885 macroseismic data points for the VIPE (light grey) and Pea08 (grey) configurations. The overall higher values of the VIPE distribution reflects the larger uncertainty in the predictions due to the additional uncertainty of the GMICEs.

(a)



(b)

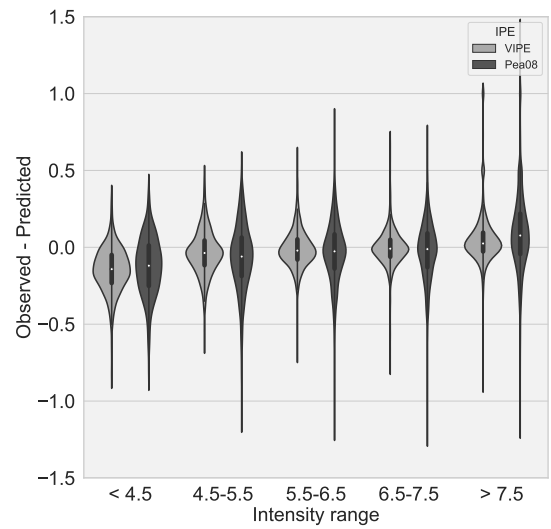


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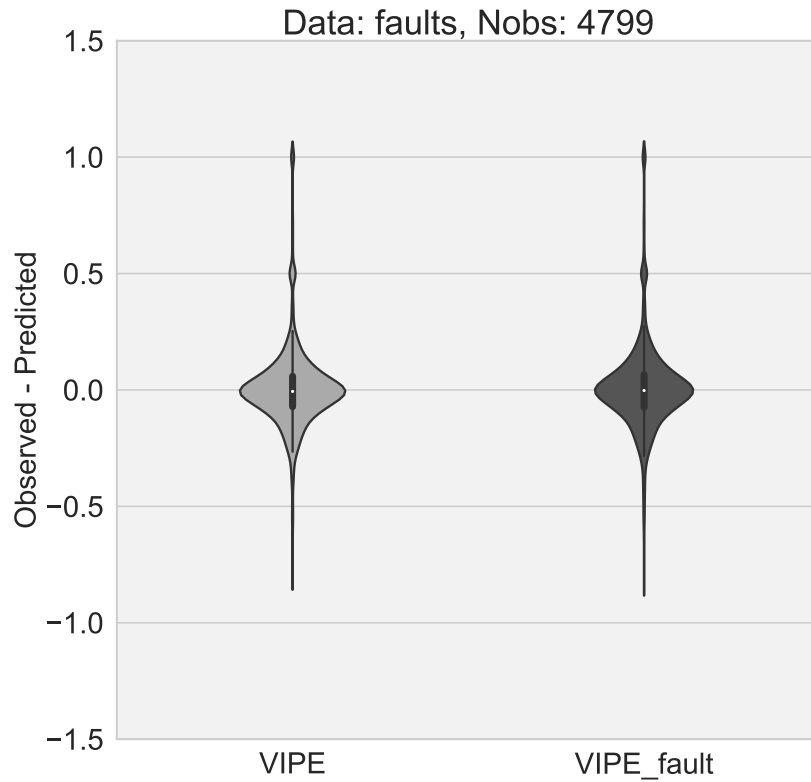


Figure 7: Violin plot diagram of the differences between observed and ShakeMap predicted intensity values for a subset of the validation dataset (16 of 79 earthquakes) resulting from the leave-one-out cross-validation test, using the VIPE configuration with and without the fault geometry. The faults and the focal mechanism parameters are provided by the DISS database (see Data and Resources; DISS Working Group, 2021).

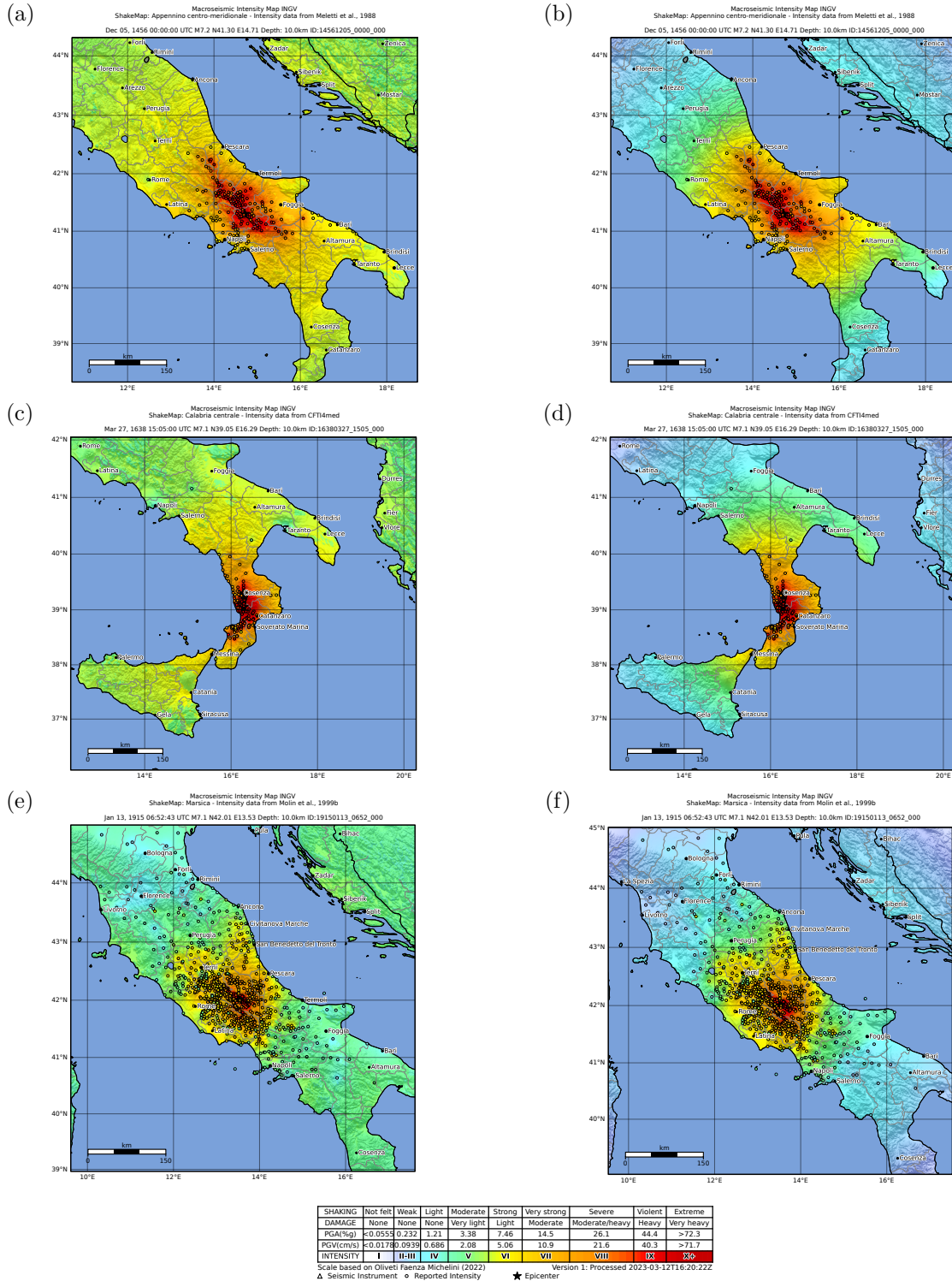


Figure 8: Intensity maps for the 5 December 1456 M 7.2, 27 March 1638 M 7.1, and 13 January 1915 M 7.1 earthquakes. The maps have been created using the v.4 of the U.S. Geological Survey (USGS)-ShakeMap software with the VIPE (a,c,e) and Pea08 (b,d,f) configurations.

Table 1: List of the selected seismic events: event number and ID, time, epicenter, magnitude, number of macroseismic data, name of the epicentral area and the reference macroseismic study for each event are indicated.

No.	Event-ID	Origin Time	Lat	Lon	Mag	MPDs	Epicentral Area	Reference
1	11170103.1515.000	1117-01-03T15:15:00Z	45.267	11.015	6.5	19	Veronese	Guidoboni et al. (2007)
2	11690204.0700.000	1169-02-04T07:00:00Z	37.215	14.949	6.5	7	Sicilia sud-orientale	Guidoboni et al. (2007)
3	11840524.0000.000	1184-05-24T00:00:00Z	39.395	16.193	6.8	6	Valle del Crati	Guidoboni et al. (2007)
4	12790430 1800.000	1279-04-30T18:00:00Z	43.093	12.872	6.2	13	Appennino umbro-marchigiano	Monachesi (1987)
5	12981201.0000.000	1298-12-01T00:00:00Z	42.575	12.902	6.3	4	Monti Reatini	Guidoboni et al. (2007)
6	13281201.0000.000	1328-12-01T00:00:00Z	42.857	13.018	6.5	11	Valnerina	Monachesi (1987)
7	13480125 1530.000	1348-01-25T00:00:00Z	46.504	13.581	6.6	17	Alpi Giulie	Caracciolo et al. (2015)
8	13490909.0000.000	1349-09-09T00:00:00Z	42.270	13.118	6.3	15	Appennino laziale-abruzzese	Guidoboni et al. (2007)
9	13490909 0815.001	1349-09-09T00:00:00Z	41.554	13.942	6.8	19	Lazio-Molise	Galli and Naso (2009)
10	13521225.0000.000	1352-12-25T00:00:00Z	43.469	12.127	6.3	7	Alta Valtiberina	Castelli et al. (1996)
11	13610717 1715.000	1361-07-17T17:15:00Z	41.205	15.561	6.0	2	Subappennino dauno	Guidoboni et al. (2007)
12	13891018.0000.000	1389-10-18T00:00:00Z	43.527	12.299	6.0	7	Alta Valtiberina	Castelli et al. (1996)
13	14561205.0000.000	1456-12-05T00:00:00Z	41.302	14.711	7.2	197	Appennino centro-meridionale	Meletti et al. (1988)
14	14611127 2105.000	1461-11-27T00:00:00Z	42.308	13.543	6.5	7	Aquilano	Tertulliani et al. (2009)
15	15010605 1000.000	1501-06-05T10:00:00Z	44.519	10.844	6.1	14	Modenese	Guidoboni et al. (2007)
16	15110326 1440.000	1511-03-26T15:30:00Z	46.209	13.216	6.3	73	Friuli-Slovenia	Camassi et al. (2011)
17	15420613 0215.000	1542-06-13T02:15:00Z	44.006	11.385	6.0	45	Mugello	Guidoboni et al. (2007)
18	15421210 1515.000	1542-12-10T15:15:00Z	37.215	14.944	6.7	26	Sicilia sud-orientale	Guidoboni et al. (2007)
19	15610731 1945.000	1561-07-31T19:45:00Z	40.650	15.389	6.3	21	Vallo di Diano	Castelli et al. (2008)
20	15610819 1550.000	1561-08-19T15:50:00Z	40.563	15.505	6.7	32	Vallo di Diano	Castelli et al. (2008)
21	15991106 0125.000	1599-11-06T01:25:00Z	42.724	13.021	6.1	13	Valnerina	Guidoboni et al. (2007)
22	16260404 1245.000	1626-04-04T12:45:00Z	38.851	16.456	6.1	7	Calabria centrale	Guidoboni et al. (2007)
23	16270730 1050.000	1627-07-30T10:50:00Z	41.737	15.342	6.7	47	Capitanata	Guidoboni et al. (2007)

24	16270807 1640_000	1627-08-07T16:40:00Z	41.758	15.328	6.0	5	Capitanata	Guidoboni et al. (2007)
25	16380327 1505_000	1638-03-27T15:05:00Z	39.048	16.289	7.1	212	Calabria centrale	Guidoboni et al. (2007)
26	16380608 0945_000	1638-06-08T09:45:00Z	39.279	16.812	6.8	41	Crotonese	Guidoboni et al. (2007)
27	16391007_0000_000	1639-10-07T00:00:00Z	42.639	13.261	6.2	30	Monti della Laga	Castelli (2003)
28	16460531_0000_000	1646-05-31T00:00:00Z	41.905	15.993	6.7	28	Gargano	Camassi et al. (2008)
29	16540724 0025_000	1654-07-24T00:00:00Z	41.635	13.683	6.3	37	Sorano	Guidoboni et al. (2007)
30	16591105 2215_000	1659-11-05T22:15:00Z	38.694	16.249	6.6	126	Calabria centrale	Guidoboni et al. (2007)
31	16610322 1250_000	1661-03-22T12:50:00Z	44.021	11.898	6.1	78	Appennino forlivese	Guidoboni et al. (2007)
32	16880605 1530_000	1688-06-05T15:30:00Z	41.283	14.561	7.1	169	Sannio	Guidoboni et al. (2007)
33	16901204 1400_000	1690-12-04T14:00:00Z	46.633	13.880	6.2	57	Carinthia, Villach	Guidoboni et al. (2007)
34	16930109 2100_000	1693-01-09T21:00:00Z	37.141	15.035	6.1	30	Sicilia sud-orientale	Guidoboni et al. (2007)
35	16930111 1330_000	1693-01-11T13:30:00Z	37.140	15.013	7.3	178	Sicilia sud-orientale	Guidoboni et al. (2007)
36	16940908 1140_000	1694-09-08T11:40:00Z	40.862	15.406	6.7	247	Irpinia-Basilicata	Guidoboni et al. (2007)
37	16950225 0530_000	1695-02-25T05:30:00Z	45.861	11.910	6.4	98	Asolano	Camassi et al. (2012)
38	17020314 0500_000	1702-03-14T05:00:00Z	41.120	14.989	6.6	30	Sannio-Irpinia	Guidoboni et al. (2007)
39	17030114 1800_000	1703-01-14T18:00:00Z	42.708	13.071	6.9	187	Valnerina	Guidoboni et al. (2007)
40	17030202 1105_000	1703-02-02T11:05:00Z	42.434	13.292	6.7	67	Aquilano	Guidoboni et al. (2007)
41	17061103 1300_000	1706-11-03T13:00:00Z	42.076	14.080	6.8	96	Maiella	Guidoboni et al. (2007)
42	17300512 0500_000	1730-05-12T05:00:00Z	42.753	13.120	6.0	113	Valnerina	Guidoboni et al. (2007)
43	17310320 0300_000	1731-03-20T03:00:00Z	41.274	15.757	6.3	40	Tavoliere delle Puglie	Guidoboni et al. (2007)
44	17321129 0740_000	1732-11-29T07:40:00Z	41.064	15.059	6.8	182	Irpinia	Guidoboni et al. (2007)
45	17410424 0900_000	1741-04-24T09:20:00Z	43.425	13.005	6.2	135	Fabrianese	Monachesi (1987)
46	17430220 1630_000	1743-02-20T00:00:00Z	39.847	18.774	6.7	72	Ionio settentrionale	Galli and Naso (2008)
47	17470417_0000_000	1747-04-17T00:00:00Z	43.204	12.769	6.1	61	Appennino umbro-marchigiano	Castelli (2003)
48	17510727 0100_000	1751-07-27T01:00:00Z	43.225	12.739	6.4	52	Appennino umbro-marchigiano	Guidoboni et al. (2007)
49	17810404 2120_000	1781-04-04T21:20:00Z	44.251	11.798	6.1	95	Faentino	Guidoboni et al. (2007)
50	17810603_0000_000	1781-06-03T00:00:00Z	43.596	12.512	6.5	142	Cagliese	Monachesi (1987)

51	17830205_1200_000	1783-02-05T12:00:00Z	38.297	15.970	7.1	353	Calabria meridionale	Guidoboni et al. (2007)
52	17830207_1310_000	1783-02-07T13:10:00Z	38.580	16.201	6.7	191	Calabria centrale	Guidoboni et al. (2007)
53	17830328_1855_000	1783-03-28T18:55:00Z	38.785	16.464	7.0	323	Calabria centrale	Guidoboni et al. (2007)
54	17860310_1410_000	1786-03-10T14:10:00Z	38.102	15.021	6.1	10	Golfo di Patti	Guidoboni et al. (2007)
55	17911013_0120_000	1791-10-13T01:20:00Z	38.636	16.268	6.1	75	Calabria centrale	Guidoboni et al. (2007)
56	17990728_2205_000	1799-07-28T22:05:00Z	43.193	13.151	6.2	53	Appennino marchigiano	Guidoboni et al. (2007)
57	18050726_2100_000	1805-07-26T21:00:00Z	41.500	14.474	6.7	208	Molise	Guidoboni et al. (2007)
58	18180220_1815_000	1818-02-20T18:15:00Z	37.603	15.140	6.3	121	Catanese	Guidoboni et al. (2007)
59	18320113_1300_000	1832-01-13T13:00:00Z	42.980	12.605	6.4	91	Valle Umbra	Guidoboni et al. (2007)
60	18320308_1830_000	1832-03-08T18:30:00Z	39.079	16.919	6.7	99	Crotonese	Guidoboni et al. (2007)
61	18360425_0020_000	1836-04-25T00:00:00Z	39.567	16.737	6.2	42	Calabria settentrionale	Guidoboni et al. (2007)
62	18460814_1200_000	1846-08-14T12:00:00Z	43.470	10.562	6.0	103	Colline Pisane	Guidoboni et al. (2007)
63	18510814_1320_000	1851-08-14T13:20:00Z	40.960	15.669	6.5	97	Vulture	Guidoboni et al. (2007)
64	18540212_1750_000	1854-02-12T17:50:00Z	39.256	16.295	6.3	87	Cosentino	Guidoboni et al. (2007)
65	18571216_2115_001	1857-12-16T21:15:00Z	40.352	15.842	7.1	314	Basilicata	Guidoboni et al. (2007)
66	18701004_1655_000	1870-10-04T16:55:00Z	39.220	16.331	6.2	53	Cosentino	Guidoboni et al. (2007)
67	18730629_0358_000	1873-06-29T03:58:00Z	46.159	12.383	6.3	187	Alpago Cansiglio	Guidoboni et al. (2007)
68	18870223_0521_000	1887-02-23T05:21:50.00	43.89	7.992	6.3	1366	Liguria occidentale	Guidoboni et al. (2007)
69	18941116_1752_000	1894-11-16T17:52:00Z	38.288	15.870	6.1	299	Calabria meridionale	Guidoboni et al. (2007)
70	19050908_0143_000	1905-09-08T01:43:00Z	38.811	16.000	7.0	766	Calabria centrale	Galli and Molin (2007)
71	19081228_0420_000	1908-12-28T04:20:27.00	38.14	15.68	7.1	766	Stretto di Messine	Guidoboni et al. (2007)
72	19150113_0652_000	1915-01-13T06:52:43.00	42.01	13.53	7.0	886	Marsica	Molin et al. (1999)
73	19190629_1506_000	1919-06-29T15:06:13.00	43.95	11.48	6.3	484	Mugello	Guidoboni et al. (2007)
74	19200907_0555_000	1920-09-07T05:55:40.00	44.18	10.27	6.5	688	Garfagna	Guidoboni et al. (2007)
75	19280327_0832_000	1928-03-27T08:32:00Z	46.372	12.975	6.0	289	Carnia	Barbano et al. (1990)
76	19300723_0008_000	1930-07-23T00:00:00Z	41.068	15.318	6.7	496	Irpinia	Galli et al. (2002)
77	19361018_0310_000	1936-10-18T03:10:00Z	46.089	12.380	6.1	247	Alpago Cansiglio	Barbano et al. (1986)

78	19620821 1819.000	1962-08-21T18:19:00Z	41.230	14.953	6.2	560	Irpinia	Gizzi (2012)
79	19680115 0201.000	1968-01-15T02:01:9.00Z	37.756	12.981	6.4	161	Valle del Belice	Guidoboni et al. (2007)

Table 2: **Statistical results in terms of median, mean, standard deviation (sd), first (Q1) and third quartiles (Q3) of the residuals. The comparison between the two Shakemap configurations (i.e, using VIPE and Pea08) is shown for the entire dataset and for data grouped according to the EC8 site classes.**

Dataset	median		mean		sd		Q1		Q3	
	VIPE	Pea08	VIPE	Pea08	VIPE	Pea08	VIPE	Pea08	VIPE	Pea08
All data	-0.0048	0.0034	0.0024	0.0075	0.16	0.23	-0.07	-0.12	0.06	0.13
EC8 class A	0.13	0.01	0.17	0.02	0.22	0.24	0.06	-0.11	0.23	0.14
EC8 class B	-0.0018	0.0017	0.002	0.005	0.15	0.24	-0.06	-0.13	0.06	0.13
EC8 class C	-0.040	0.005	-0.035	0.013	0.17	0.23	-0.11	-0.12	0.03	0.14

Table 3: **List of the selected faults from DISS: strike, dip and rake for each source are indicated.**

Event-ID	DISS-ID	Strike(deg)	Dip(deg)	Rake(deg)
11170103.1515.000	ITIS140	248	40	90
16930111.1330.000	ITIS074	57	45	70
16950225.0530.000	ITIS102	240	35	80
17030202.1105.000	ITIS015	132	50	270
17321129.0740.000	ITIS006	275	64	237
17810404.2120.000	ITIS093	108	35	90
17810603.0000.000	ITIS047	134	30	270
17830205.0000.000	ITIS012	30	30	270
17830207.1310.000	ITIS011	30	30	270
18050726.2100.000	ITIS004	304	55	270
18320113.1300.000	ITIS061	330	30	270
19081228.0420.000	ITIS013	30	29	270



19150113.0652.000	ITIS002	135	60	270
19190629.1506.000	ITIS086	298	40	270
19200907.0555.000	ITIS050	305	40	270
19300723.0008.000	ITIS088	280	64	237

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